Space — the ultimate frontier. As our heroic astronauts from the daring crew of the Swine-trek, push back the outer limits of knowledge through extraterres trial exploration, they find wonders beyond belief...

nothing has prepared you for....

* PIGS IN SPACE

Magnetic remote sensing of planetary oceans and the importance of tidal heat in preventing freezing

Robert Tyler

-Planetary Geodynamics Branch,
NASA Goddard Space Flight Center
-Astronomy Department,
University of Maryland College Park

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Case for oceans:

surface features

magnetic data



plumes



heat flux

Observed heat flux 0.1 (W/m²)

Enceladus Flyby

Cassini Returns to the Plumes

August 11, 2008



Magnetic remote sensing of planetary oceans

Two examples:

Induction response of Europa (Khurana, et al., *Nature*, 1998)



Magnetic fields of Earth's ocean tides (Tyler, et al., *Science*, 2003)





Crude inversion of M2 tide (complex amplitude) from magnetic data at satellite altitude



Tide which is

correct \rightarrow

coastline).

Magnetic field of ocean circulation: toroidal component (within ocean)



MAGNETIC FIELD (T); z=-1000 m



Magnetic field of ocean circulation: poloidal component (satellite altitude)

ampl. of magnetic field (tesla) at 500 km altitude, year=10 max 1.602e-09





Back to the case for oceans in Outer Solar System:

surface features

magnetic data



heat flux

Observed heat flux 0.1 (W/m^2)

August 11, 2008



Enceladus Flyby

Cassini Returns to the Plumes

plumes



Central Enigma

Many of these observations can be explained if there are subsurface oceans with a relatively large source of heat

But what is the heat source?

Radiogenic

But what about Enceladus!

- Tidal flexing
- Shear heating in ice shell
- Antifreeze

credits:

NASA, Outer Planets Research Program NASA, Earth Surface and Interiors Program

Tyler R., Strong ocean tidal flow and heating on the outer moons, *Nature (Dec. 13, 2008*)

Bills B., Tidal Flows in Satellite Oceans, Nature Geoscience (Jan. 2009)

Tyler R., Ocean tides heat Enceladus, Geophys. Res. Lett. (2009)

Tyler R., Dynamical considerations suggest that oceans may be common in the universe, J. Cosmology (2010)

Tyler R., Tidal dynamical consideration constrain the state of an ocean on Enceladus, *Icarus (2011)*

Tyler R., Water Worlds and Oceans May be Common in the Universe. *Chapter in book:* The Biological Big Bang: Panspermia and the Origins of Life, edited by C. Wickramasinghe, Cosmology Science Publishers (2010 In Press)

Tyler R., Magnetic remote sensing of Europa's ocean, Icarus (2011)

First contribution:

Ocean tides can easily supply the required heat (for "resonant" ocean configurations)

Second contribution:

These "resonant" ocean configurations are not just possible but appear inevitable (for a satellite ocean attempting to freeze)

Third contribution:

-Not just satellites...

(also smaller and larger planetary bodies under

synchronous or non-synchronous rotation)

-Not just barotropic modes...

(stratified response can be resonantly excited)

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Resonance...









Moons and Men



Enceladus



Isaac Newton

Theory of gravitation (equilibrium tidal response)

Pierre-Simon Laplace

Laplace tidal equations (dynamical tidal response)

Carl-Gustaf Rossby

Rossby waves (important in tidal response)





Previous calculations of the ocean tidal response on the outer moons:

Sagan and Dermott (Titan) --neglected Coriolis forces

Moore and Schubert (Europa, Ganymede, Callisto)

-- assumed equilibrium response (i.e. neglected Coriolis forces *and* ocean inertia)

obeyed Newton

disregarded Laplace

did not consider Rossby at all!







Case for ocean tides as heat source:

- Calculate the tidal forces on the ocean (What is the time-dependent gravitational pull of the planet on the moon's ocean?)
- 2) Calculate the ocean's response to these forces(How does the ocean slosh around as a result?)
- 3) Calculate the dissipation (heat) associated with this response(How much friction generated by flow?)

