

Be the professor! First circle **True** or **False** for each statement below (there may be several correct statements or none at all). Then, for all **False** statements, cross out the part of the sentence that is wrong. Finally, for Extra Credit, provide a one or two word correction for the **False** statements, if possible.

Chapter 5: Planetary Atmospheres

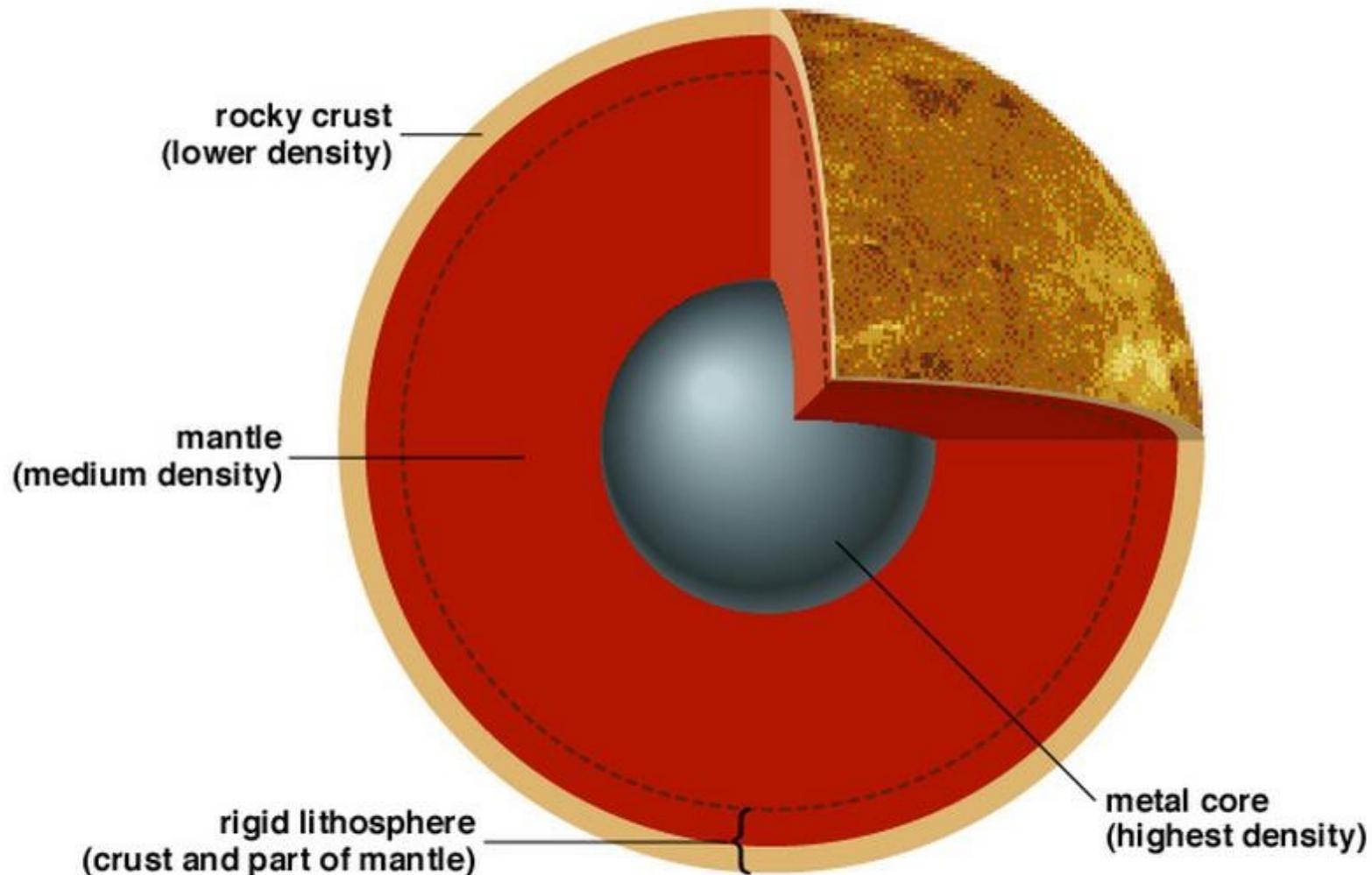
- 3 True or False The exponential change in atmospheric pressure with height derived in class also applies to the interiors of planets. *gas giant +1*
- 3 True or False Ozone absorbs ultraviolet light and is the primary molecule responsible for heating Earth's ~~Mesosphere~~ *stratosphere +1*
- 3 True or False Water clouds are found on the following planets: ~~Mercury, Venus,~~ Earth, Mars, Jupiter, and Saturn.
- 3 True or False Carbon Dioxide clouds are found on the following planets: ~~Mercury, Venus,~~ Earth, Mars, ~~Jupiter,~~ and ~~Saturn.~~
- 3 True or ~~False~~ Gas molecules high in a planetary atmosphere have a Maxwellian distribution of speeds; the fastest ones undergo Thermal, or Jeans, escape.
- 3 True or False Giant planet atmospheres are composed primarily of H and He; other abundant gases include H_2O , ~~O_2 , CO_2 ,~~ H_2S , CH_4 , and NH_3 .

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Planetary Interiors

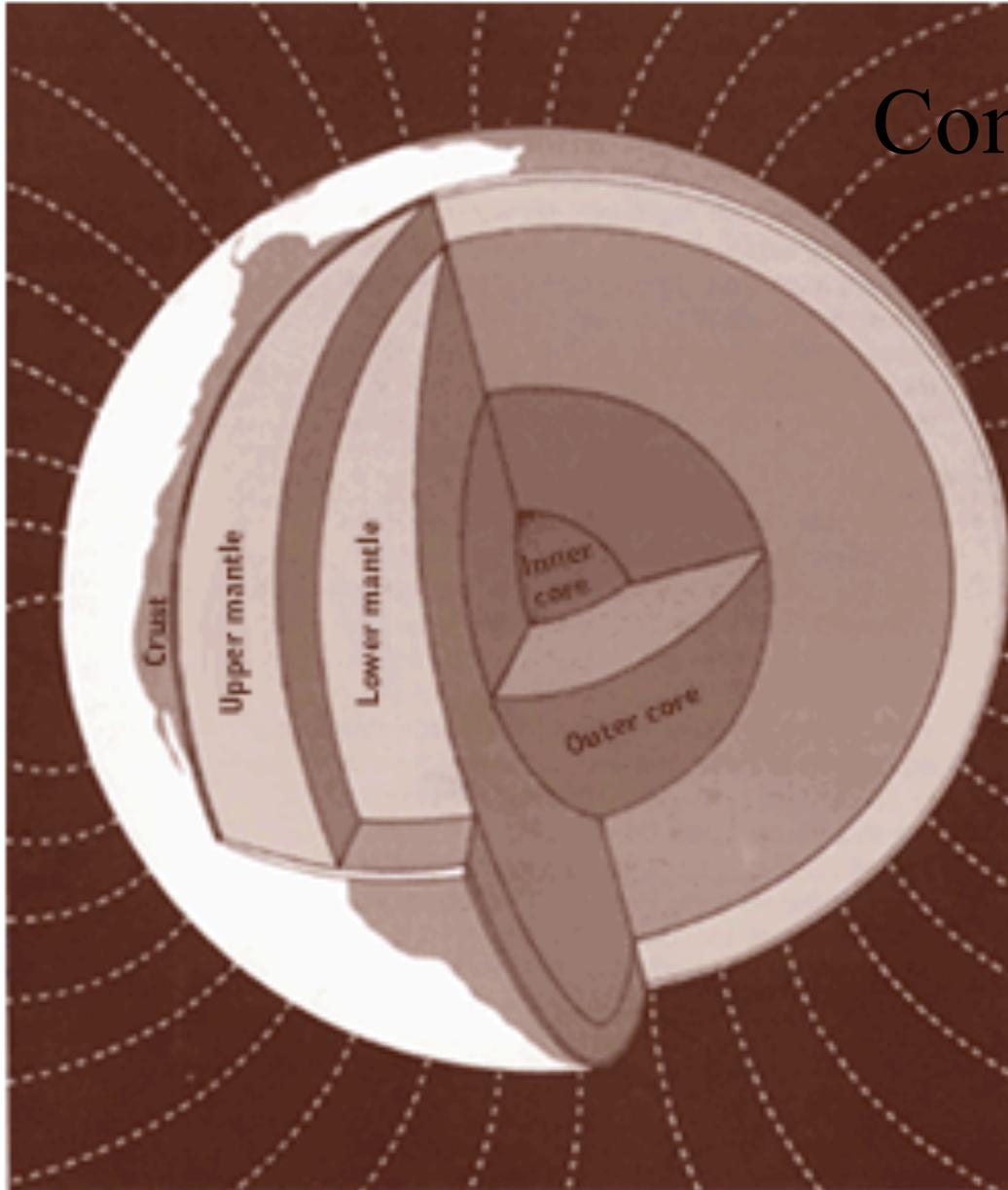
1. Interior Structure of the Earth
2. Interiors of Other Planetary Bodies
3. Accretion and Radioactive Heating
4. Plate Tectonics

Earth's Interior



The Interior of the Earth

by Eugene C. Robertson



Components:

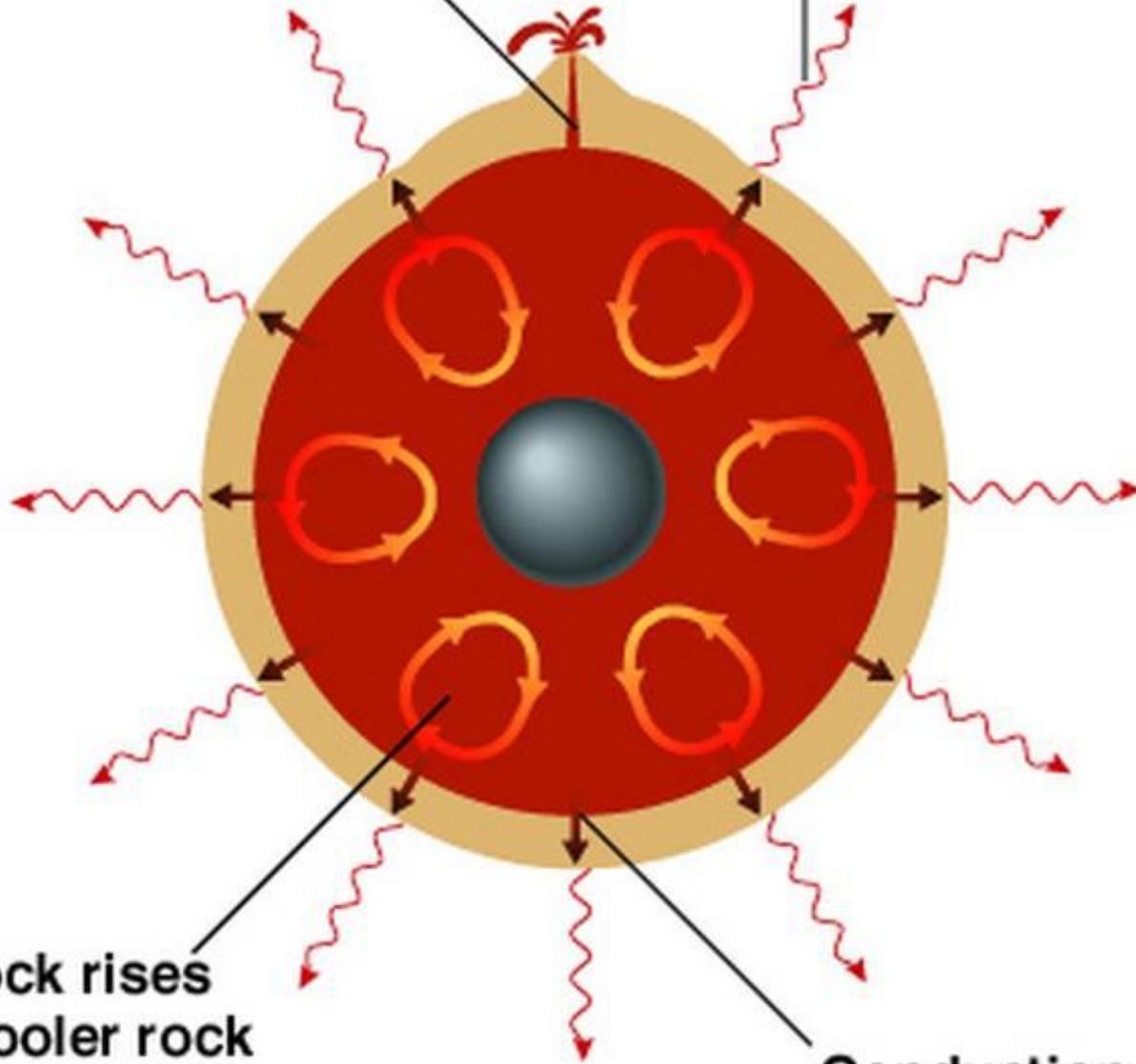
Crust

Mantle

Core

Eruptions bring hot lava up from the interior.

Energy is radiated into space.



Hot rock rises and cooler rock falls in a mantle convection cell.

Conduction carries heat in the rigid lithosphere.

Earthquake Waves

1. P Waves and S Waves

2. Planetary Size vs. Cooling Rate

Crustal and mantle materials. This boundary is known as the Moho or 'Moho' (Box 2.4). This important seismic boundary provides excellent evidence for a fundamental compositional layering structure within our planet.

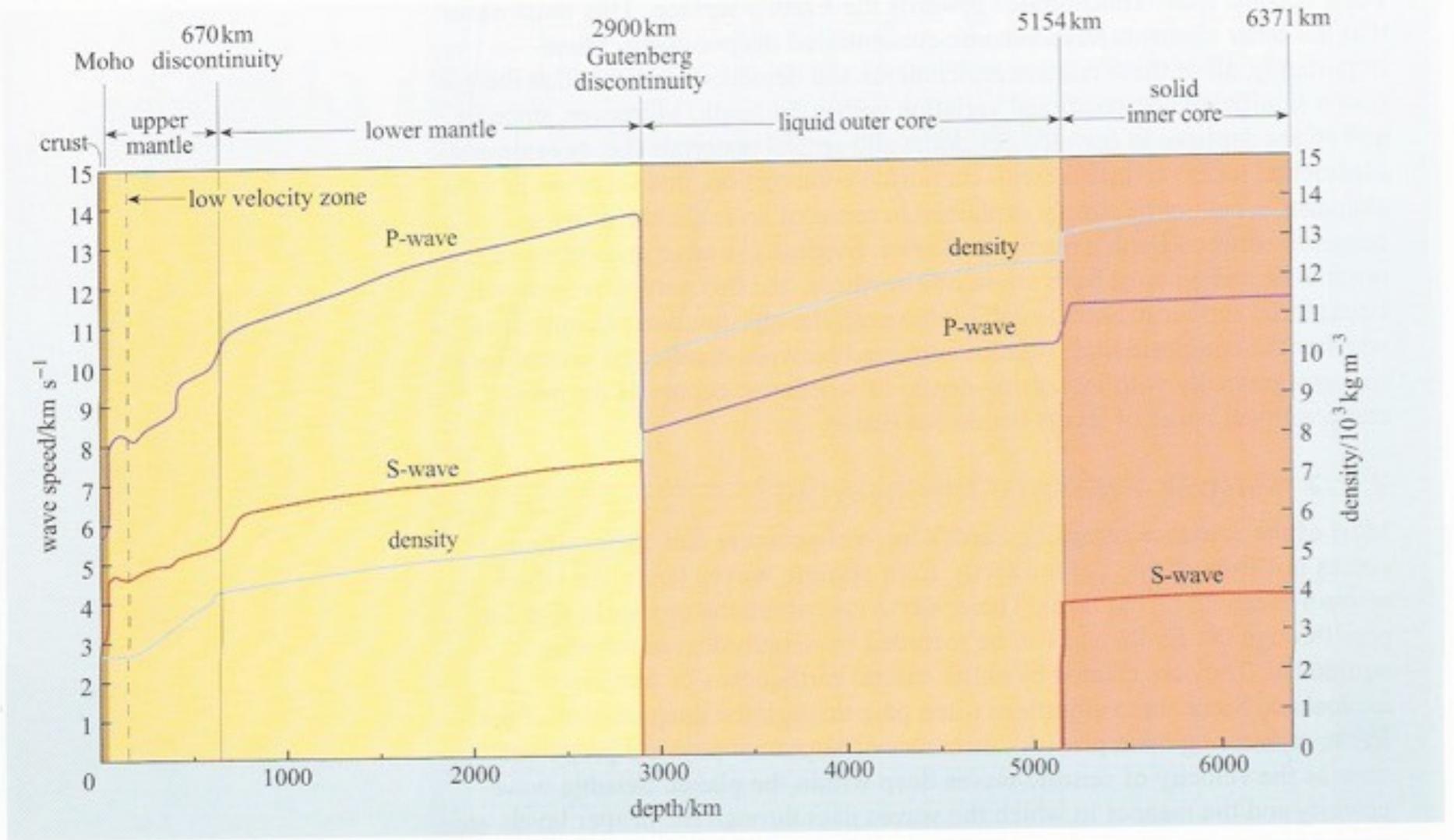
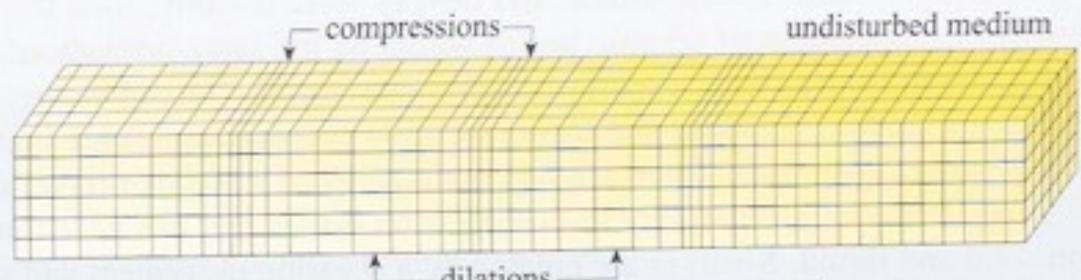
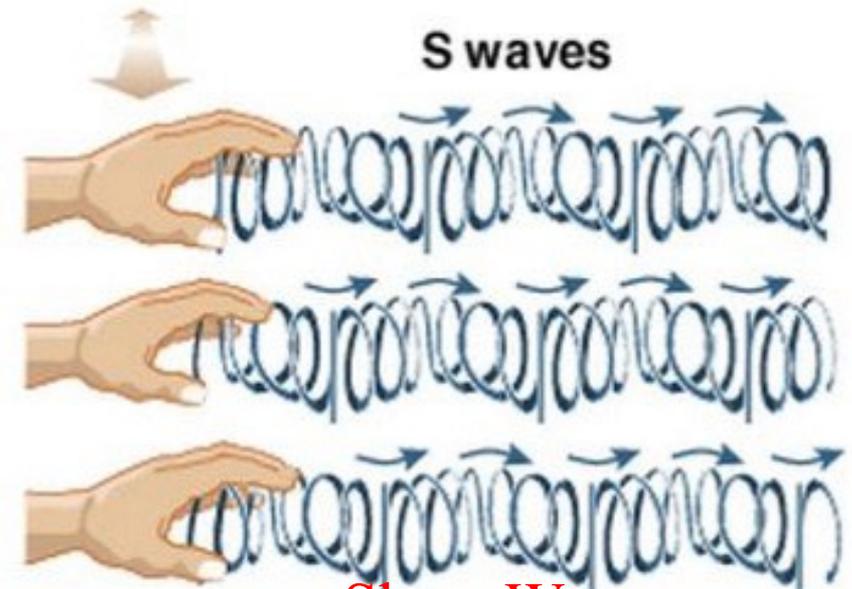
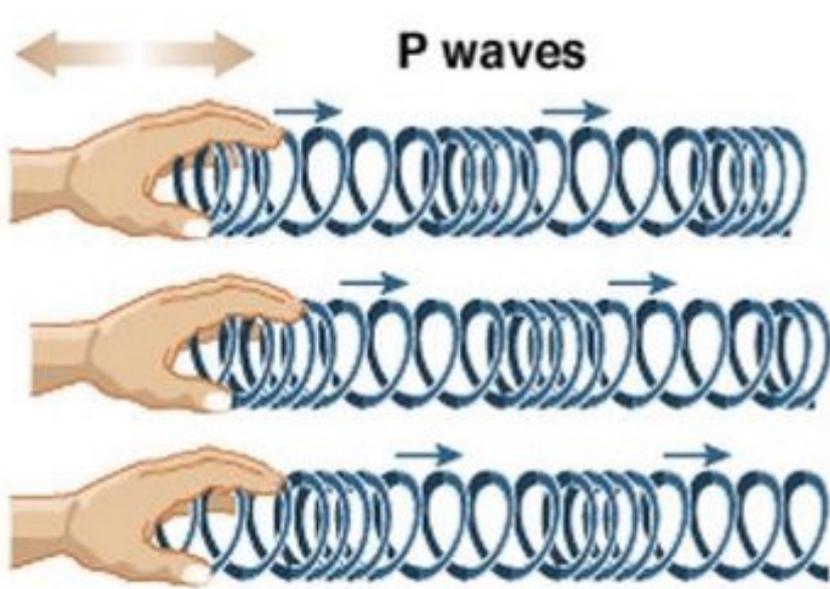


Figure 2.4 Velocity profiles of P-waves and S-waves within the Earth, and inferred densities. The term 'velocity profile' refers to the changes in velocity of seismic waves with increasing depth.

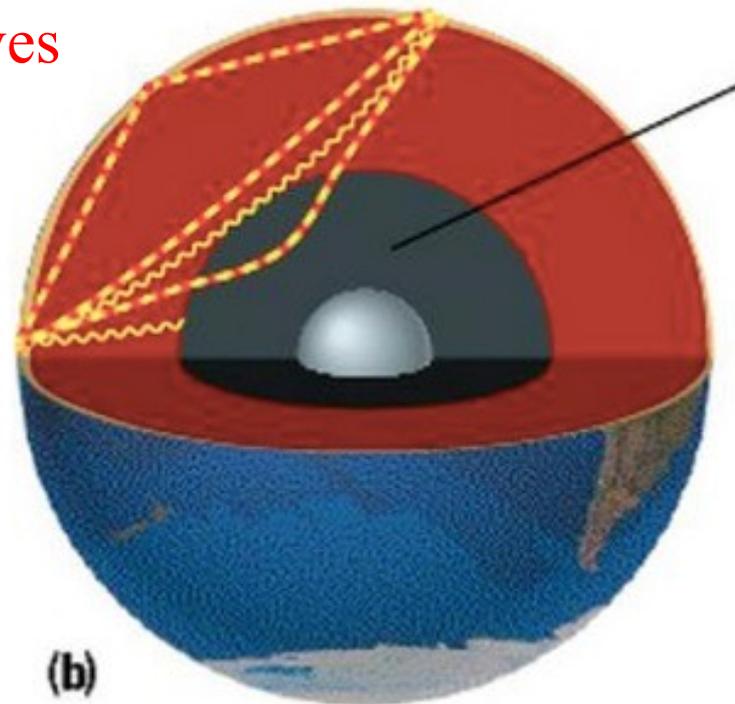


P Waves and S Waves



(a) Pressure Waves
Primary Waves

Shear Waves
Secondary Waves



Molten outer core
stops S waves,
bends P waves.