Gravitational force (gradient of gravitational potential)

$$\Phi = -(GM_J/q_o)(q_o/d)$$

$$\frac{q_o}{d} \approx \{1 + e \cos \Omega t\}
+ \frac{a}{q_o} \{(\cos \phi + 2e \cos(\phi - \Omega t)) \sin \theta + \theta_o ((1 + 2e \cos \Omega t) \cos \theta \cos(\Omega t - \lambda_o)))\}
+ \frac{a^2}{q_o^2} \left\{ \frac{1}{2} (-1 + 3 \cos^2 \phi \sin^2 \theta) + e \left(\frac{3}{2} \cos \Omega t (-1 + 3 \cos^2 \phi \sin^2 \theta) + 6 \cos \phi \sin^2 \theta \sin \phi \sin \Omega t \right) \right\}
+ \frac{a^2}{q_o^2} \left\{ \theta_o 3 \cos \theta \sin \theta \cos \phi \cos(\Omega t - \lambda_o) \right\}
+ \frac{a^2}{q_o^2} \left\{ e \theta_o 3 \cos \theta \sin \theta \cos(\Omega t - \lambda_o) (3 \cos \phi \cos \Omega t + 2 \sin \phi \sin \Omega t) \right\}$$
(9)

Laplace Tidal Equations (with forcing and dissipation term added):

$$\partial_t \mathbf{s} - f \mathbf{s} \times \hat{\mathbf{r}} = -c^2 \nabla (m - m_F) - \alpha \mathbf{s},$$

 $\partial_t m + \nabla \cdot \mathbf{s} = 0,$
 $m = \rho_o \eta$

Helmholtz decomposition:

$$\mathbf{s} = \nabla \Phi + \nabla \times (\Psi \hat{\mathbf{r}})$$

Governing equations:

$$\partial_t \left\{ (\partial_t + \alpha) \nabla^2 \Phi + f \nabla^2 \Psi + 2 \frac{\Omega}{a^2} \partial_\phi \Phi + 2 \frac{\Omega}{a^2} \sin^2 \theta \frac{\partial \Psi}{\partial \cos \theta} \right\} - c^2 \nabla^4 \Phi = \partial_t \left\{ c^2 \nabla^2 m_F \right\} + \partial_t R_F$$

$$\left(\partial_t + \alpha\right) \nabla^2 \Psi - f \nabla^2 \Phi + 2 \frac{\Omega}{a^2} \partial_\phi \Psi - 2 \frac{\Omega}{a^2} \sin^2 \theta \frac{\partial \Phi}{\partial \cos \theta} = R_2,$$

Solution method:

Solve using spherical-harmonic expansion of variables, and numerical inversion of resulting coefficient matrix

Dissipation (heating rate) calculated from an energy equation:



Calculating the Ocean Tidal Response



Dissipative heating rate (Q,h)



Caveats in this discussion:

- Not all (Q, h) combinations allowed
- Log10 scales
- Large axes range:
 - $Q \rightarrow$ damping by thick ice
 - $h \rightarrow$ "equivalent depth"
- Small contour plot range 0.001 1 (W/m²)
- Obliquity usually unknown (use 0.1 deg ref.)



Europa

Ganymede

Callisto

log₁₀ ecc. tidal work (W/m²); max=0.28685 \log_{10} ecc. tidal work (W/m²); max=0.75501 \log_{10} ecc. tidal work (W/m²); max=164.0059 10⁸ 10⁷ 10⁸ -2 -1 -3 0 -2 -2 -1 -1 -3 0 0 10⁷ 10⁶ 10⁷ equilibrium tide valid 10⁶ 10⁵ 10⁶ 10⁵ 10⁴ 10⁵ ROSON 10⁴ 0 10³ 10⁴ damped cteepine now ε 10³ μ $\widehat{\underbrace{E}}_{\mathbf{F}} 10^2$ Ê 10³ 10² gravity 10² 10¹ reso 10¹ 10⁰ 10¹ 10⁰ 10⁰ 10⁻¹ 10⁻¹ 10⁻² 10⁻¹ radio=0.0043466 radio=0.0052506 radio=0.096642 10⁻² 10⁻² 10⁻³ flex=0.0031609 flex=0.0021693 flex=0.08368 10⁻³ 10⁻³ 10^{-4} 10¹ 10⁻³ 10² 10³ 10⁻² 10⁻¹ 10⁻³ 10⁰ 10^{-2} 10^{-1} 10^{0} 10^{1} 10² 10³ 10⁴ 10^{-3} 10^{-2} 10^{-1} 10⁰ 10² 10³ 10⁴ 10⁴ 10¹ Q Q Q



Ocean tidal response depends sensitively on ocean thickness h (m), suggesting h can be inferred from observations of response. \log_{10} ecc. tidal work (W/m²); max=139.2312





eccentricity tidal flow ellipses and dissipative heat (W/m

0.3 0.4

0.1 0.2 0.6

0.7 0.8 0.9

5.5 x 10

Result of this research:

It is theoretically impossible to freeze an idealized global liquid ocean on a synchronously rotating moon subject to appreciable tidal forces

Combine with the following:

synchronously rotating moons are probably ubiquitous

• Liquid water may be a common feature of planetary bodies, at least primordially or episodically

• The time scale for ocean tides to extract available energy from the orbit is very long

General Conclusion:

Liquid oceans may be very common in the universe

Importance to Astrobiology and NASA: Follow the water!