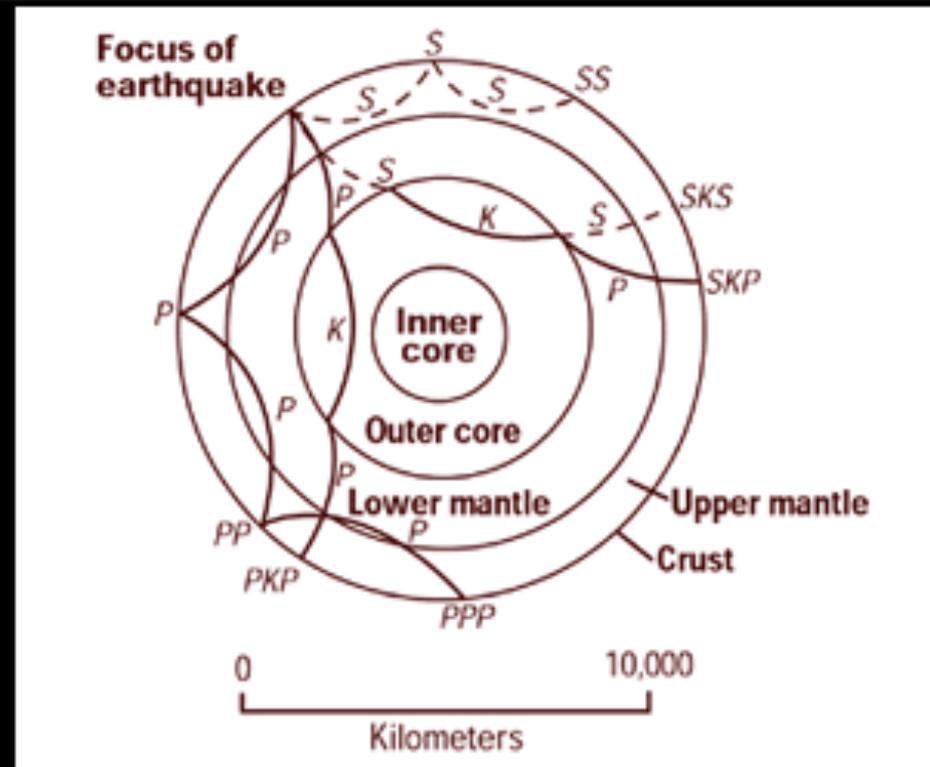
(b)

How do we know this?

Data on the Earth's Interior

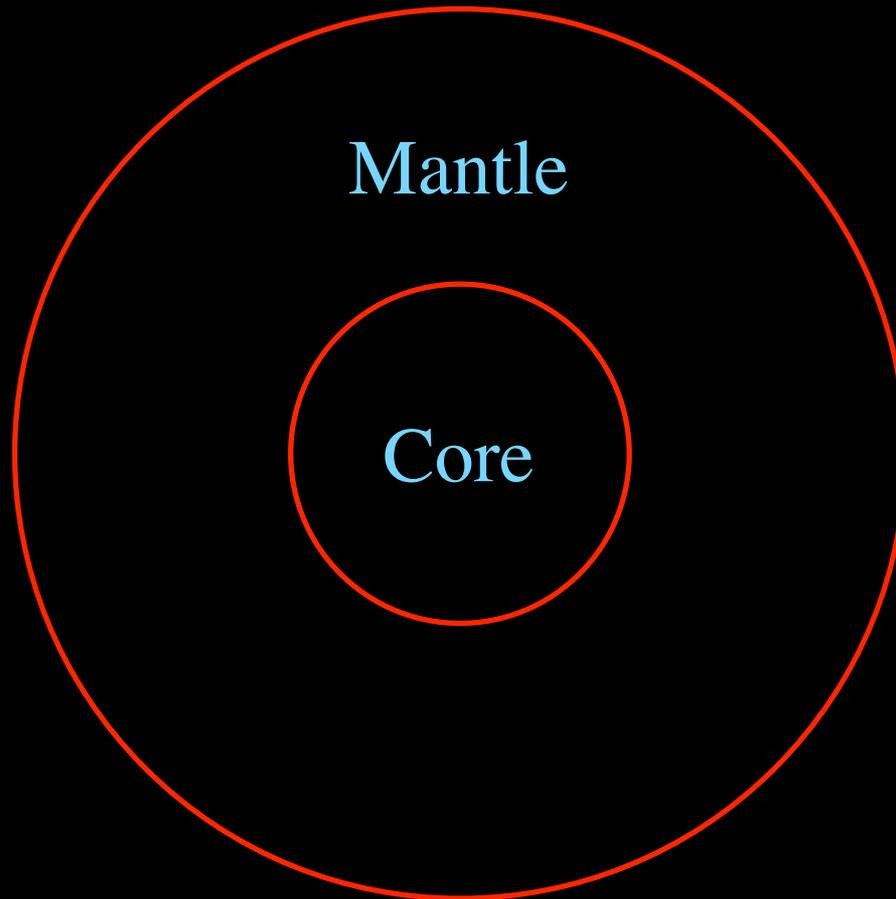
	<u>Thickness (km)</u>	<u>Density (g/cm³)</u>		<u>Types of rock found</u>
		<u>Top</u>	<u>Bottom</u>	
Crust	30	2.2 —	— 2.9	Silicic rocks. Andesite, basalt at base.
Upper mantle	720	3.4 —	— 4.4	Peridotite, eclogite, olivine, spinel, garnet, pyroxene. Perovskite, oxides.
Lower mantle	2,171	4.4 —	— 5.6	Magnesium and silicon oxides.
Outer core	2,259	9.9 —	— 12.2	Iron+oxygen, sulfur, nickel alloy.
Inner core	1,221	12.8 —	— 13.1	Iron+oxygen, sulfur, nickel alloy.
Total thickness	6,401			

Our accurate thickness and composition data comes mainly from monitoring earthquakes

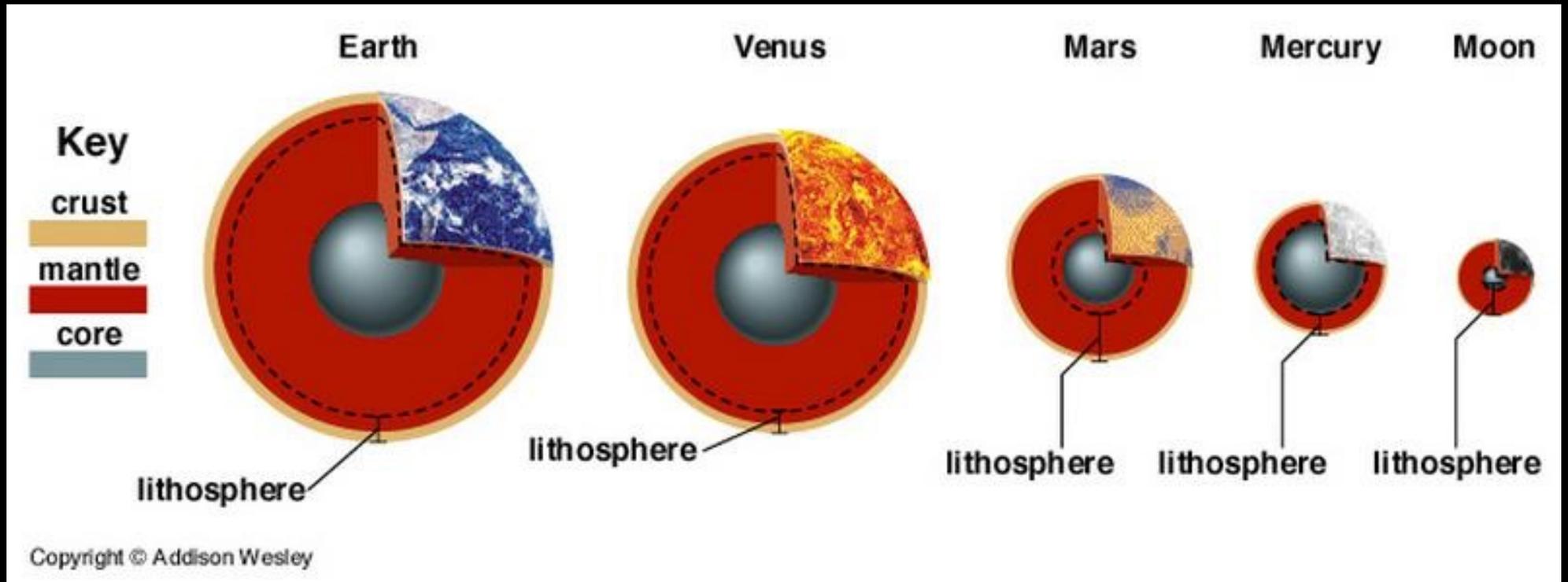


Dealing with Core/Mantle

$$V_{mantle} = \frac{4}{3}\pi r_{outer}^3 - \frac{4}{3}\pi r_{inner}^3$$

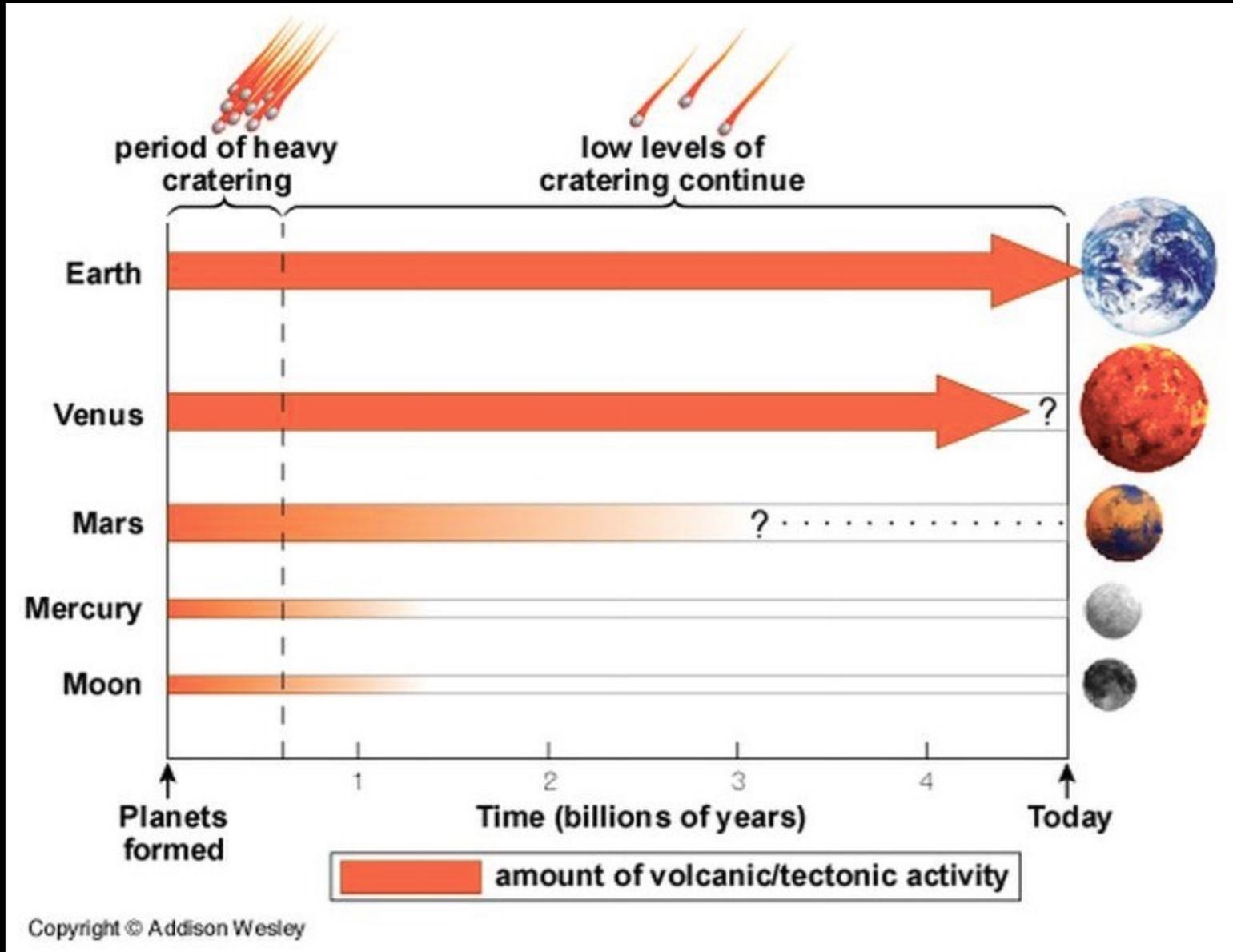


Compositions of the Terrestrial Bodies

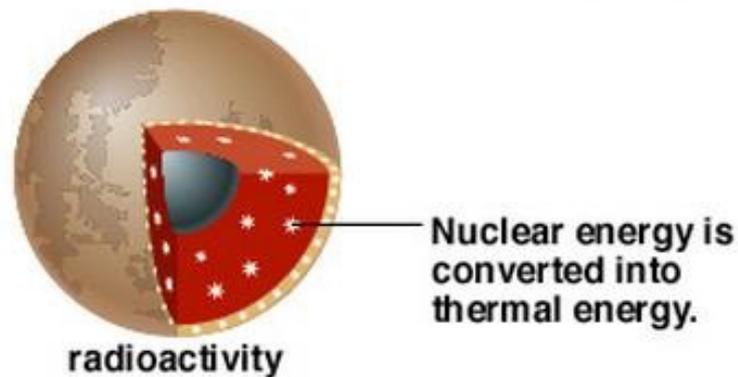
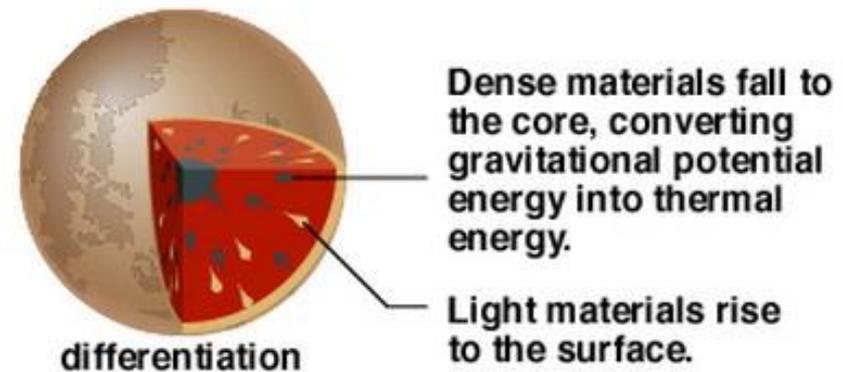
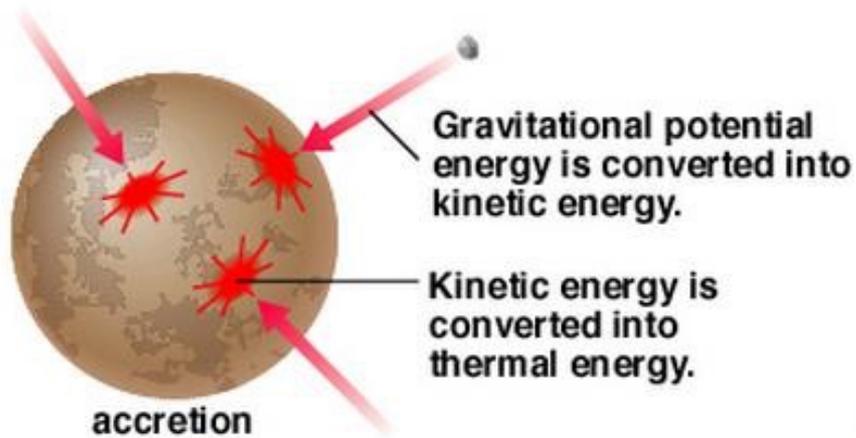


How do we know this?

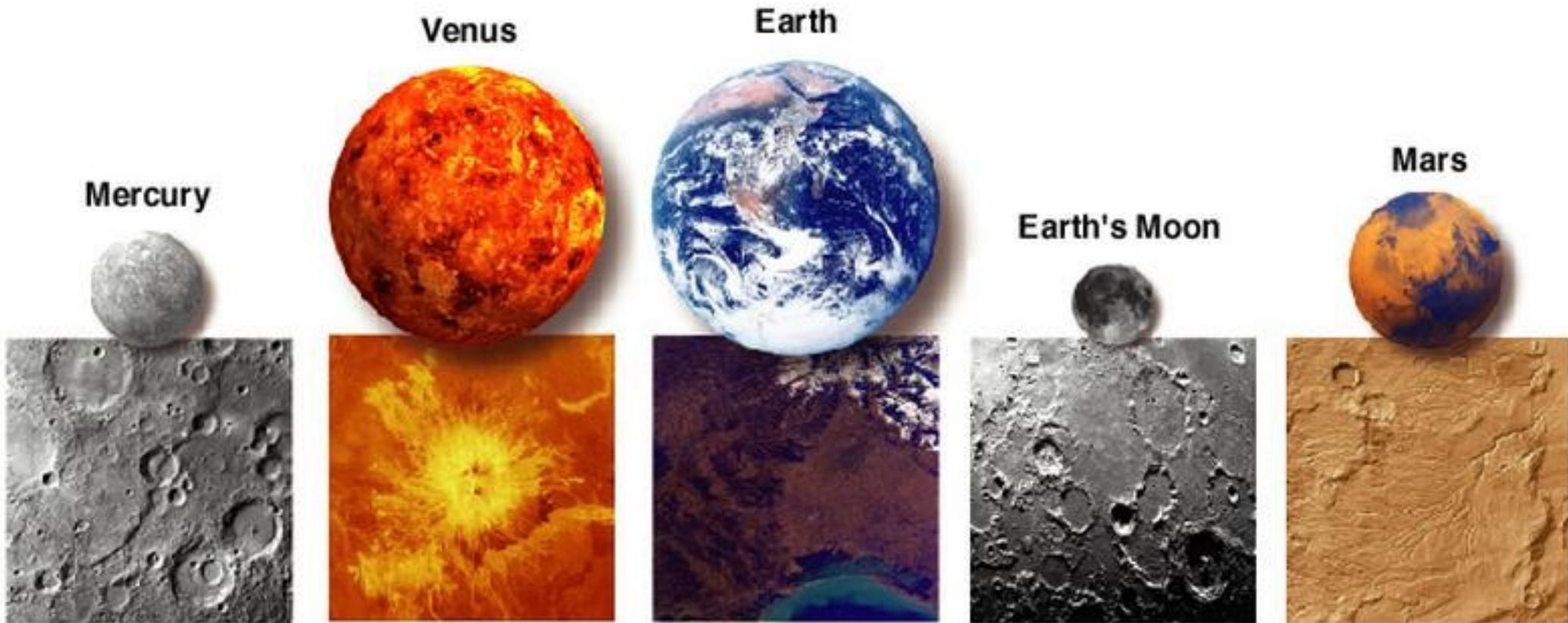
A Planet's Size Correlates with its Tectonic Activity



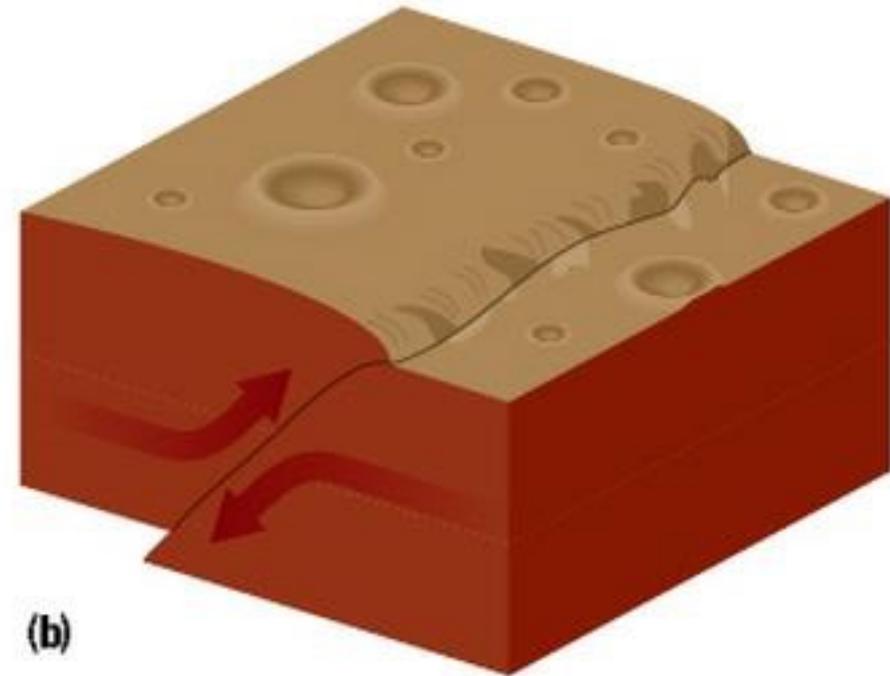
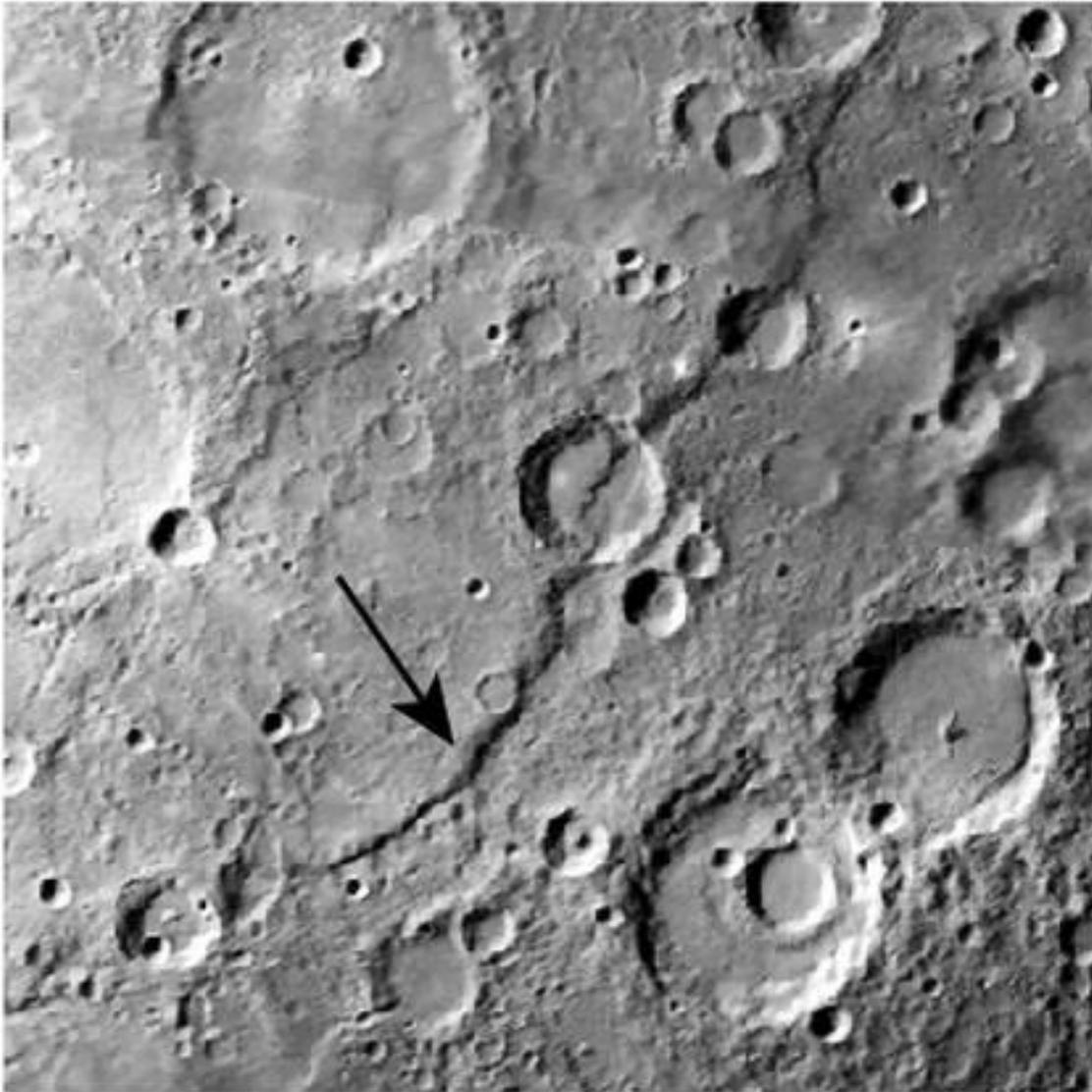
Processes leading to core formation



Planetary Surfaces



Contraction of Mercury



(b)

(a)

Planetary Surfaces



(b)



(d)



(e)

Question:

What evidence do we have to support Plate Tectonics?

Question:

What evidence do we have to support Plate Tectonics?

1. Mid Atlantic ridge, new seafloor, hydrothermal vents
2. Subduction zones
3. Fossils
4. Locations of volcanos, ring of fire + hotspots
5. Earthquakes
6. Mountain ranges

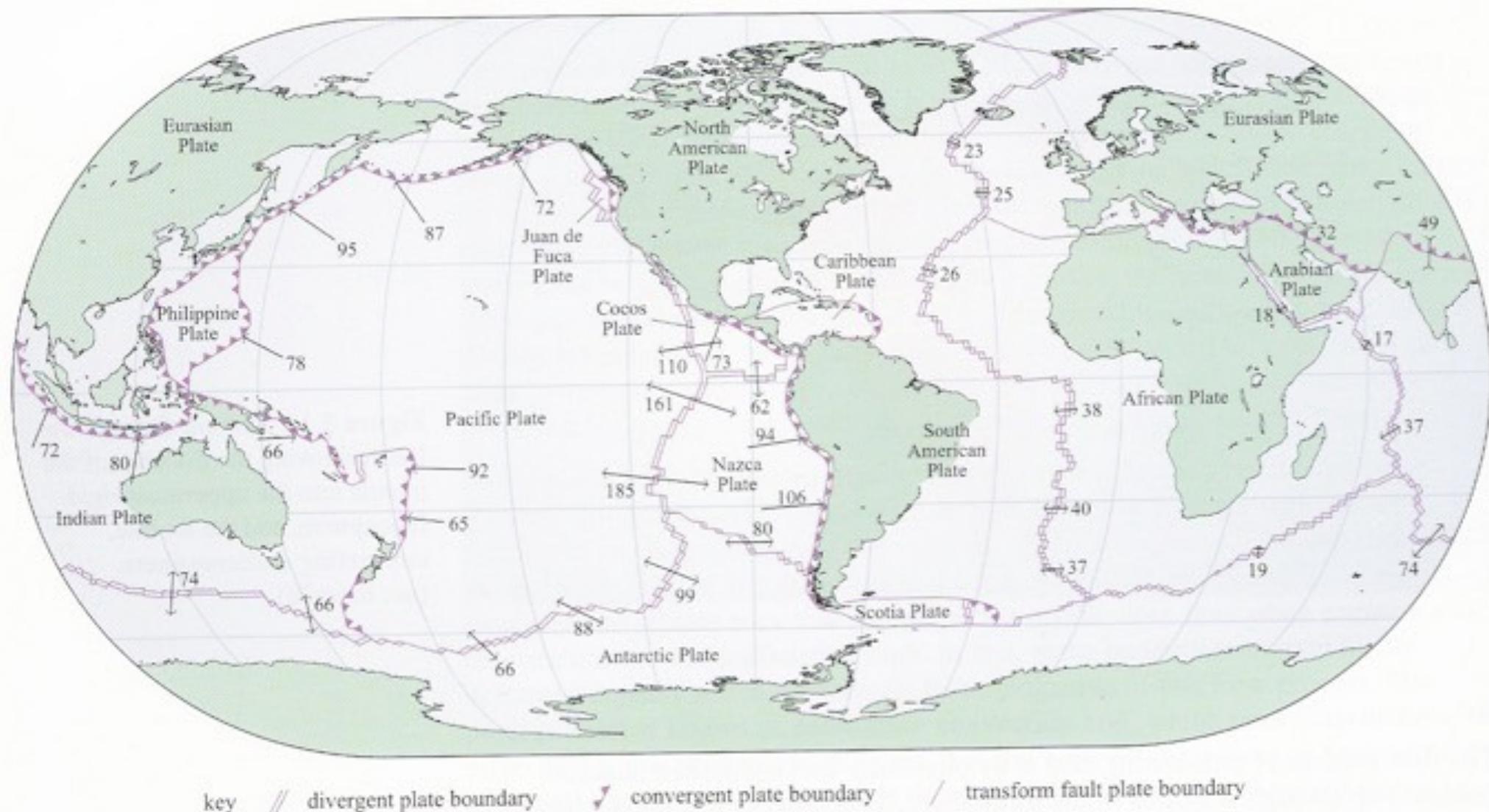


Figure 2.15 Map showing the global distribution of plates and plate boundaries. The black arrows and numbers give the direction and speed of relative motion between plates. Speeds of motion are given in mm yr^{-1} .

Crustal and mantle materials. This boundary is known as the Moho or 'Moho' (Box 2.4). This important seismic boundary provides excellent evidence for a fundamental compositional layering structure within our planet.

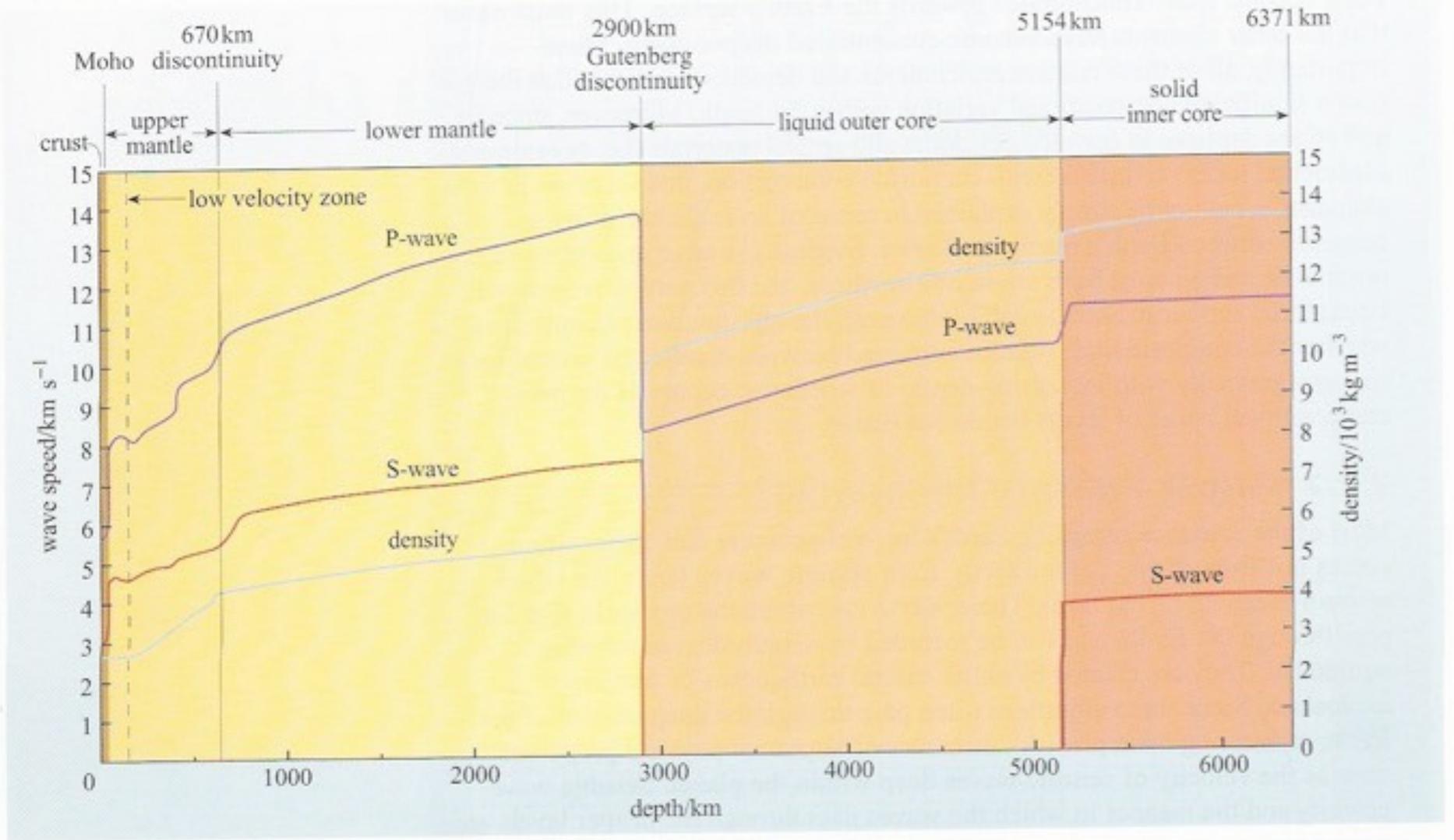
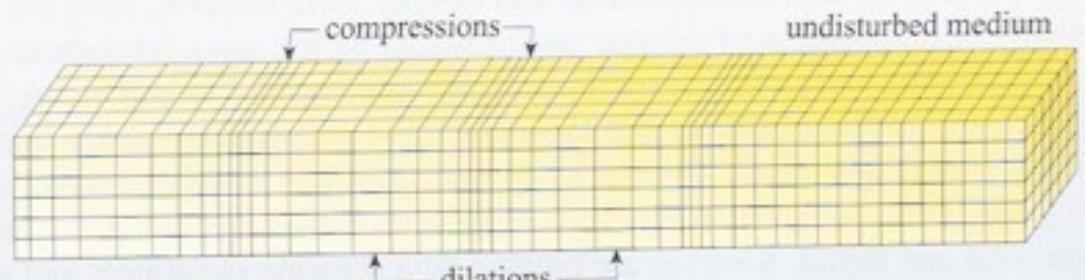
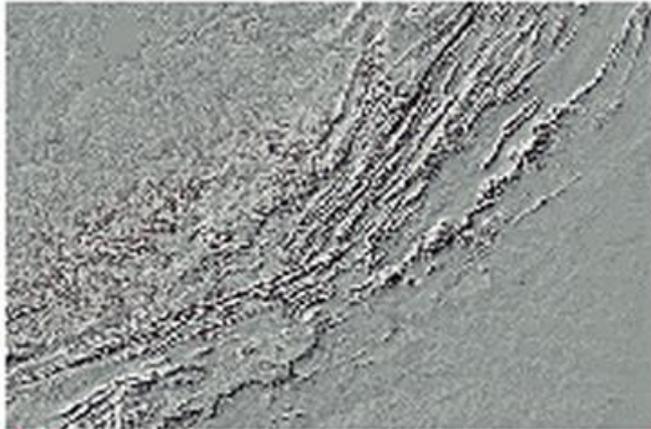


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Features due to Plate Motions



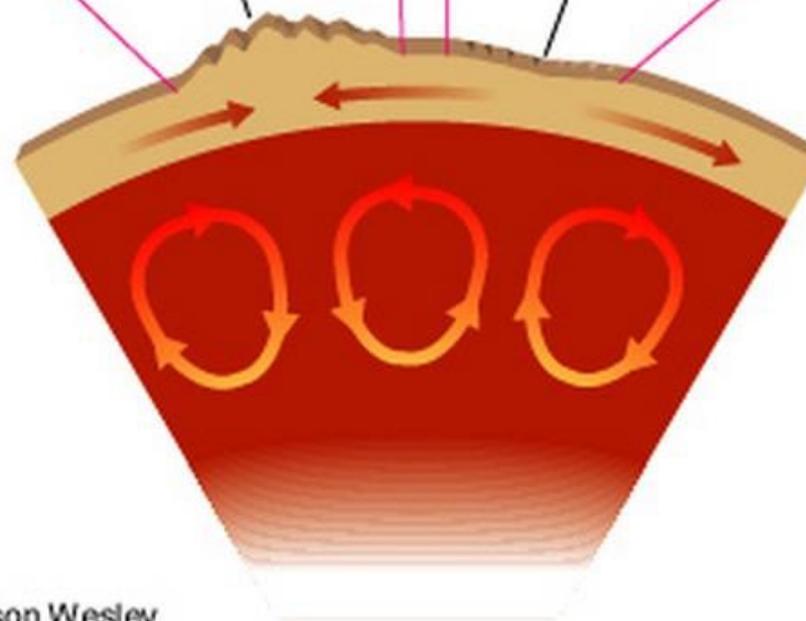
Appalachian Mountains in eastern United States



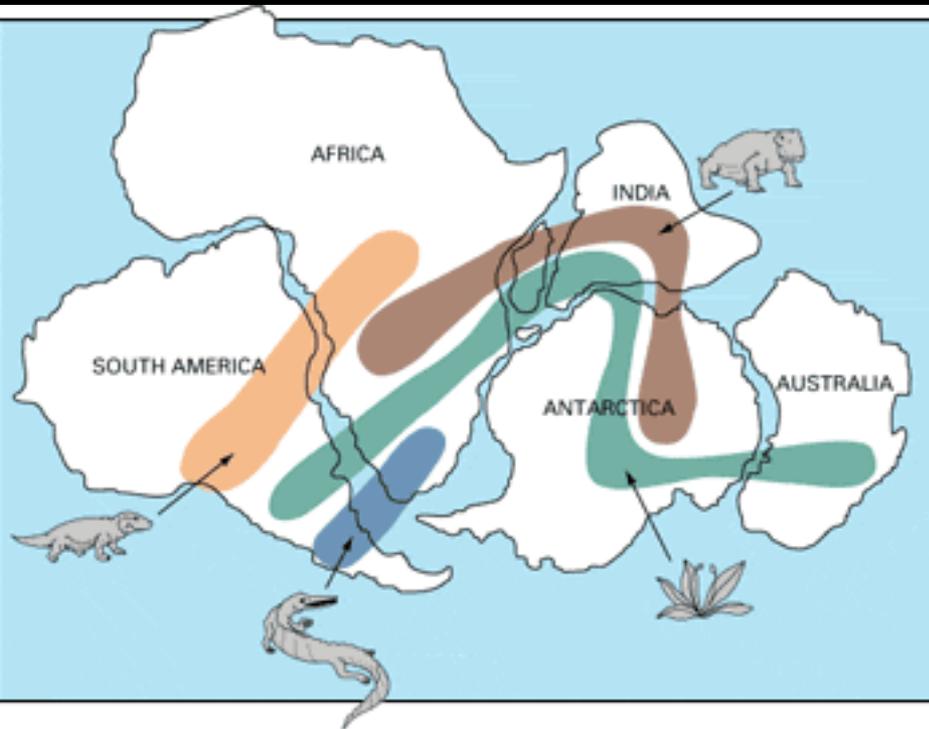
Guinevere Plains on Venus

Compression in crust can make mountains.

Extension can make cracks and valleys.



Evidence for Plate Tectonics



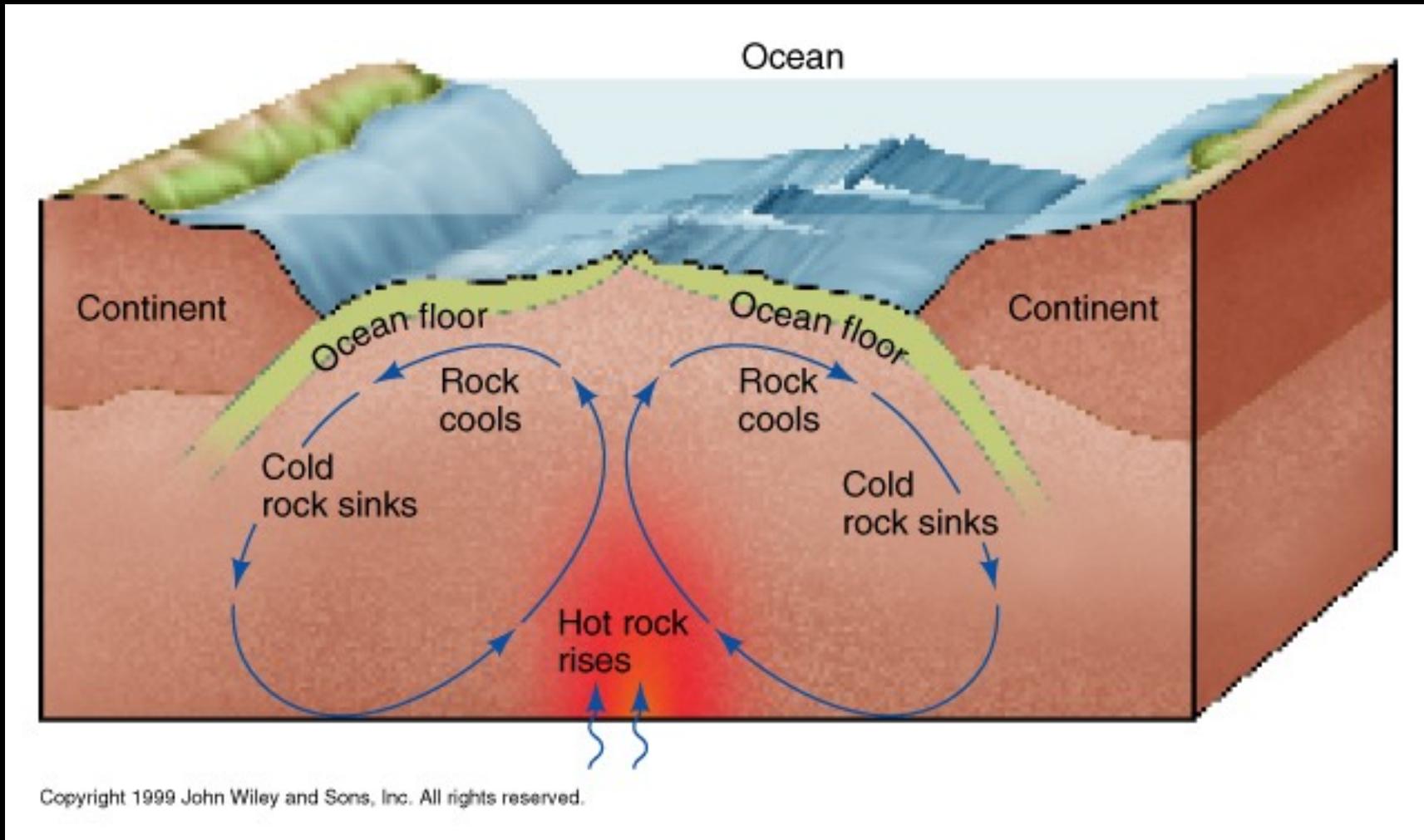
Shape Evidence: The continents look like they can be fit together like a jigsaw puzzle

Biological Evidence: The similarities between some types of ancient plants and animals can be best explained by continental motions.

Glacial Evidence: the white parts of each continent had glacial epochs at around the same time.

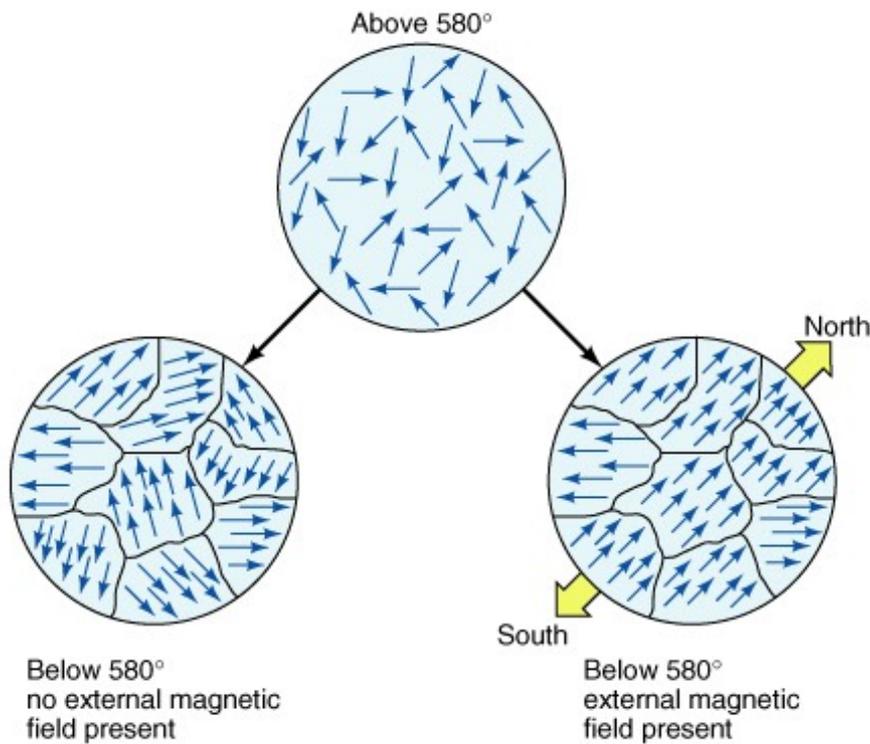


Evidence for Plate Tectonics



Seafloor Spreading: The Atlantic ocean is spreading as new molten material is raised from depth. This explains why the seafloor rock is basalt, a volcanic rock.

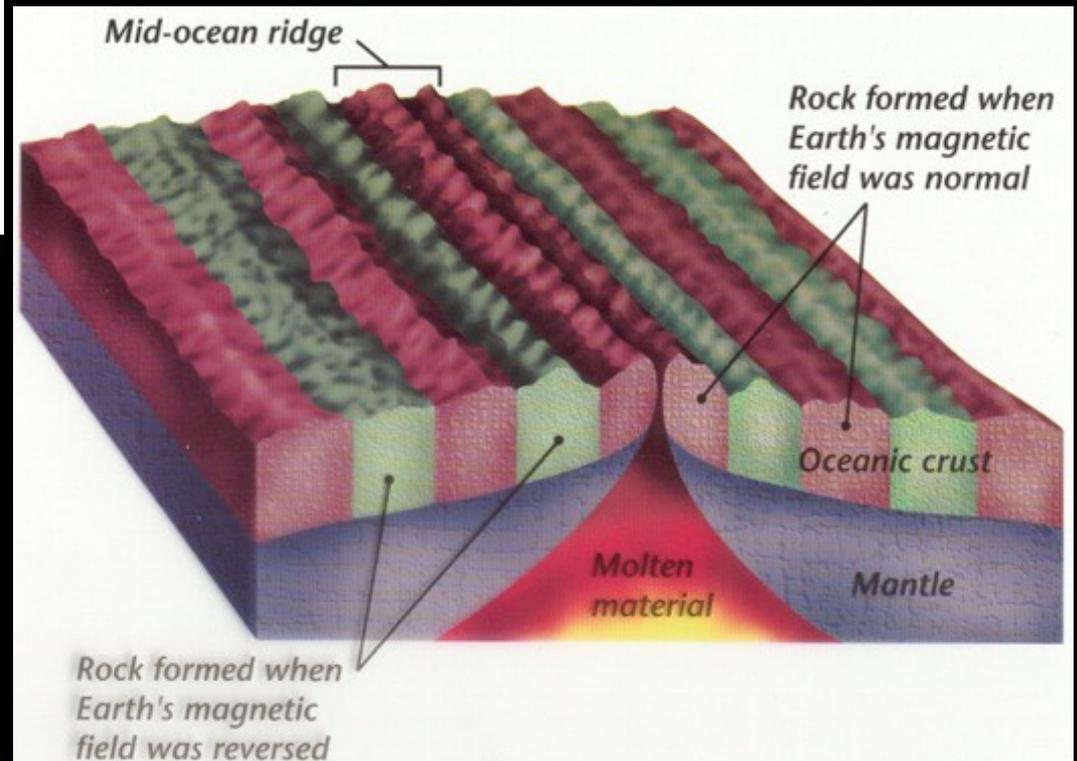
Evidence for Plate Tectonics



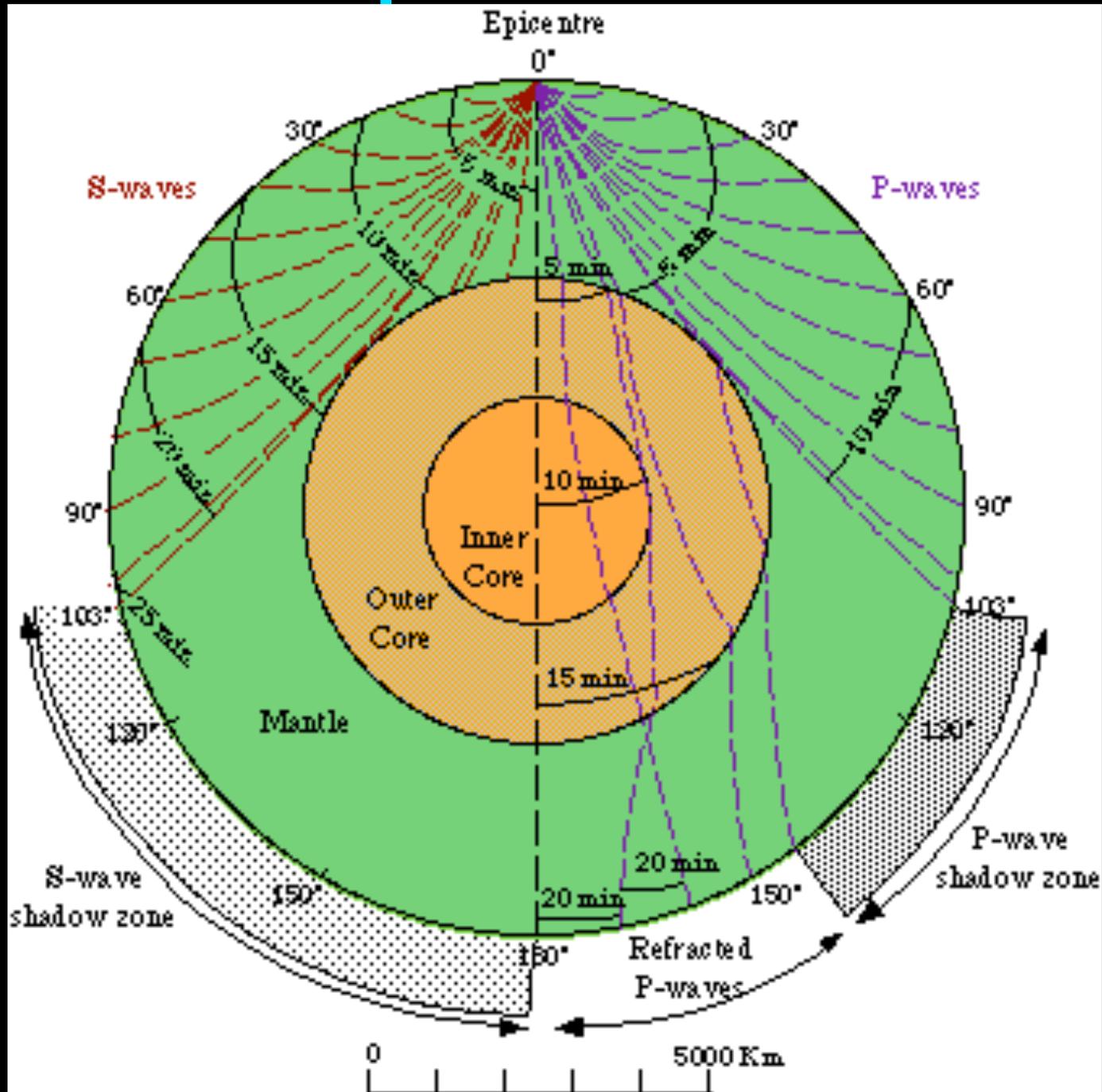
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Rock Magnetic Fields: The alignment of magnetic grains in rock depends on the orientation of the Earth's magnetic field

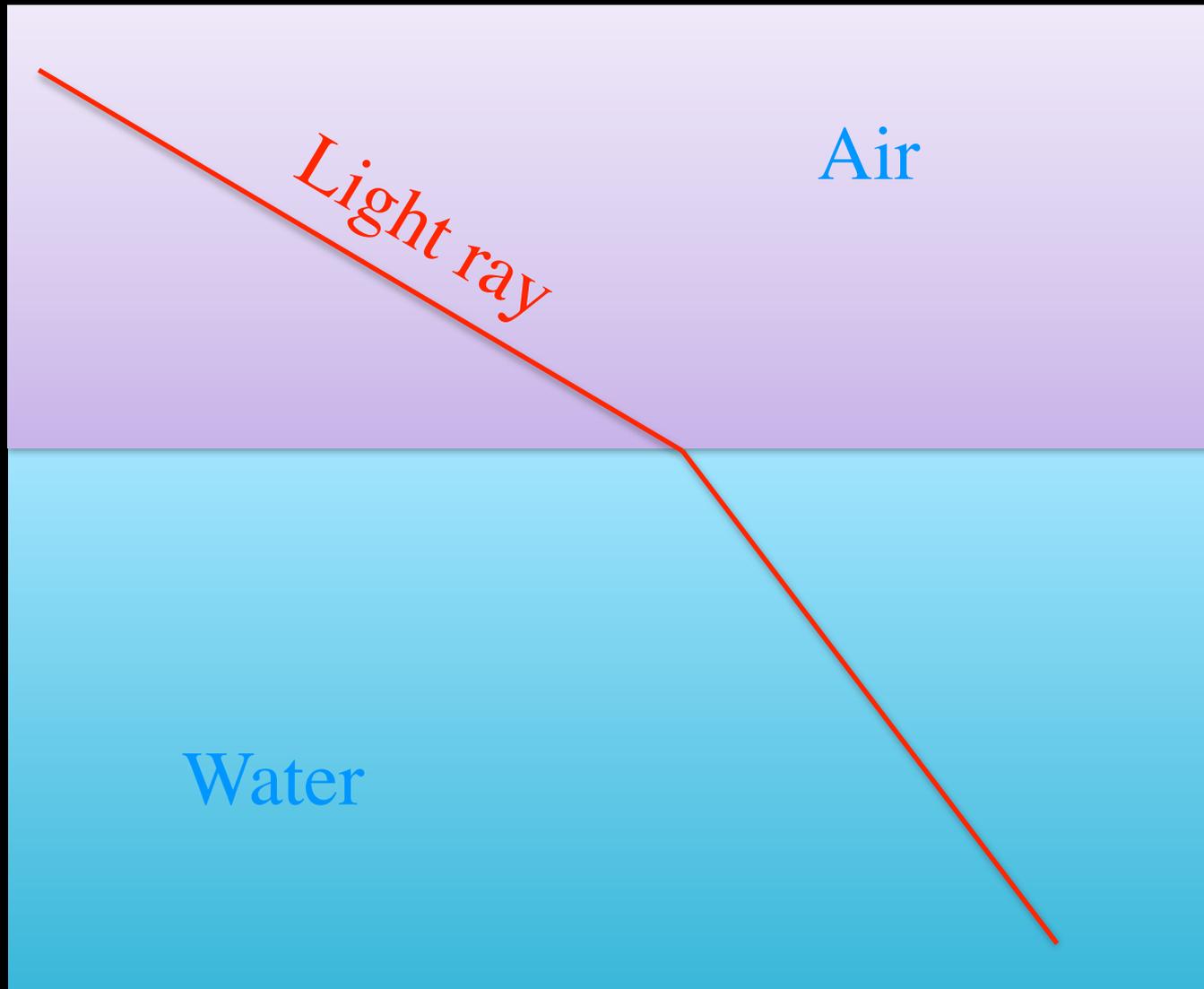
Seafloor Evidence: The rocks in the Atlantic ocean display bands of differently-oriented magnetic grains implying a “conveyer belt” motion



Earthquake Waves



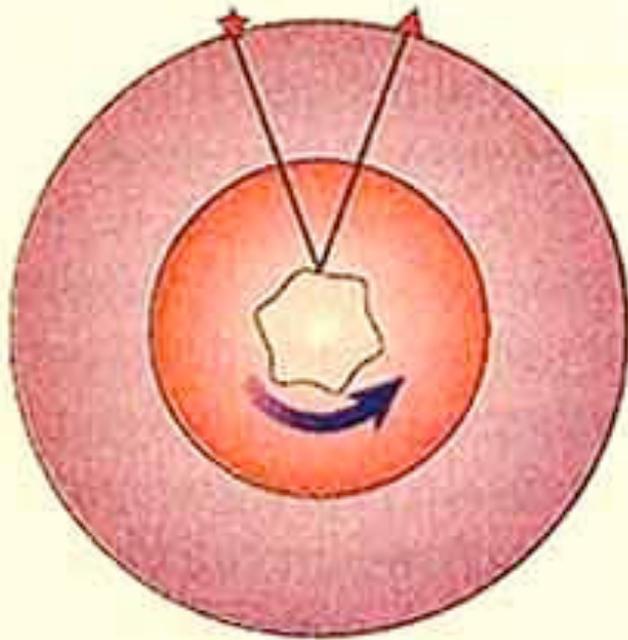
Refraction of Waves



Detecting the Spin of the Core

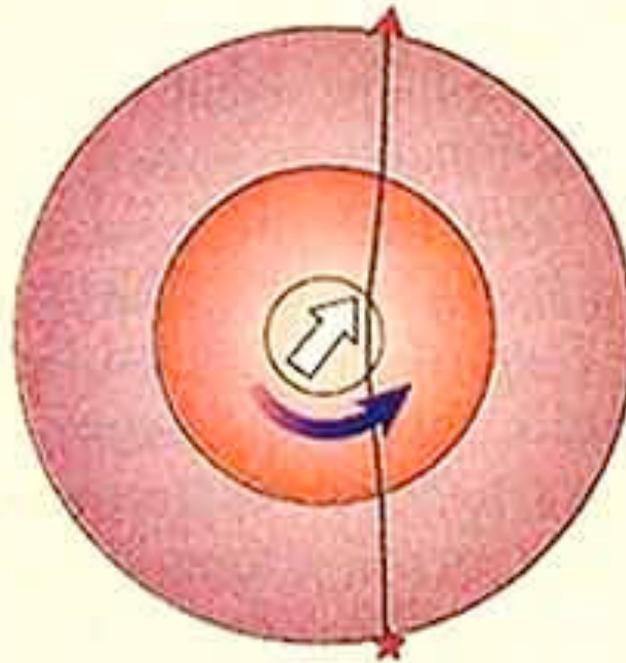
a

Seismic wave reflections from the inner-core surface



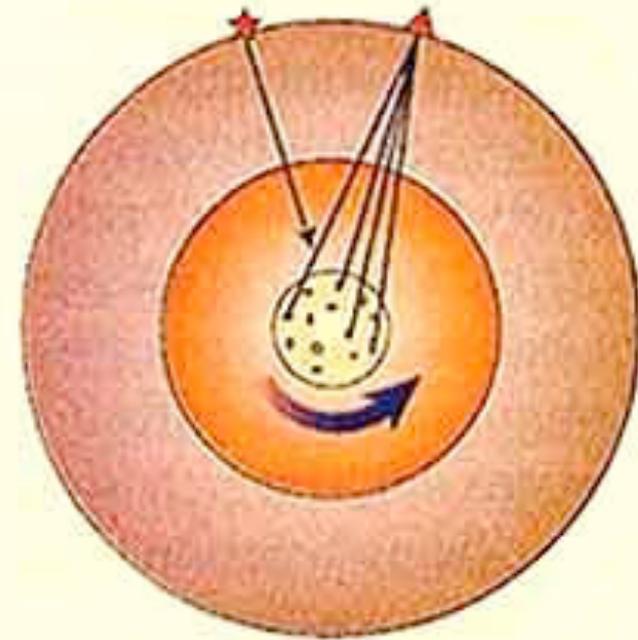
b

Seismic wave travel through the inner core



c

Seismic-wave scattering by the inner core



Two Kinds of Volcanic Activity on Earth

1. Along Plate Boundaries
2. Mantle Hot Spots

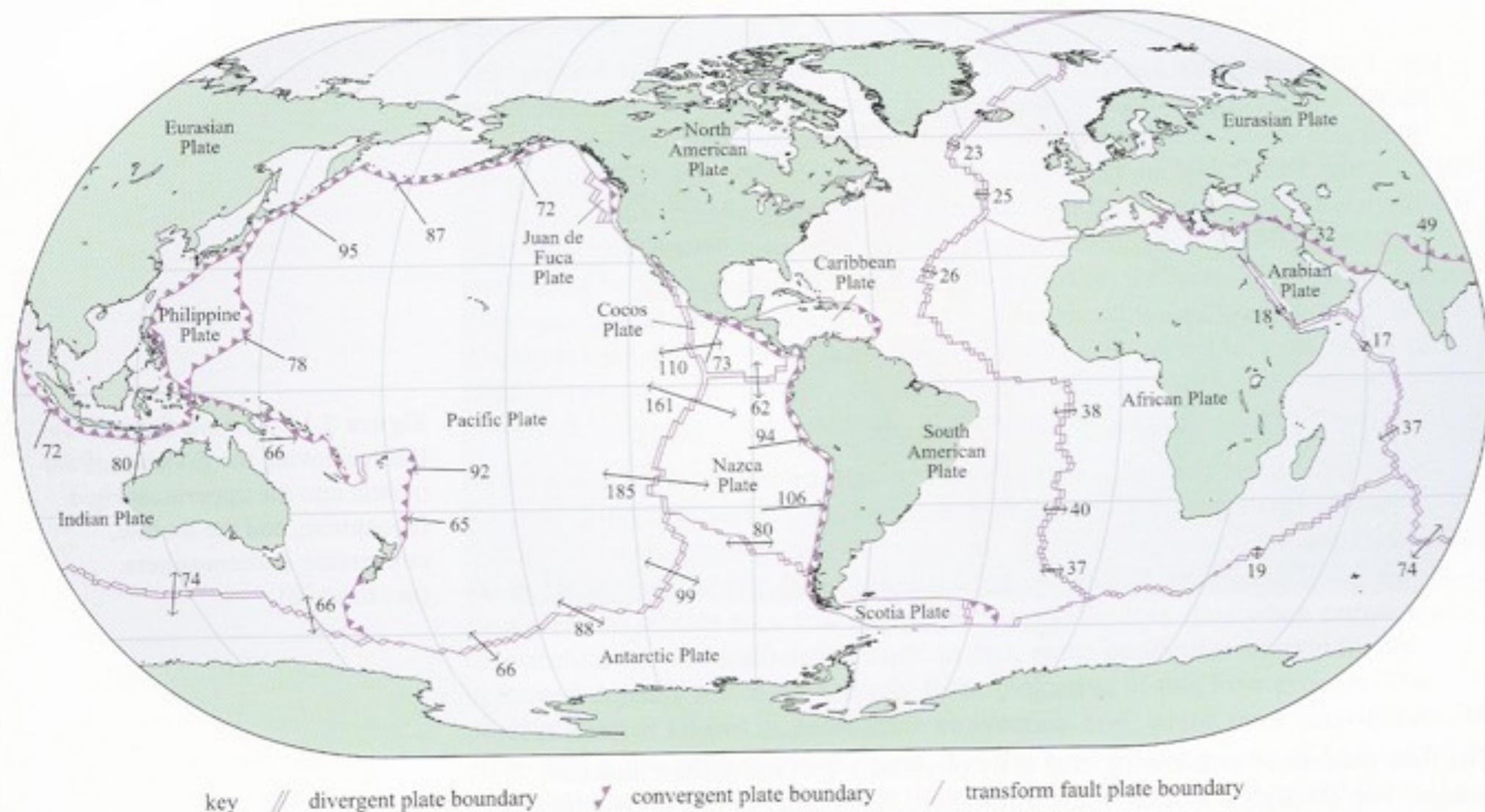
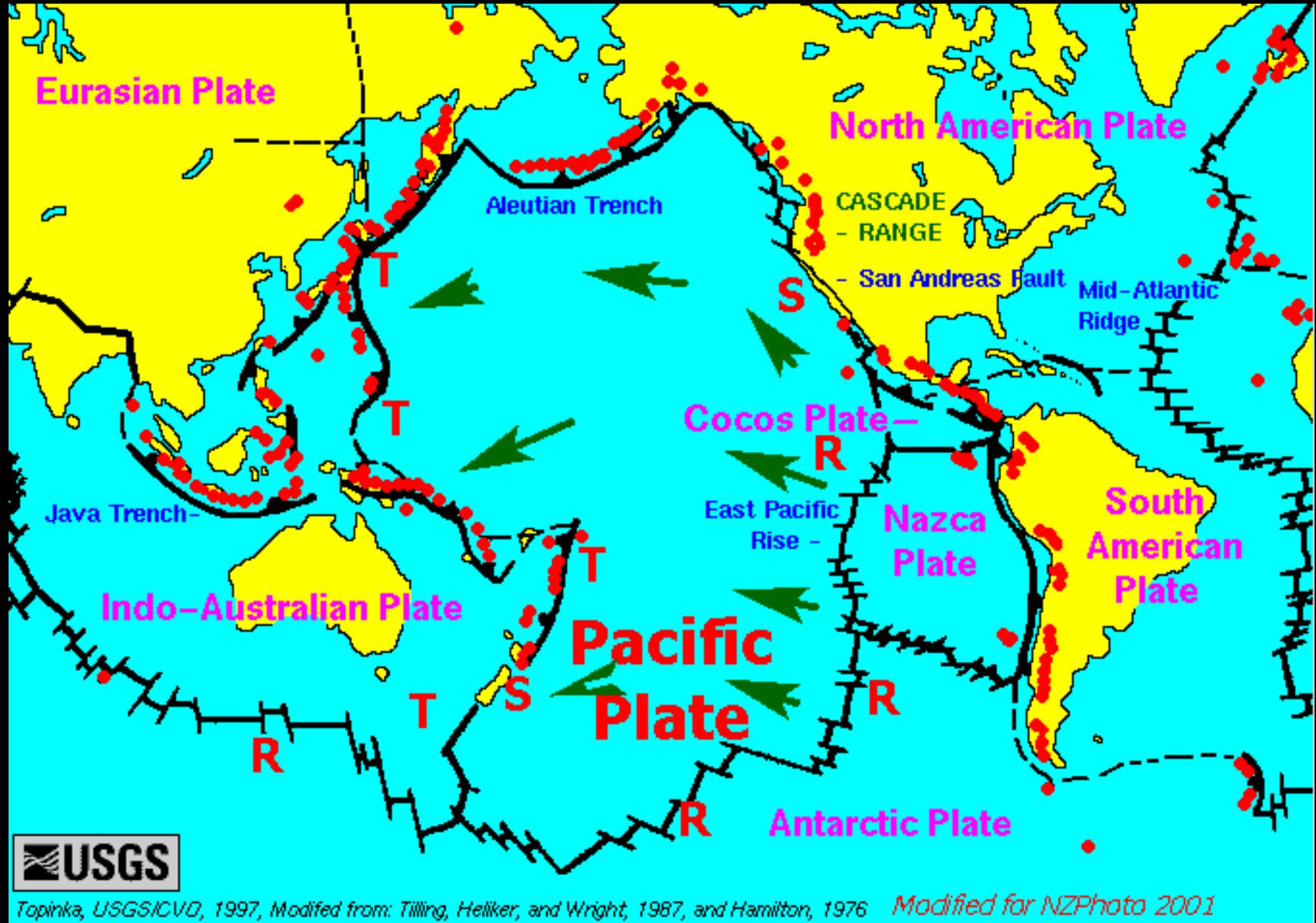
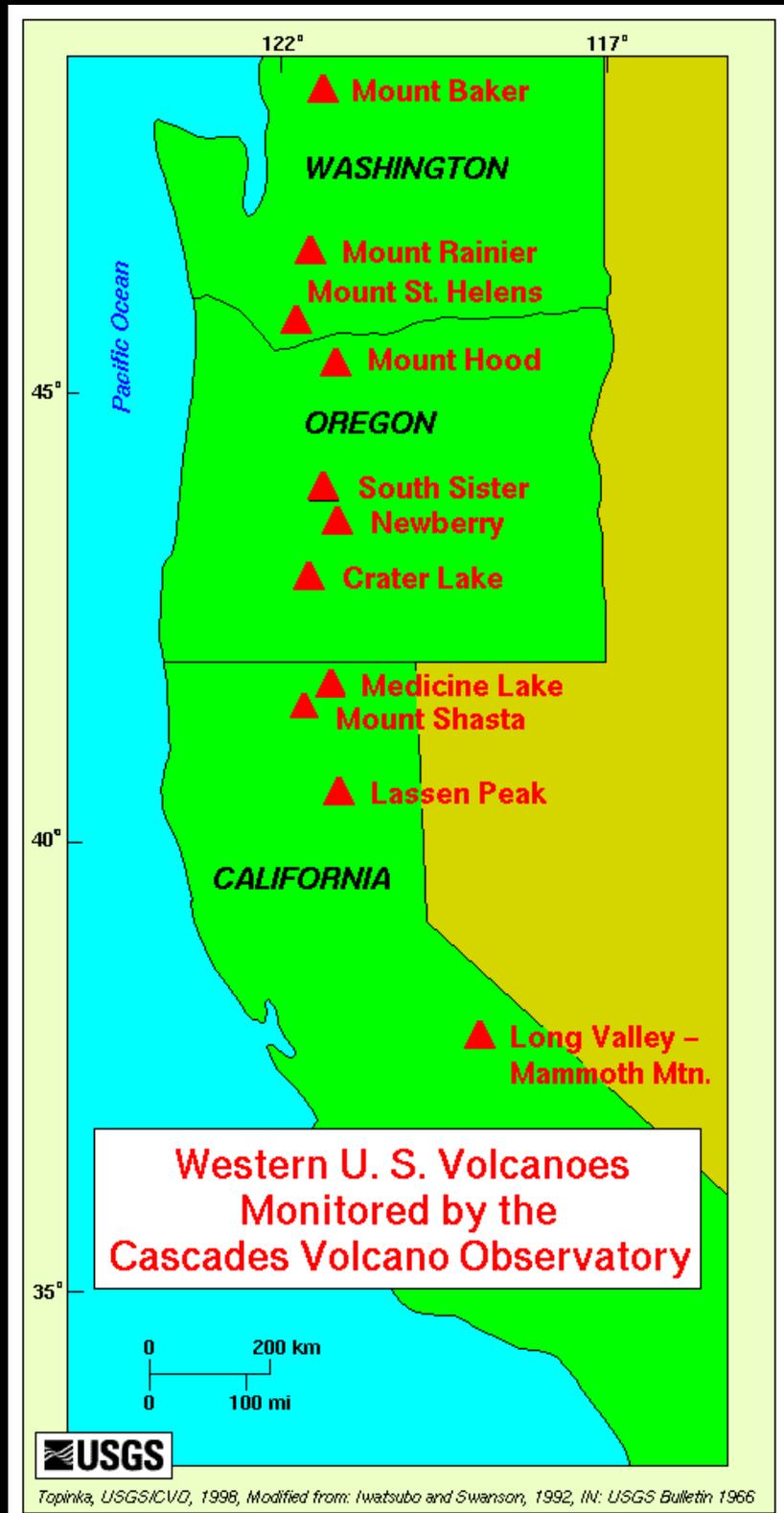
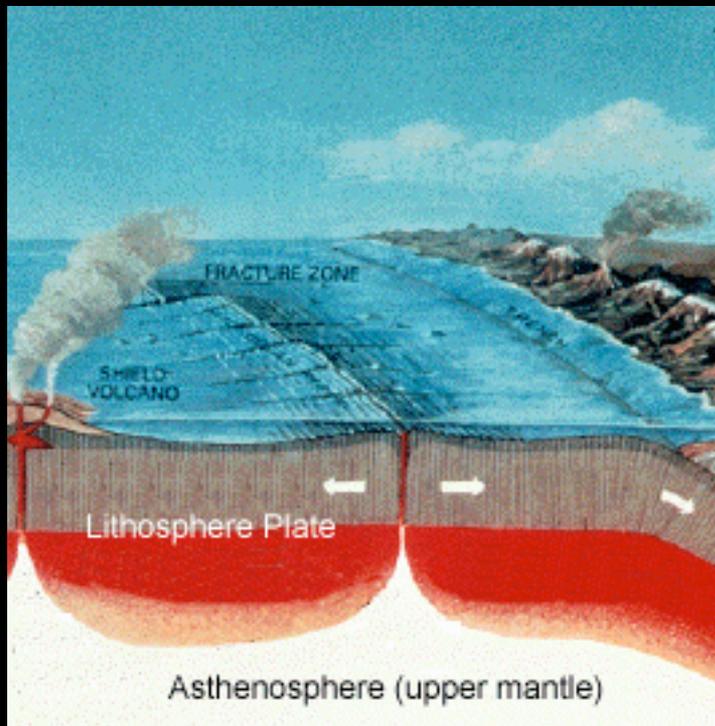


Figure 2.15 Map showing the global distribution of plates and plate boundaries. The black arrows and numbers give the direction and speed of relative motion between plates. Speeds of motion are given in mm yr^{-1} .

Volcanic Locations



Volcano Locations



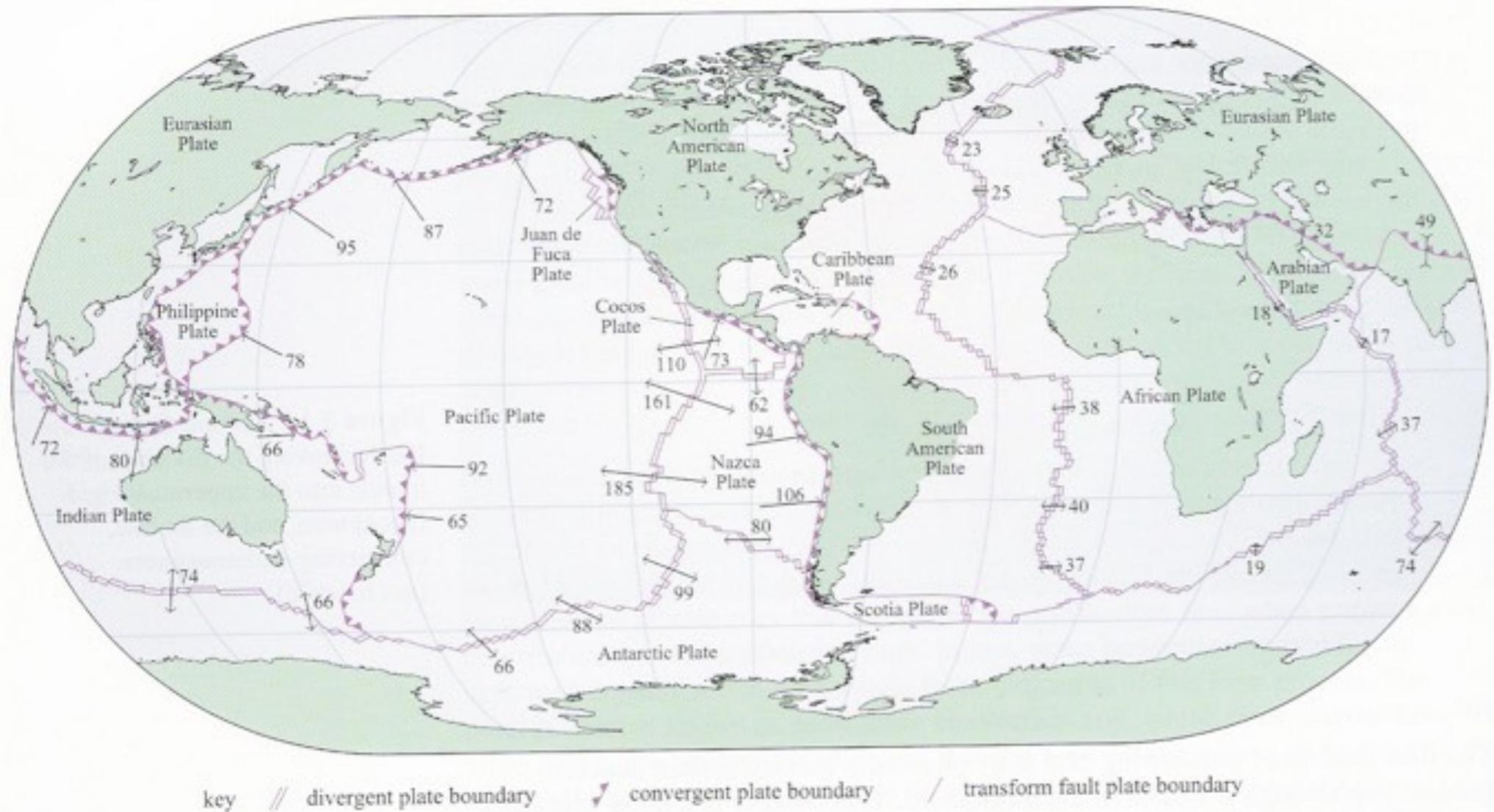


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Earthquake Locations

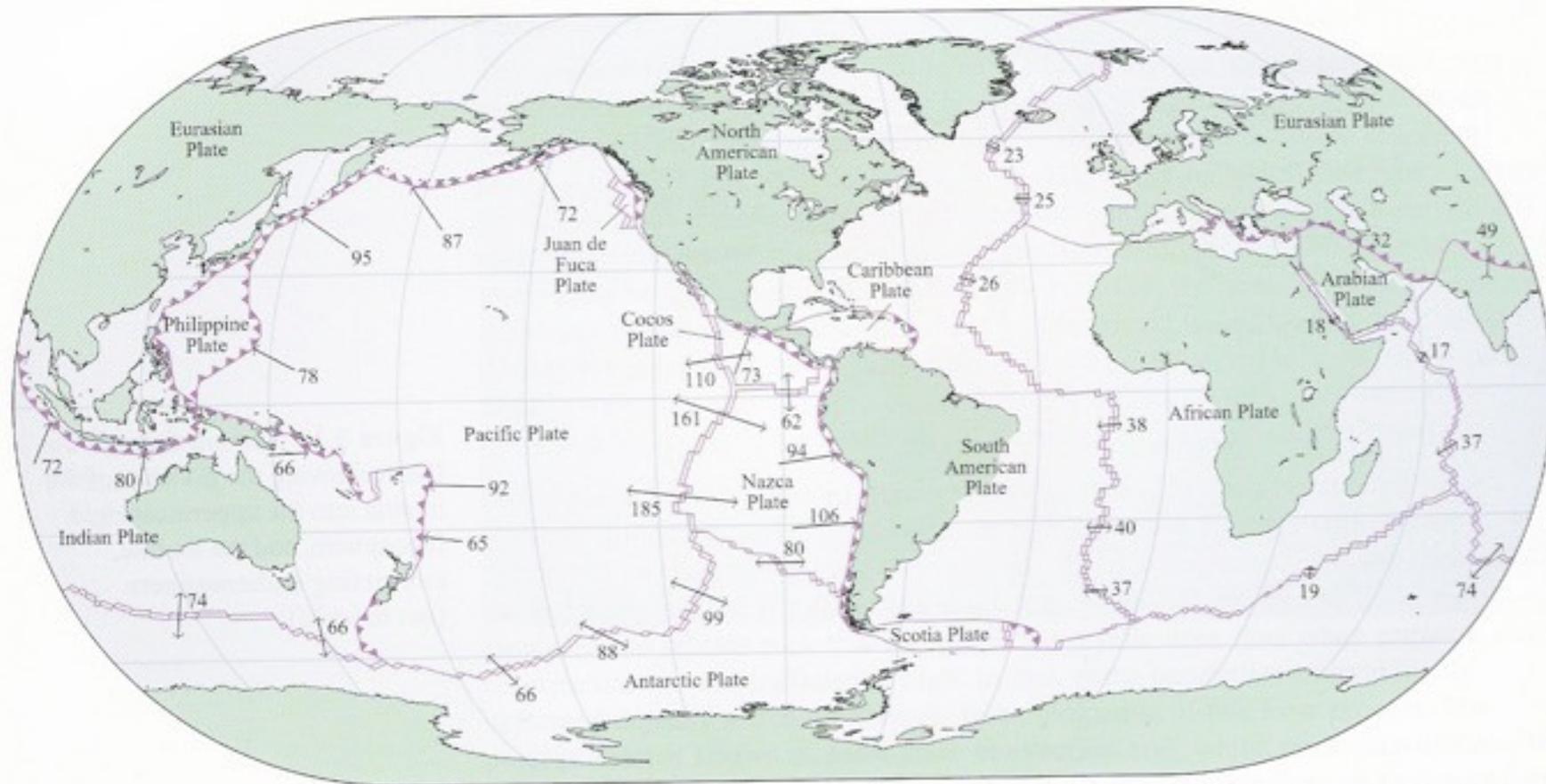


Work in Groups

20 km³/yr new basalt in oceans.

Try to estimate this from measured plate motions!

AN INTRODUCTION TO THE SOLAR SYSTEM



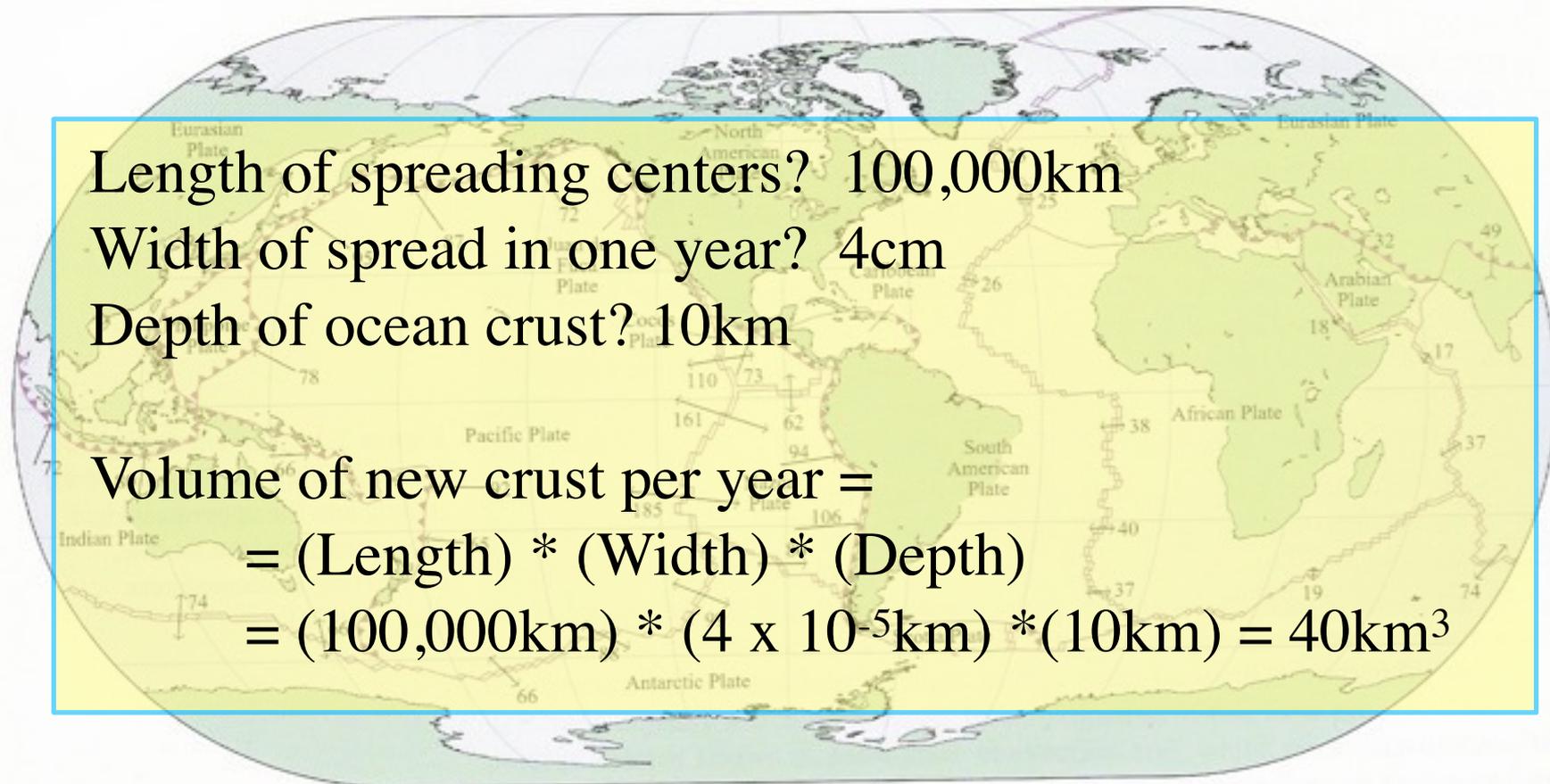
key // divergent plate boundary ↘ convergent plate boundary / transform fault plate boundary

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AN INTRODUCTION TO THE SOLAR SYSTEM

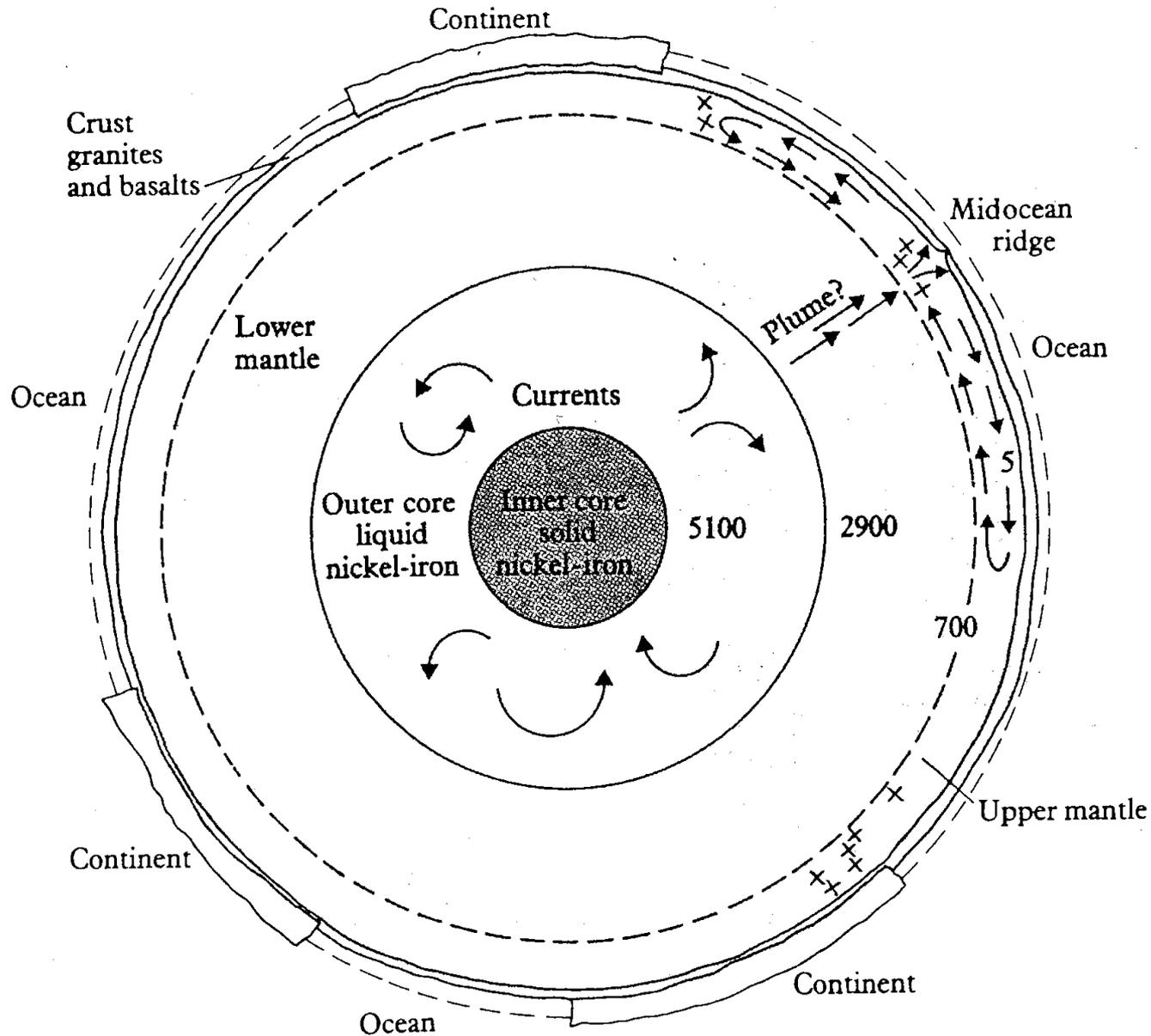


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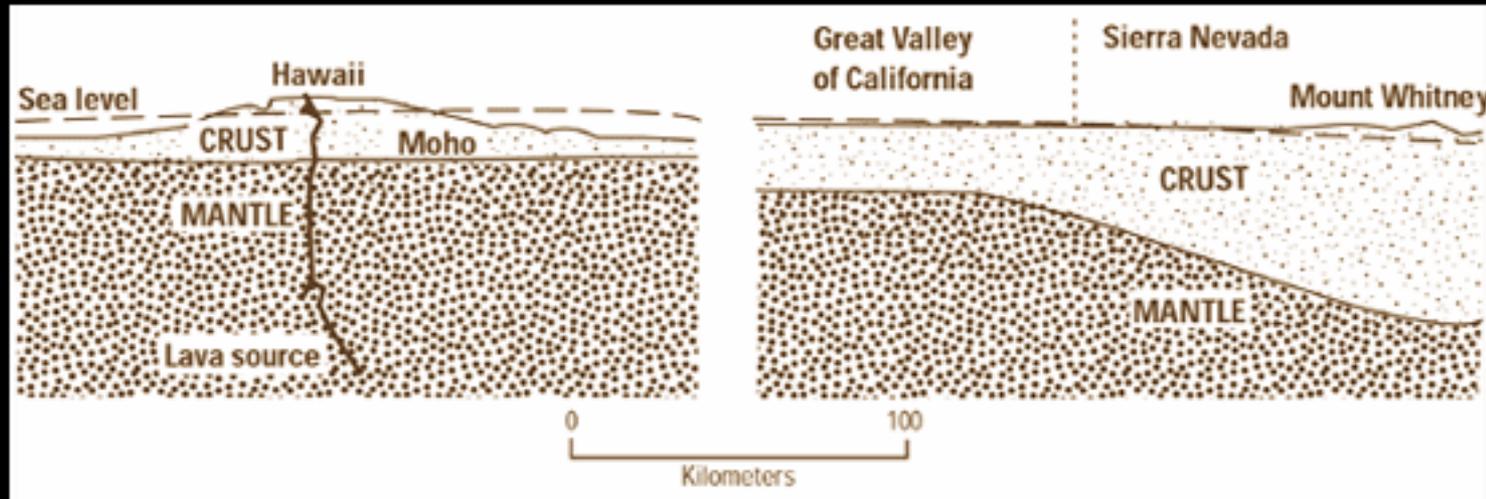
Volcanoes!

1. Mantle Hot Spots on Earth and elsewhere
2. Venus, Mars
3. Moon, Mercury

Convections Cells



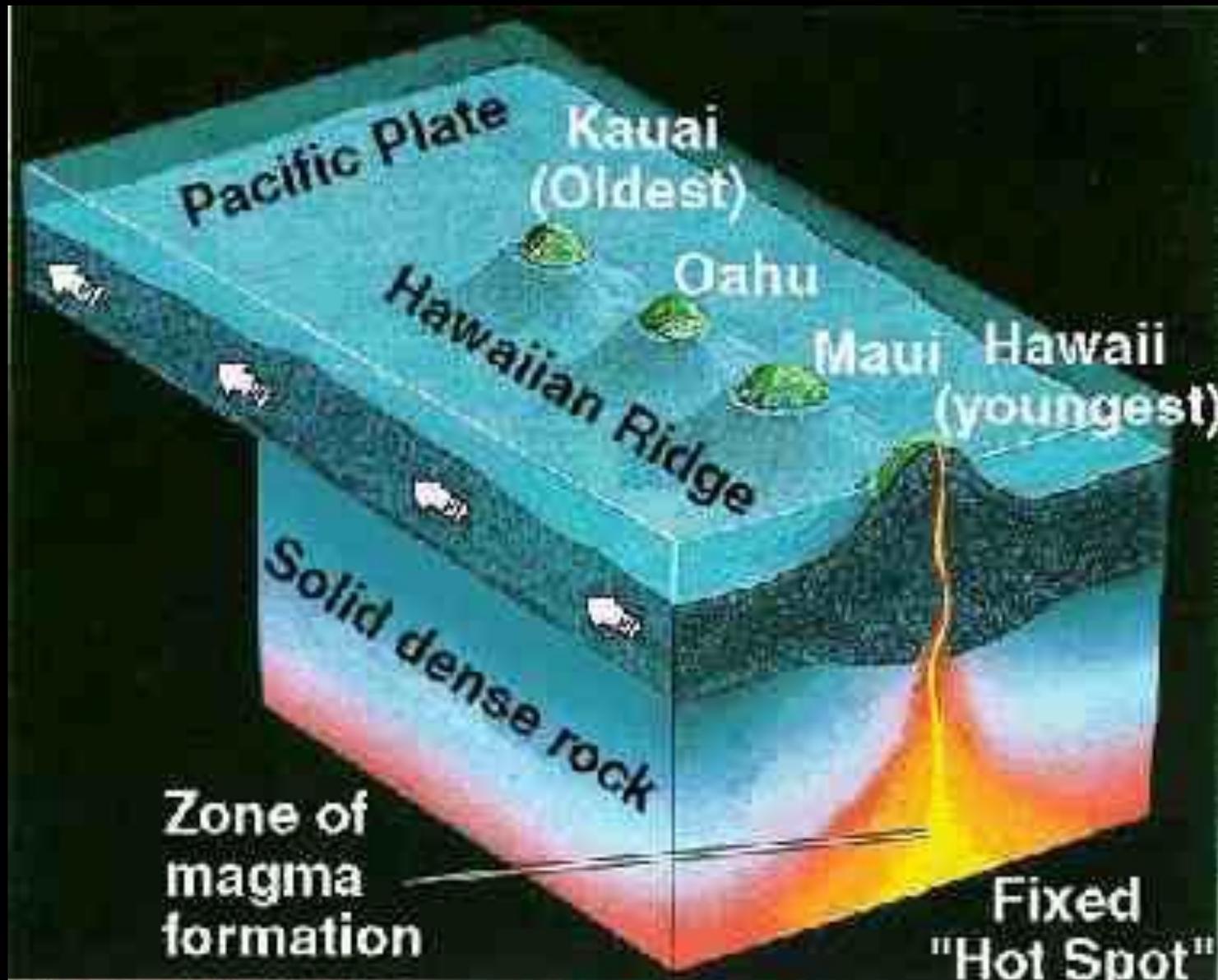
Mantle Hot Spots



Shield Volcano Hawaii forms over a mantle hot spot

Crust is thicker at plate boundaries

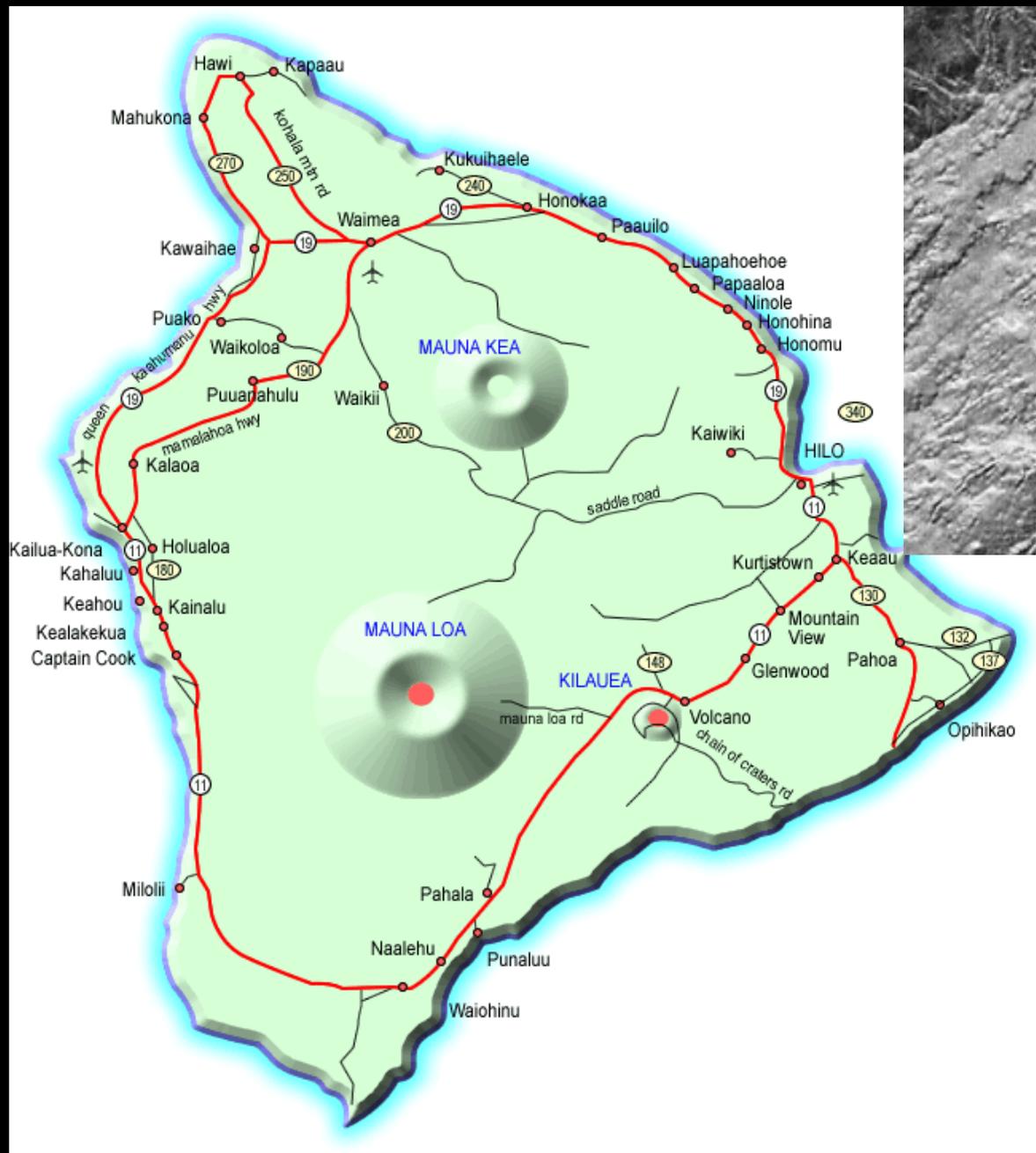
Mantle Hot Spots



Why Does Magma Rise in a “Hot Spot” Mantle Plume?

1. Hot Material is buoyant. Hot air rises, Hot water rises. Hot magma rises.
2. Decompression melting. As the magma rises, the pressure lessens, and some materials that are stable at depth, melt.

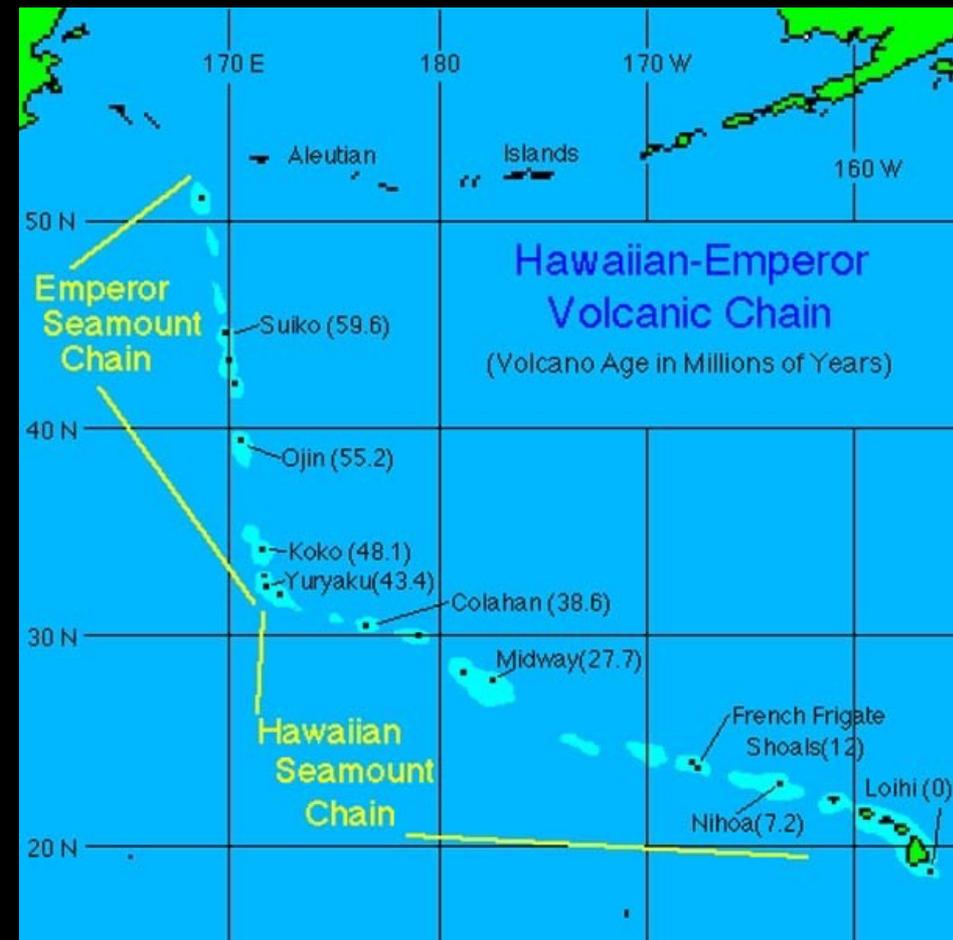
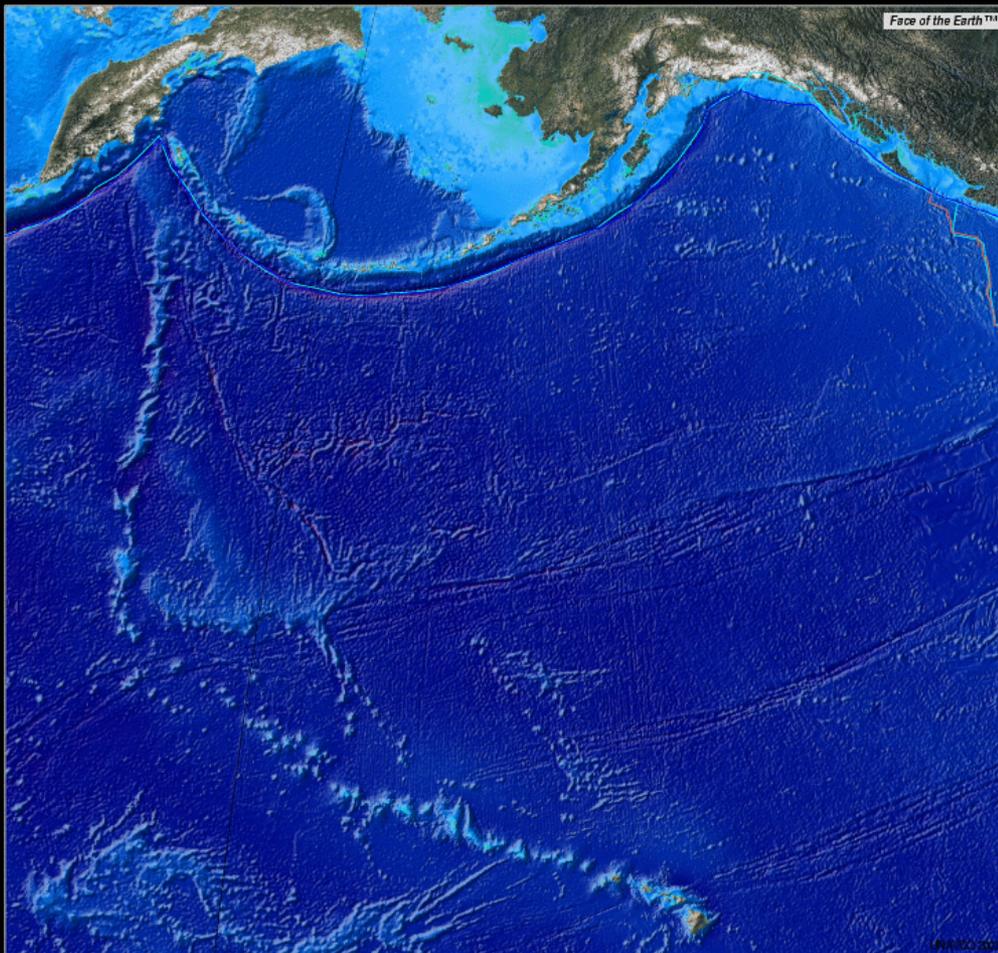
Hawaiian Volcanos



Kilauea Crater



Mantle Hot Spots



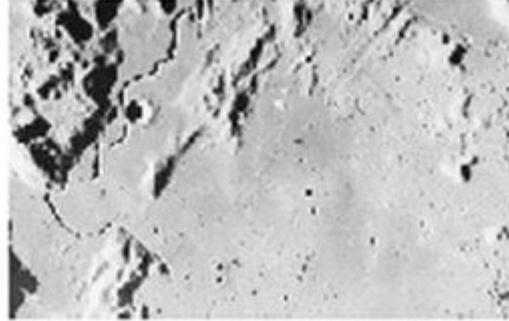
Interesting distribution of seamounts

Different Types of Lava



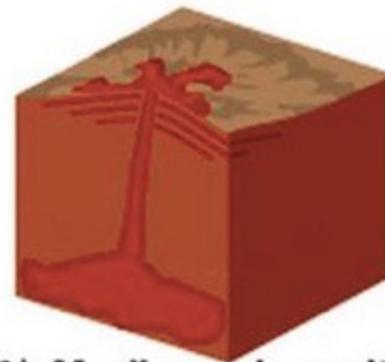
(a) Low-viscosity lava makes flat lava plains.

Lava plains (maria) on the Moon

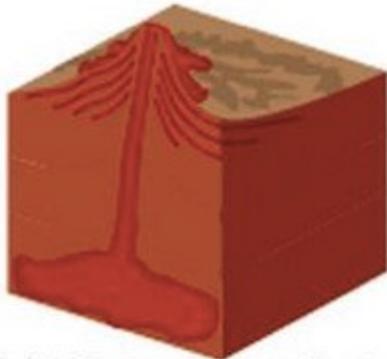


Fluid Lavas

Olympus Mons (Mars)

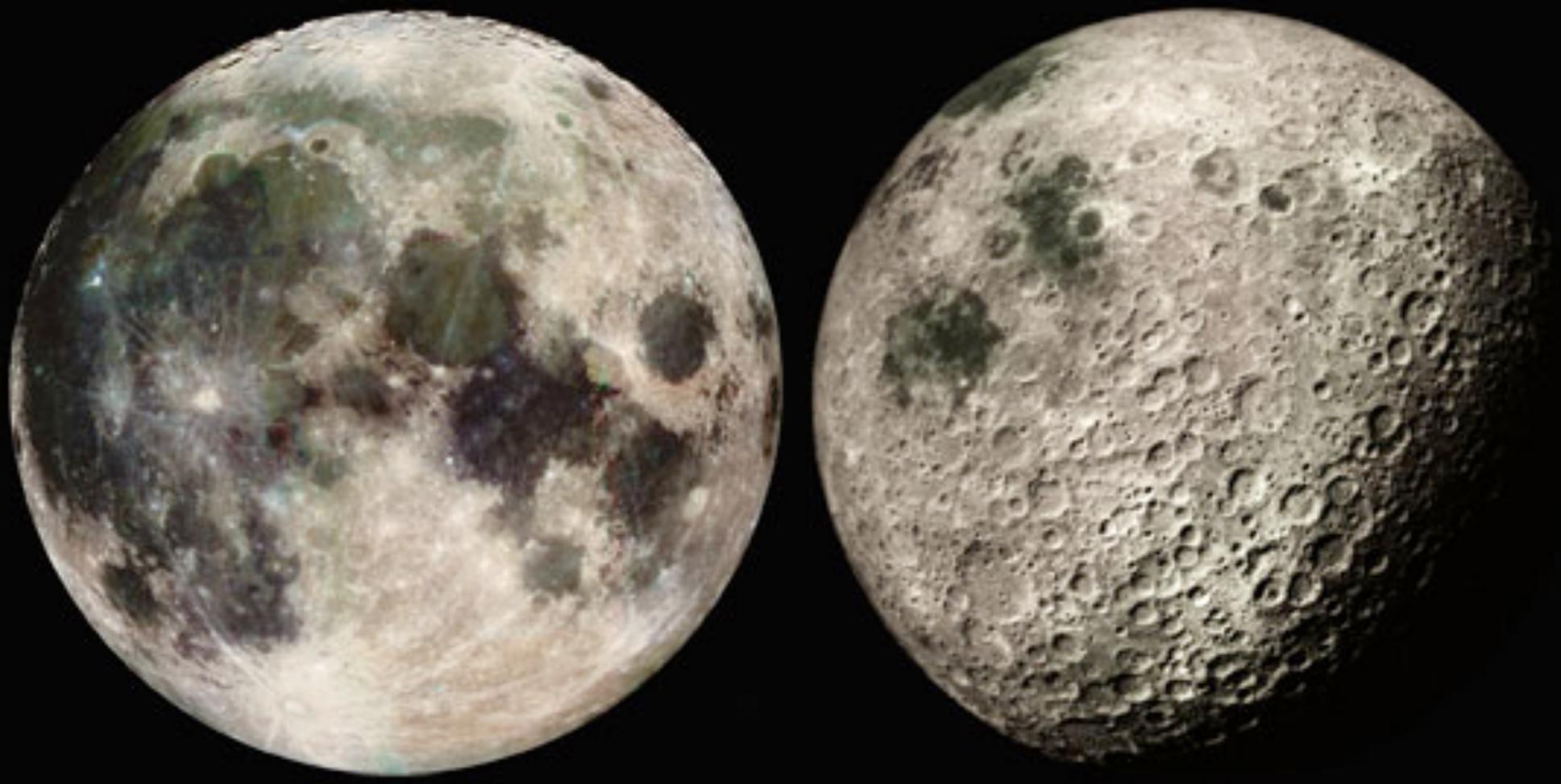


(b) Medium-viscosity lava makes shallow-sloped shield volcanoes.



(c) High-viscosity lava makes steep-sloped stratovolcanoes.

Thickness of the Moon's Crust

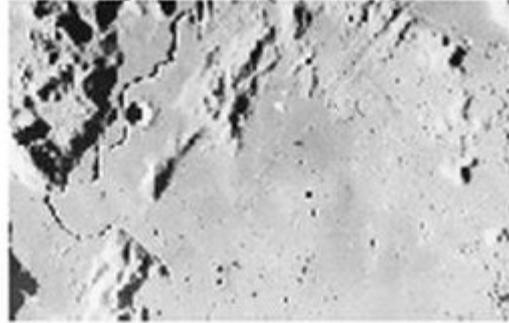


Different Types of Lava

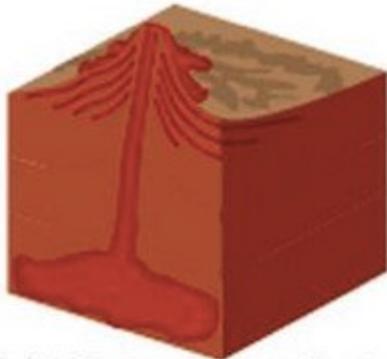


(a) Low-viscosity lava makes flat lava plains.

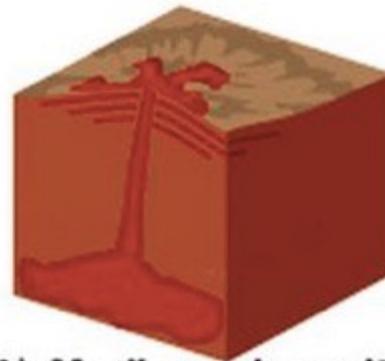
Lava plains (maria) on the Moon



Mid Range Lavas



(c) High-viscosity lava makes steep-sloped stratovolcanoes.

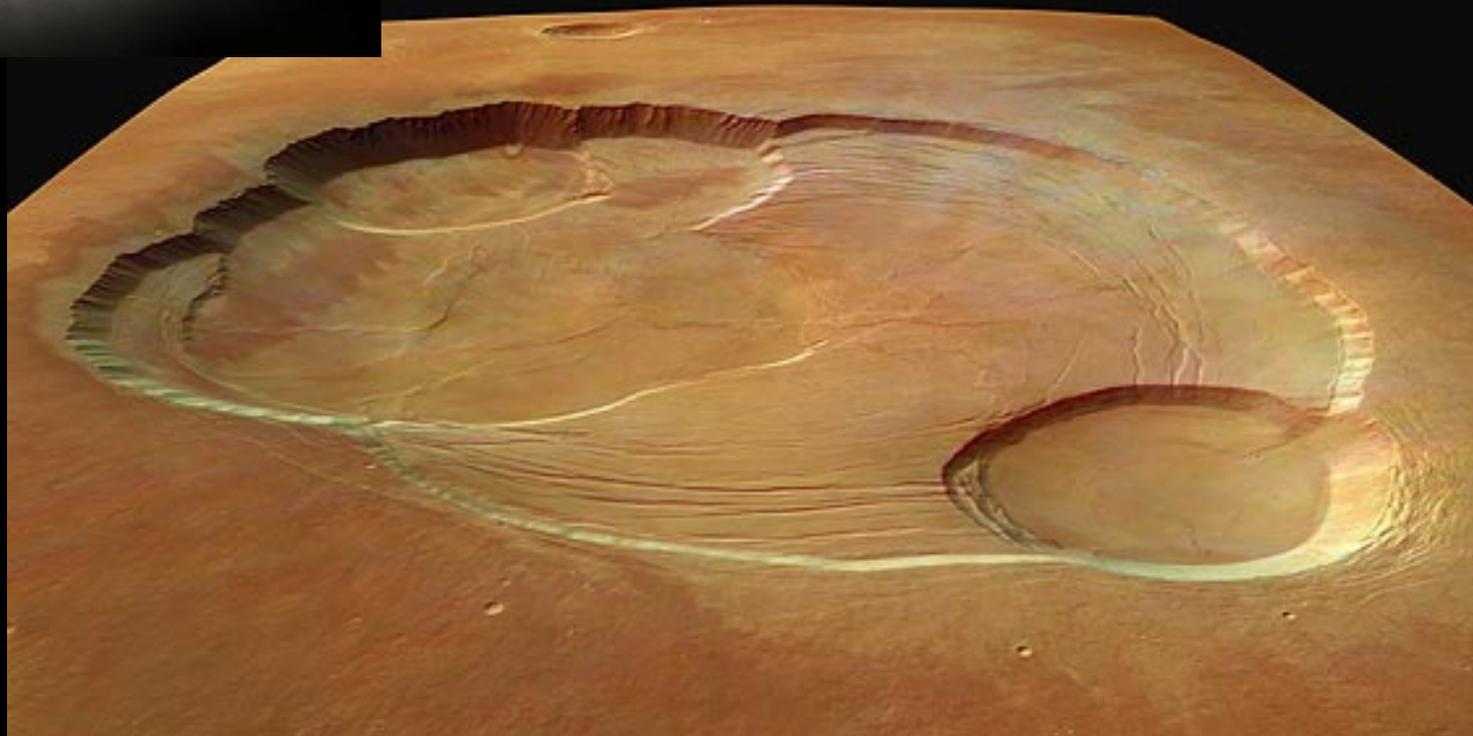
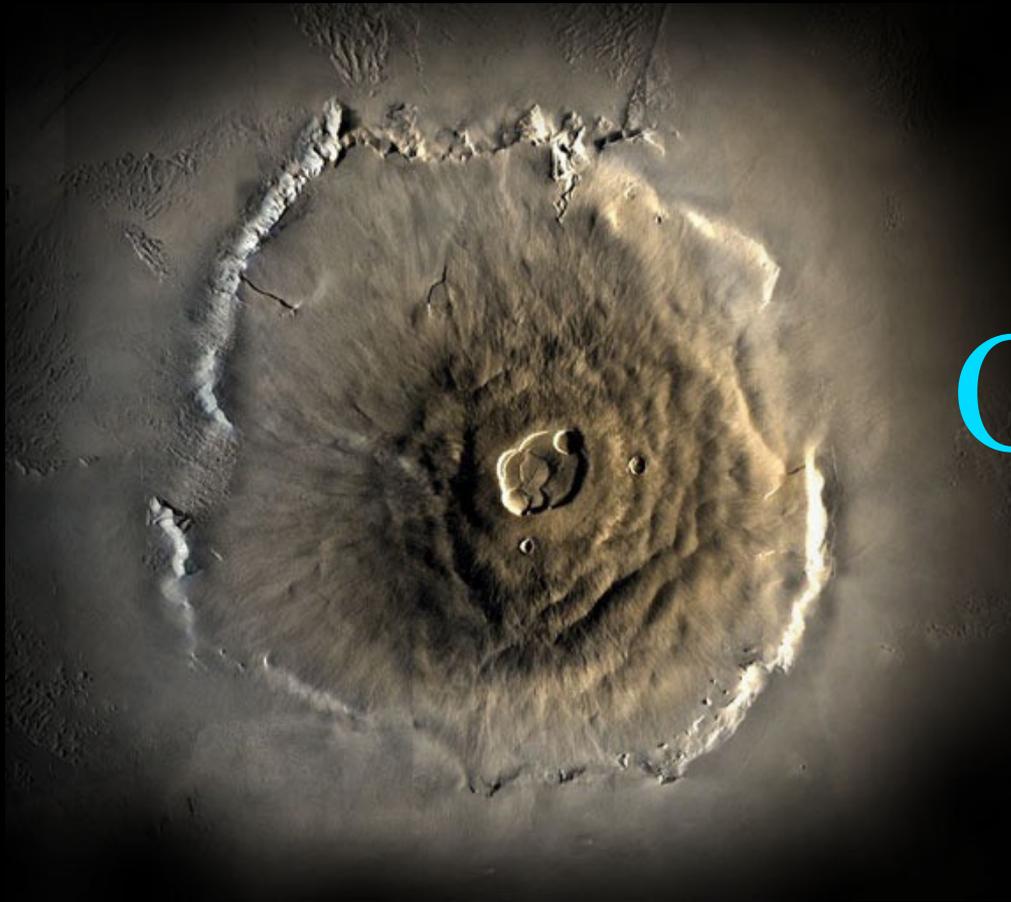


(b) Medium-viscosity lava makes shallow-sloped shield volcanoes.

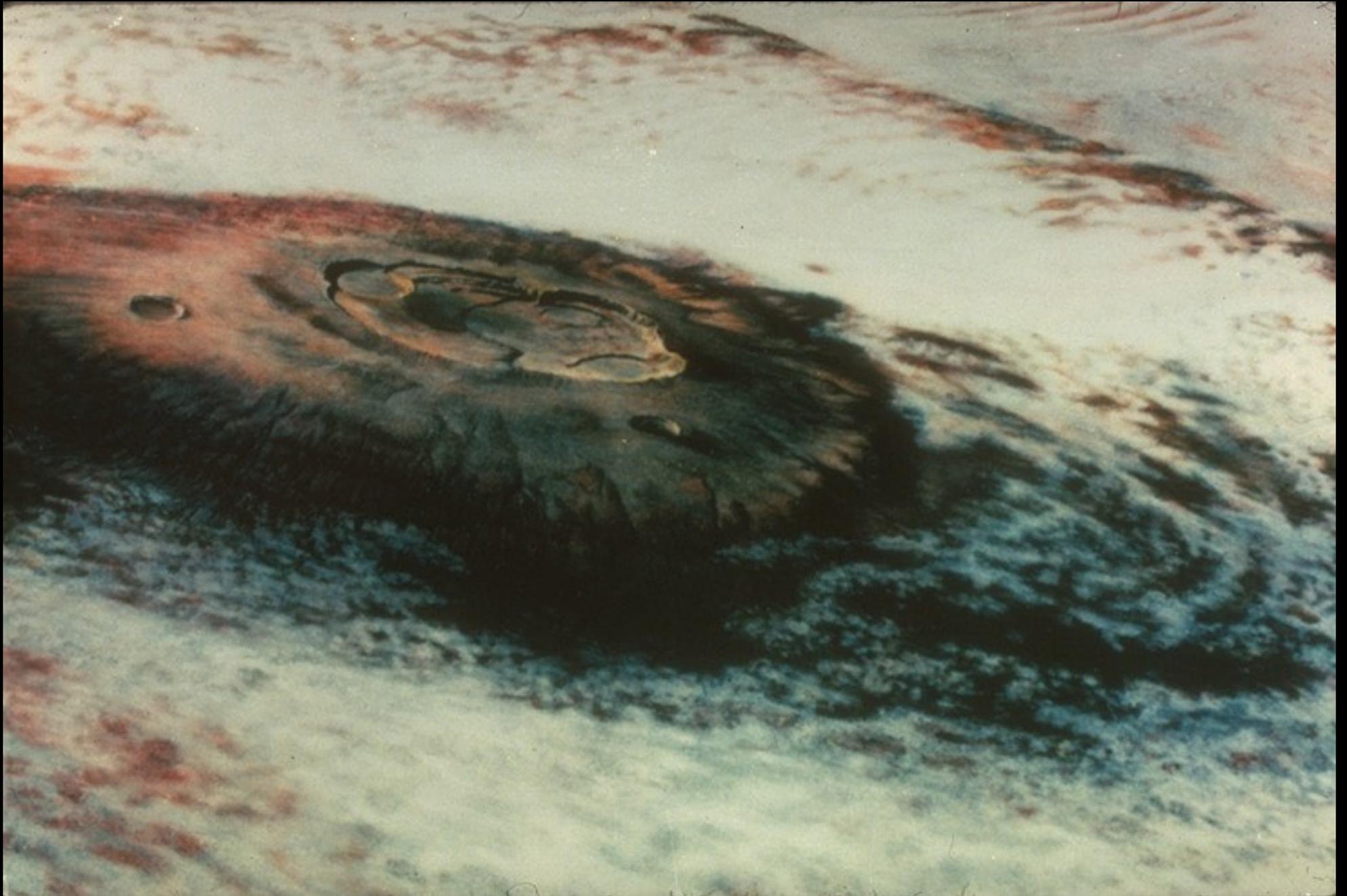
Olympus Mons (Mars)



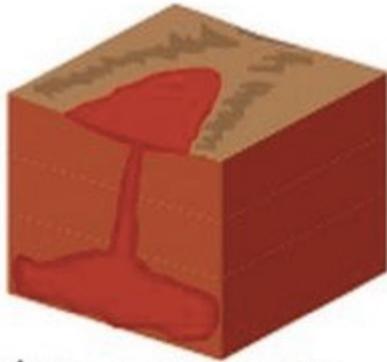
Mars: Olympus Mons



Olympus Mons

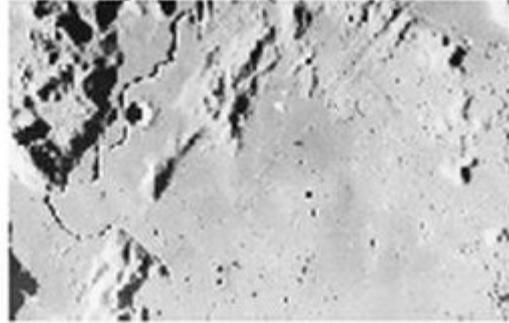


Different Types of Lava

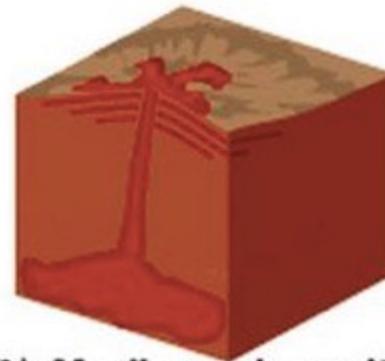


(a) Low-viscosity lava makes flat lava plains.

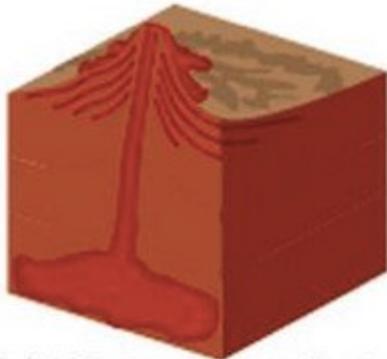
Lava plains (maria) on the Moon



Olympus Mons (Mars)



(b) Medium-viscosity lava makes shallow-sloped shield volcanoes.



(c) High-viscosity lava makes steep-sloped stratovolcanoes.

Thick Lavas

Mt. Fuji



Highest Mountains

11.2 Global Perspective

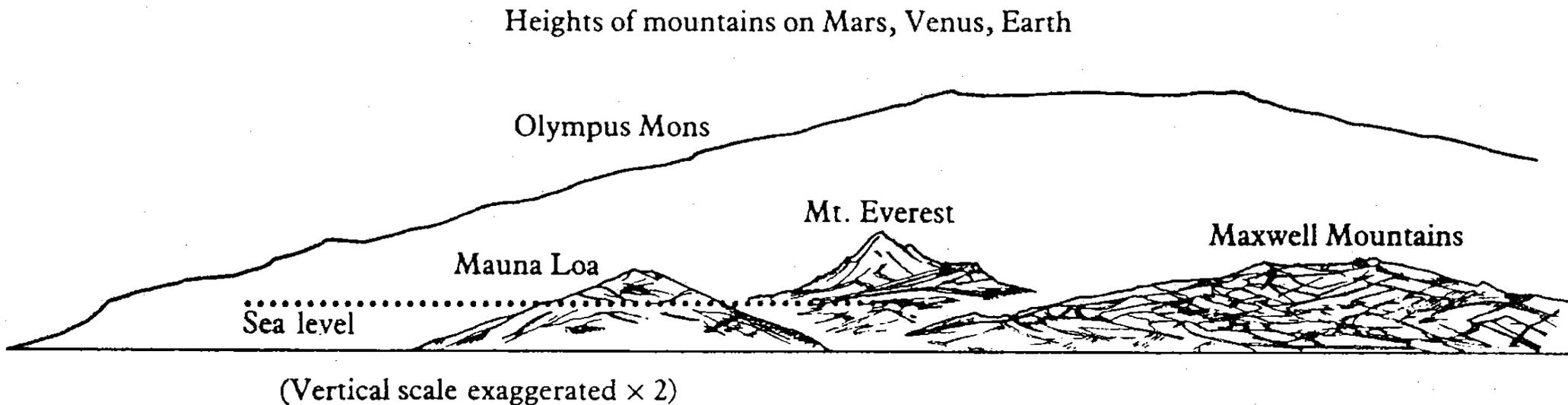
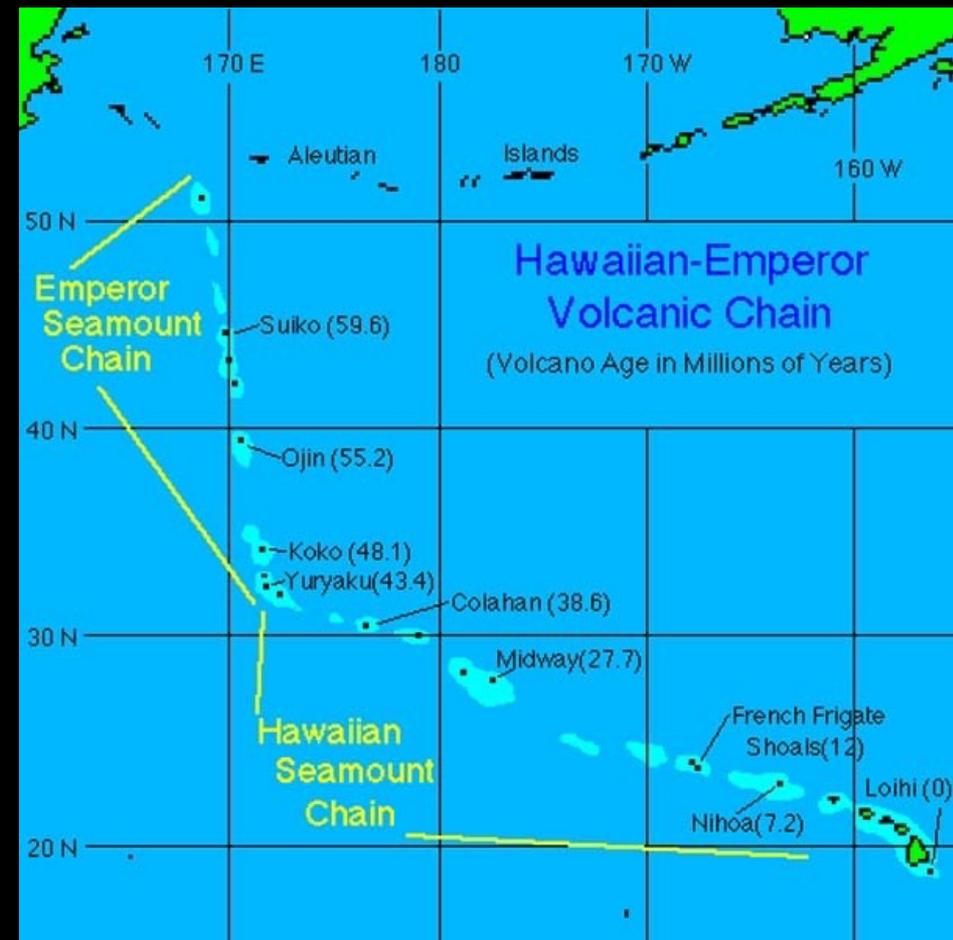
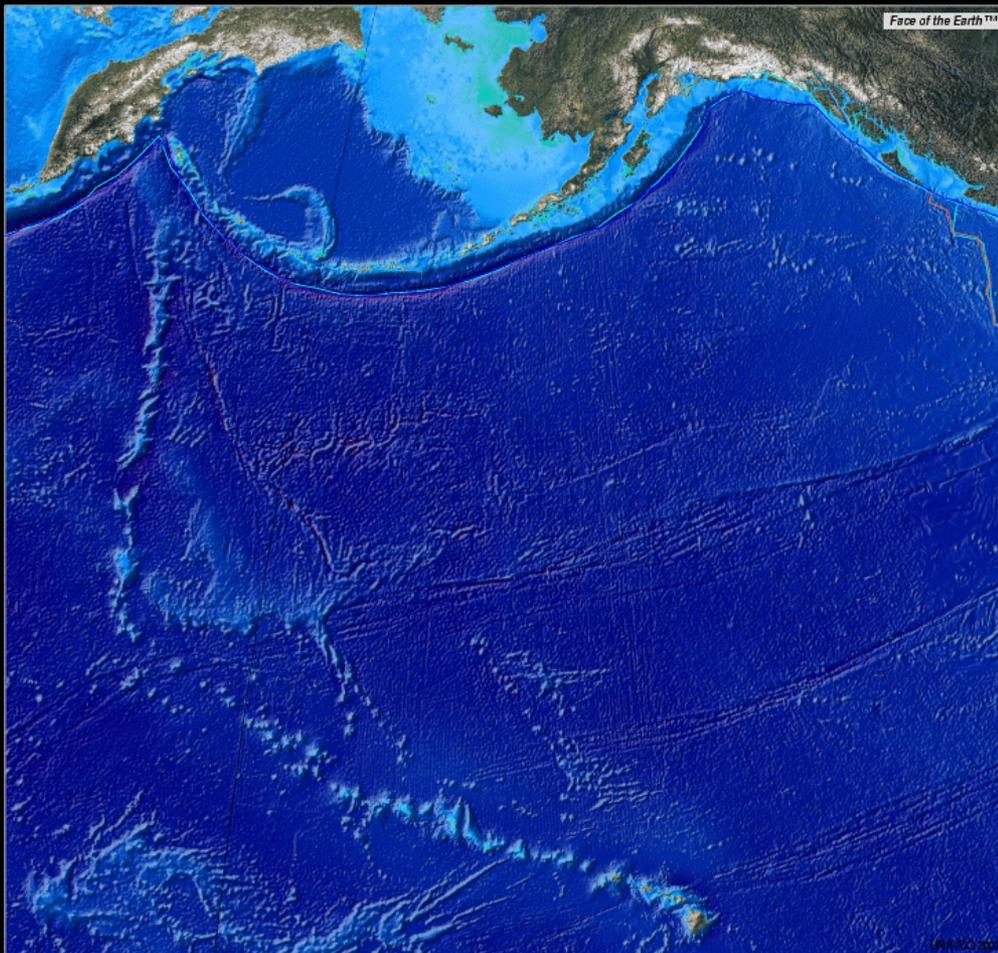
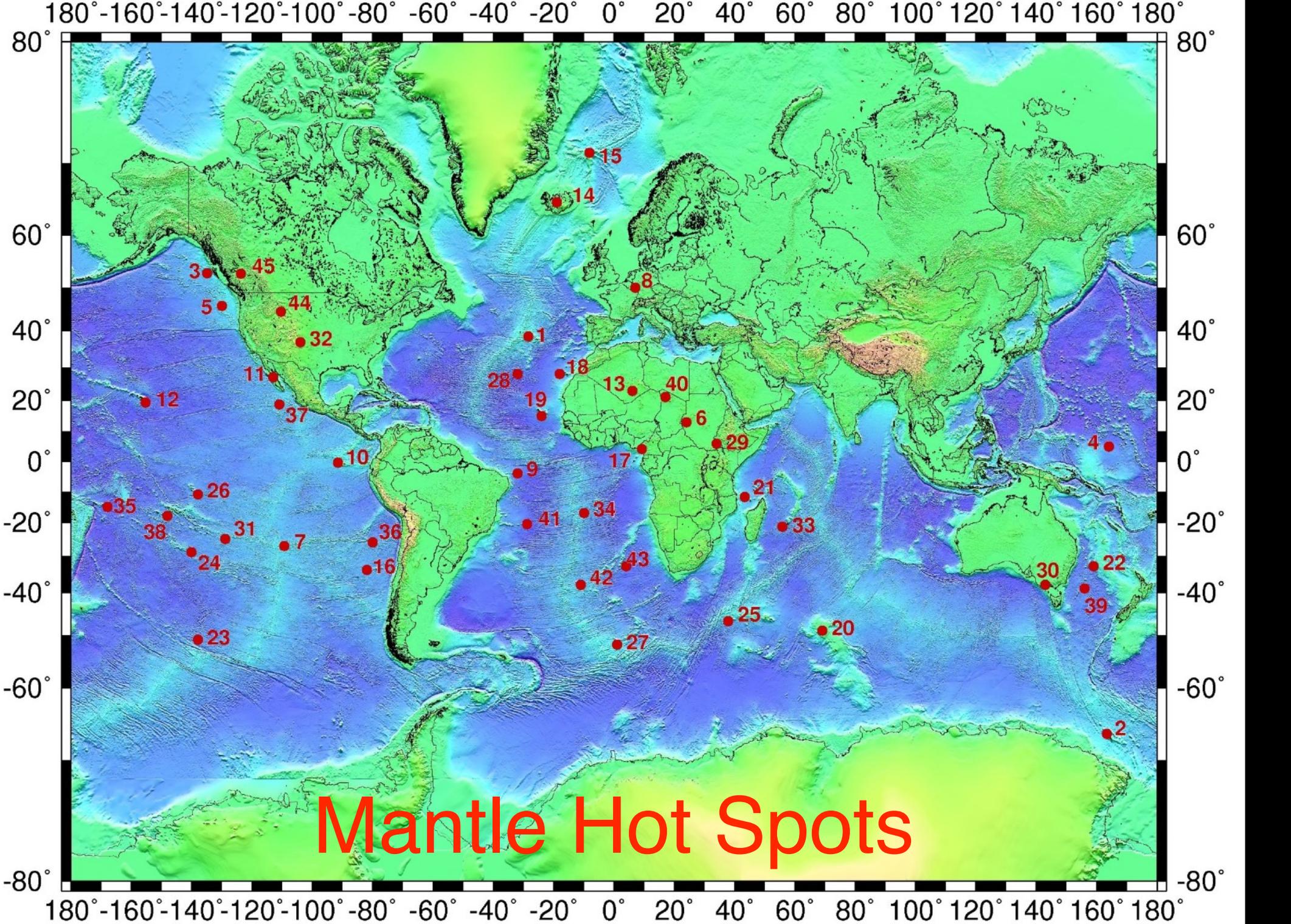


Figure 11.7 Comparison of highest elevations on Venus, Earth, and Mars. Martian mountains can grow higher because Mars has weaker gravity. Vertical scale is exaggerated by a factor of three.

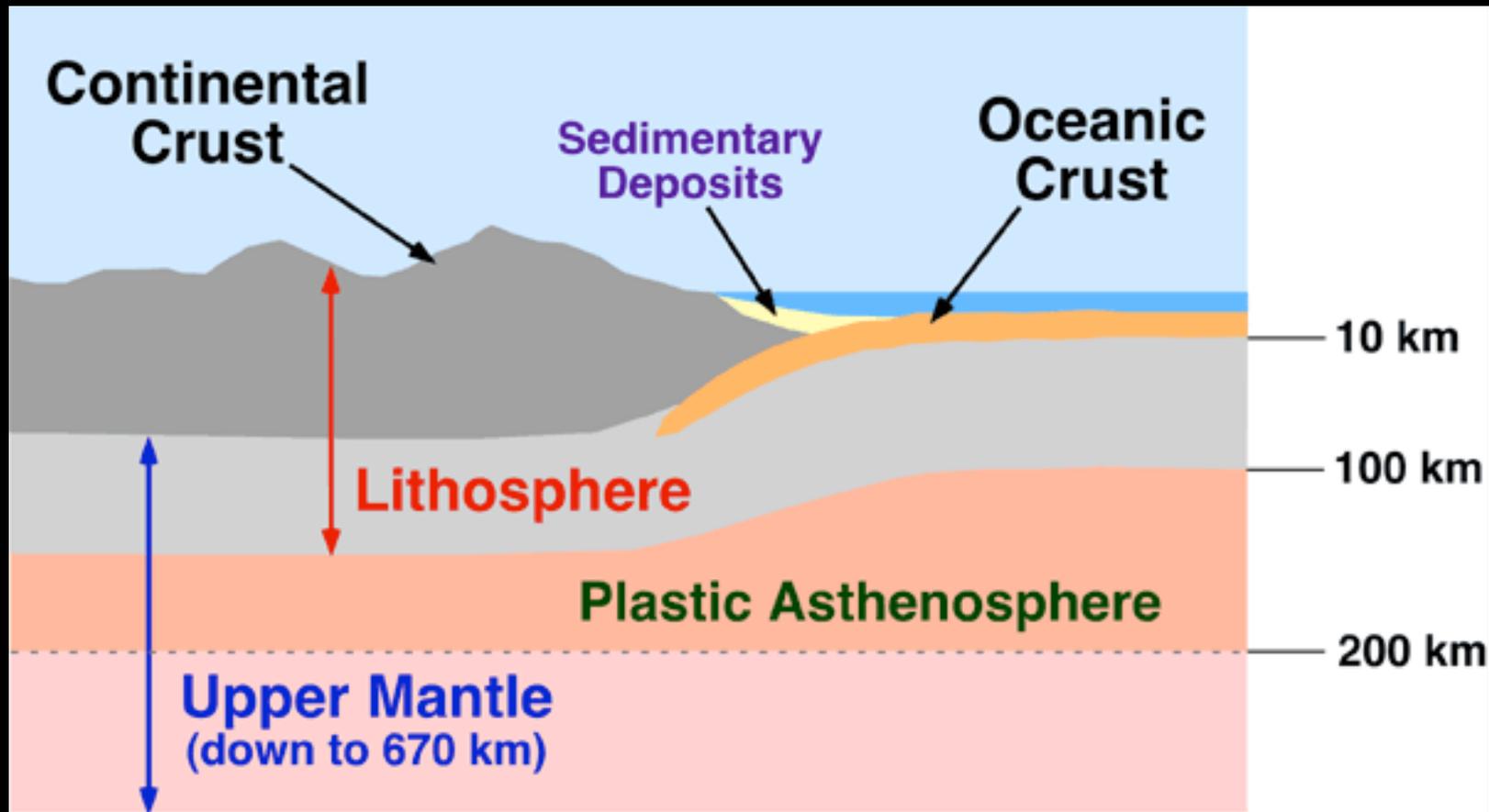
Mantle Hot Spots



Interesting distribution of seamounts



Layers of the Earth



Melting by Hydration

words, mantle altered in this manner begins to melt at lower temperatures and pressures than unaffected, or anhydrous, mantle. This is the process of hydration-induced melting.

Hydration of mantle peridotite occurs as a consequence of plate tectonic recycling (Sections 2.4.5 and 2.5.1). It takes place at depth within subduction zones where old, water-rich oceanic crust is reabsorbed into the mantle (Figure 3.9a). The resulting generation and surfaceward migration of magma produces arcuate belts of volcanoes, known as **volcanic arcs** (Figure 3.9b).

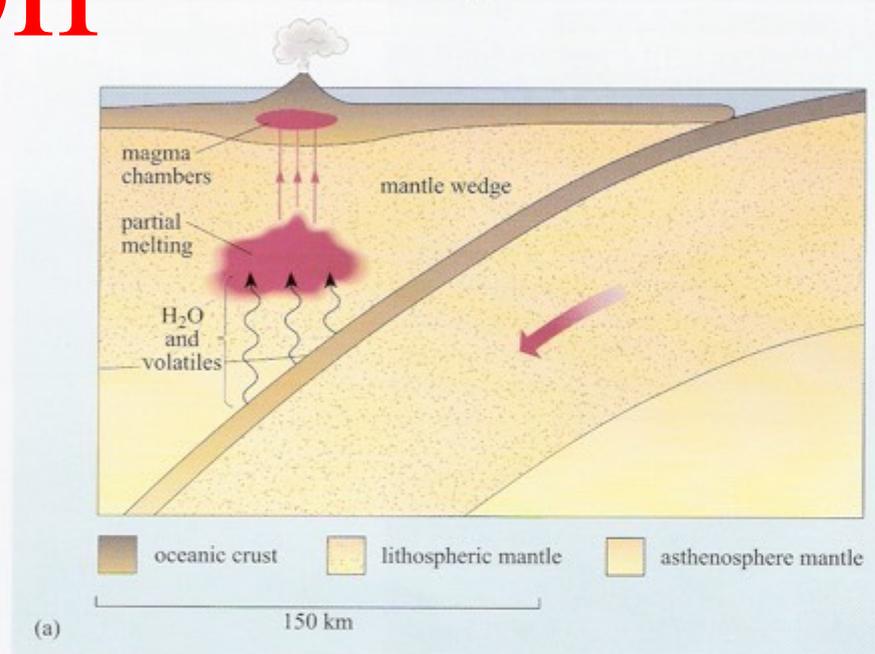
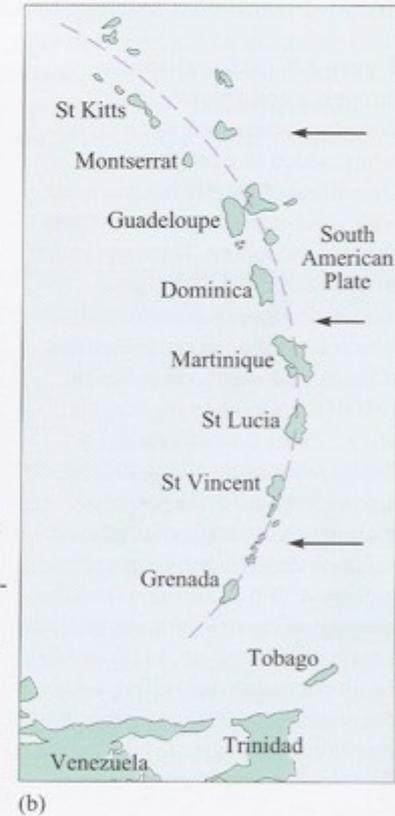


Figure 3.9 (a) Cross-section through a destructive plate boundary where old oceanic-crust material is being recycled back into the mantle during subduction. Heating of subducted material during its descent releases hydrous or gas-rich fluids which then rise and are added to the mantle wedge of the over-riding plate. These released fluids cause hydration reactions that reduce the melting point of the mantle-wedge mineral assemblage, so producing magmas. The volcanoes associated with this magmatism are typically arranged in arcuate belts (volcanic arcs) such as those of the Andes, Central America, and the Lesser Antilles of the Caribbean. (b) Sketch map of the volcanically generated islands of the Lesser Antilles that together form the volcanic arc of the eastern Caribbean. (Arrows show the relative movement of the South American Plate, which is presently being subducted beneath the Caribbean Plate.)



Partial Melting

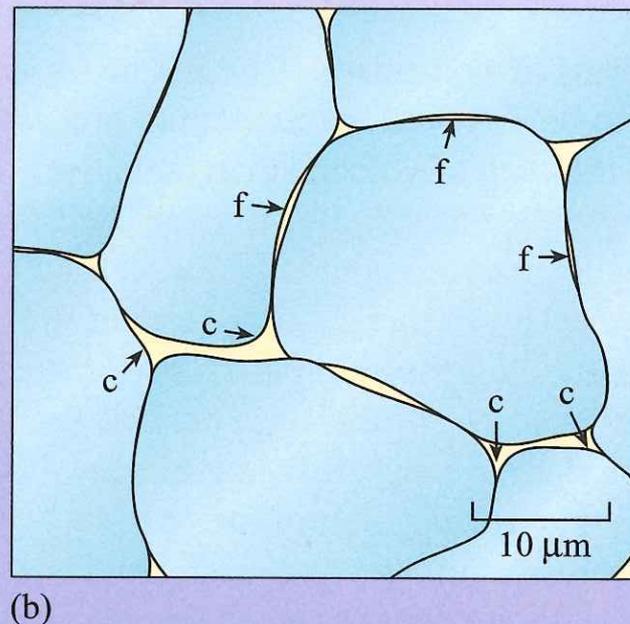
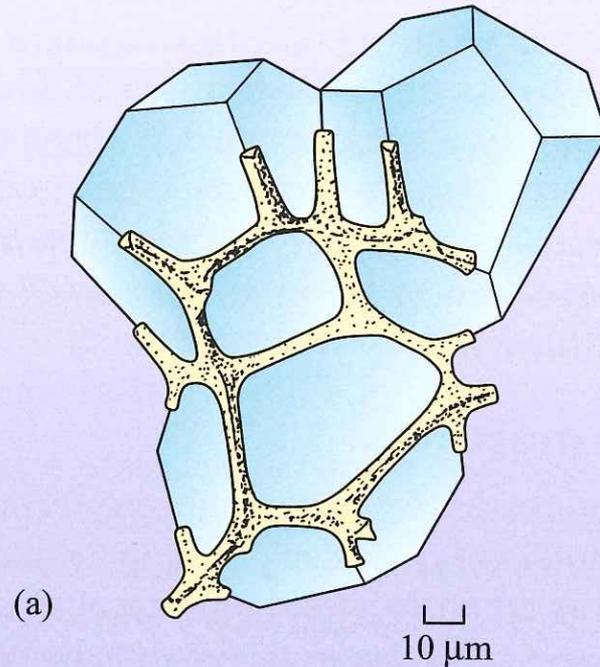


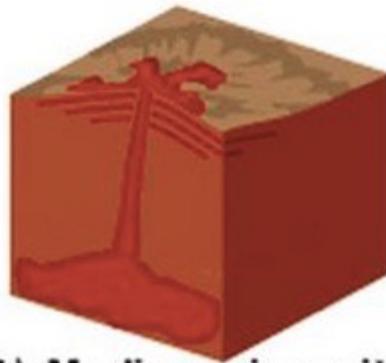
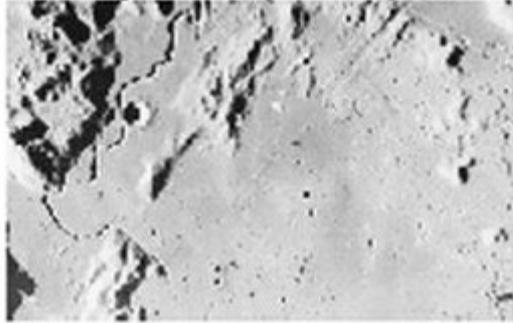
Figure 3.10 (a) Sketch showing a three-dimensional view of adjacent crystals in a rock in which partial melting has begun. The melt initially forms in the pockets (interstices) where three crystals meet and, with further melting, the pockets then become connected by channels of melt running along the grain boundaries. Once interconnected, the melt can then begin to migrate and coalesce into a larger body of magma. (b) Magnified image of a rock beginning to melt (melt: yellow, crystals: blue). Features labelled 'c' are interconnecting melt channels at crystal interstices and boundaries; features labelled 'f' are thinner films of melt coating grain boundaries that may not yet be sufficiently developed to form connections.

Different Types of Lava



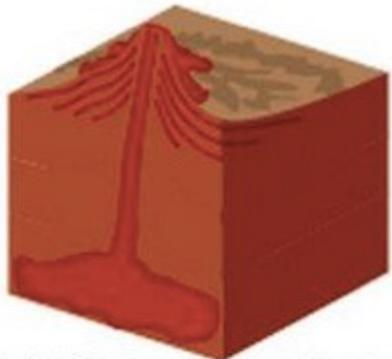
(a) Low-viscosity lava makes flat lava plains.

Lava plains (maria) on the Moon



(b) Medium-viscosity lava makes shallow-sloped shield volcanoes.

Olympus Mons (Mars)



(c) High-viscosity lava makes steep-sloped stratovolcanoes.

Basaltic (low Silicon) lavas: effusive eruptions

Silicon-rich lavas: explosive eruptions

Midterm Exam: Thursday in Class

Three Questions (Two Quantitative, One Qualitative)

- Everything Covered in Class
- Chapter 5 Atmospheres
- Chapter 6 Surfaces/Interiors

1 page of notes front/back

Effusive Eruptions



Mauna Loa

Explosive Eruptions

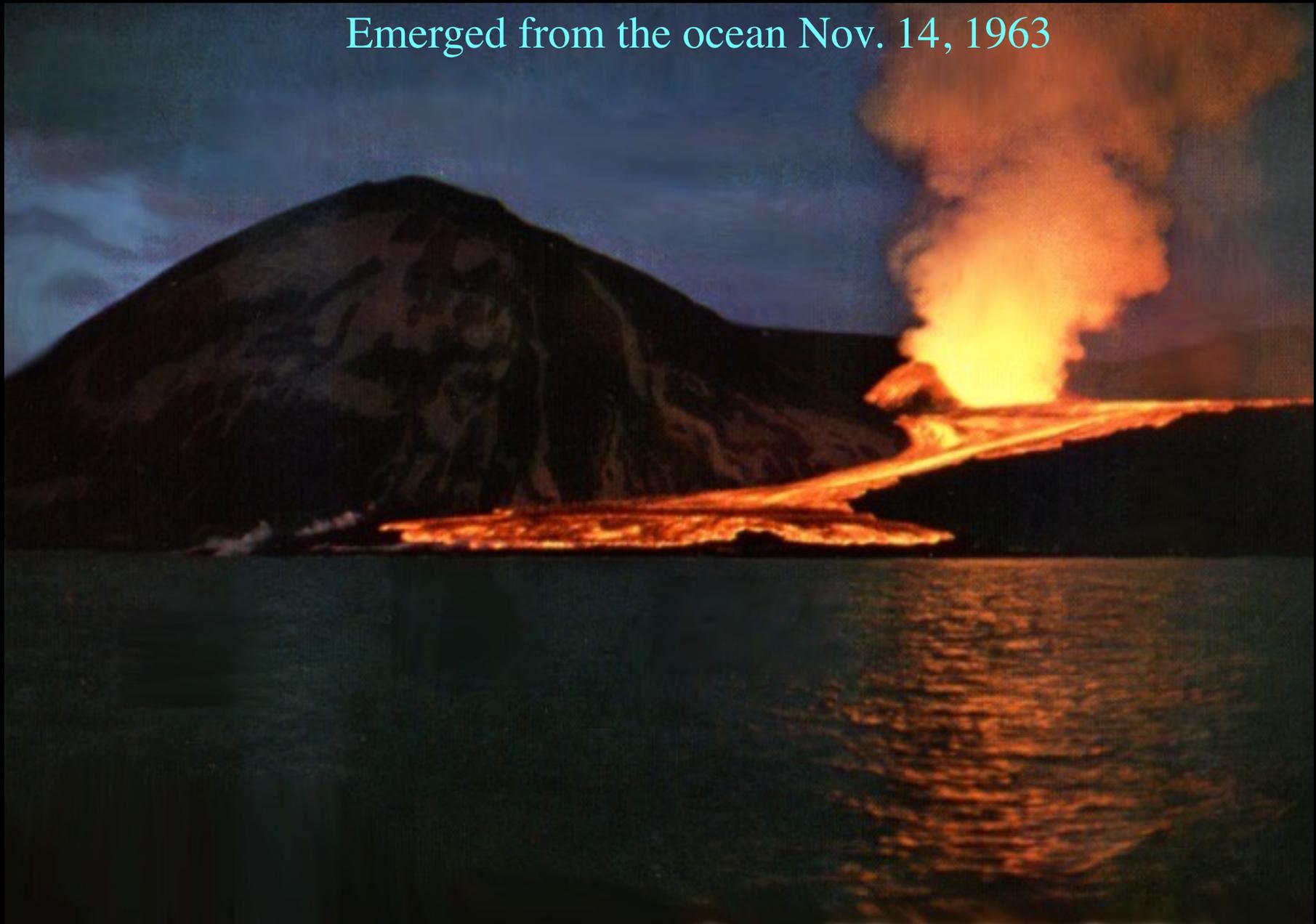


Mt. Fuji

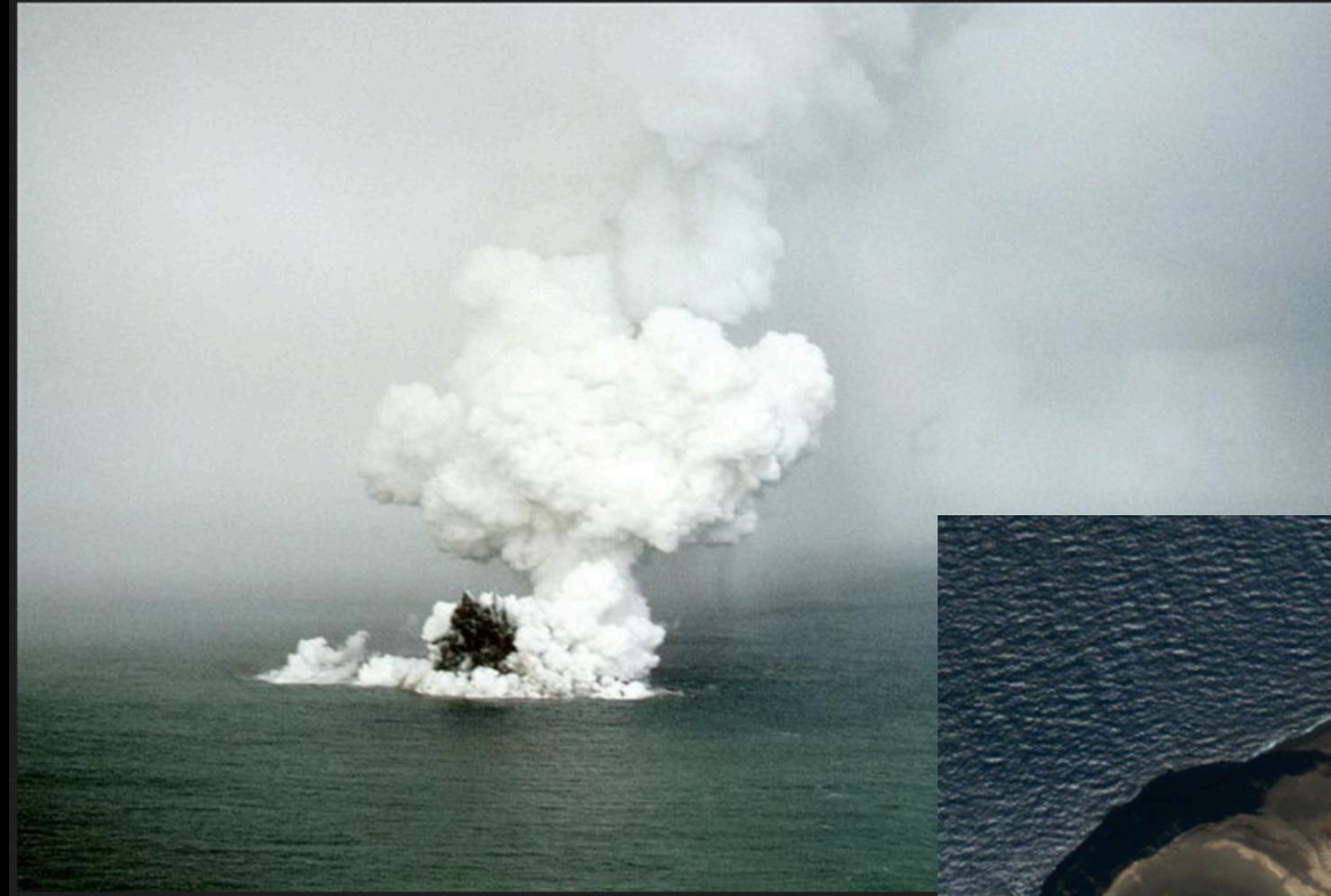


Surtsey, Iceland

Emerged from the ocean Nov. 14, 1963



Surtsey Iceland



Volcanic Craters

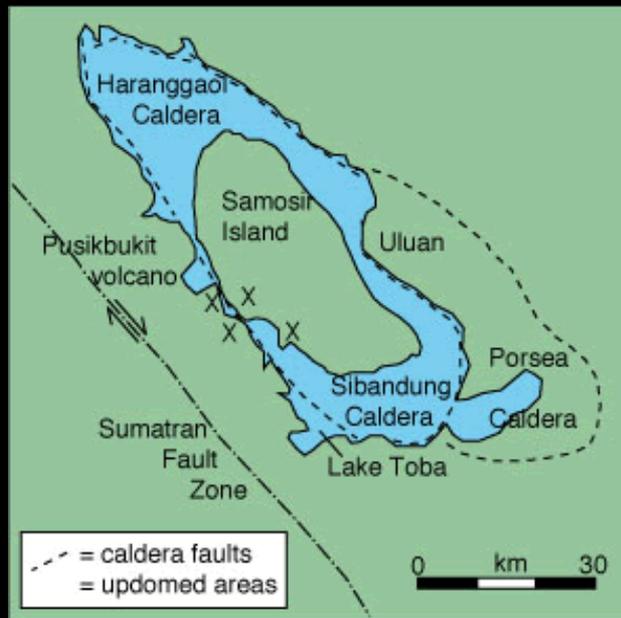
1. Erupted 7,700 years ago creating 50 km³ new volcanic material
2. 10km diameter, No rivers in or out
3. Deepest lake in U.S. at 1943 feet

Crater Lake
Caldera,
Oregon



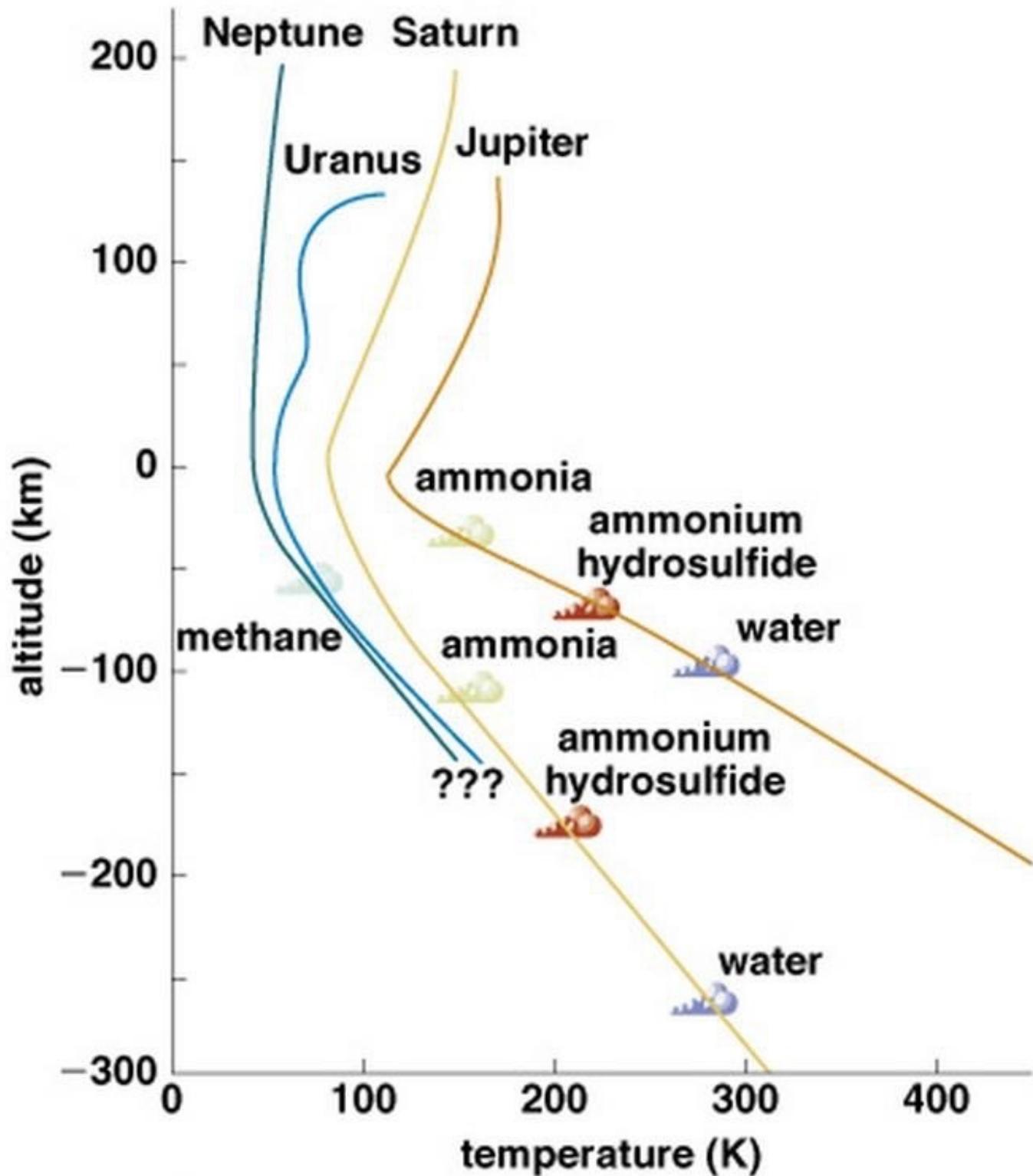
Volcanic Craters

Supervolcano
Caldera,
Toba, Sumatra



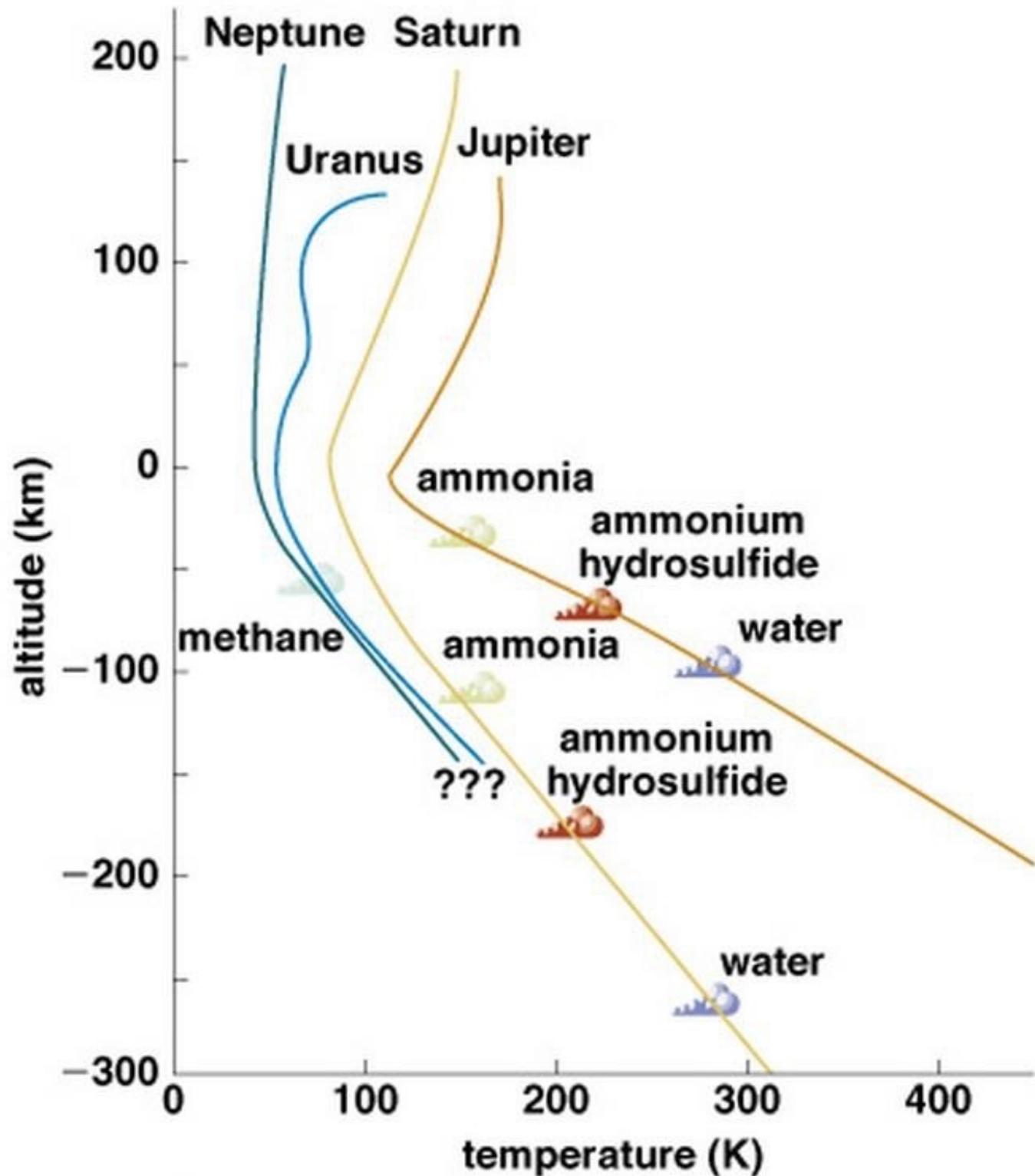
1. Erupted 75,000 year ago creating 2800 km³ new volcanic rock, 800 km³ volcanic ash, cooler temperatures
2. 6 years of Volcanic Winter, Temps down by 10-15 deg.
3. Ash found in India, Lake Malawi in Africa
4. Prolonged deforestation in South Asia
5. Genetic Bottleneck: Evidence shows human population decreased to 3,000-10,000 individuals 50-100kyr ago
6. Similar evidence for chimps, orangutans, cheetahs, tigers

Cloud Decks



Cloud Decks

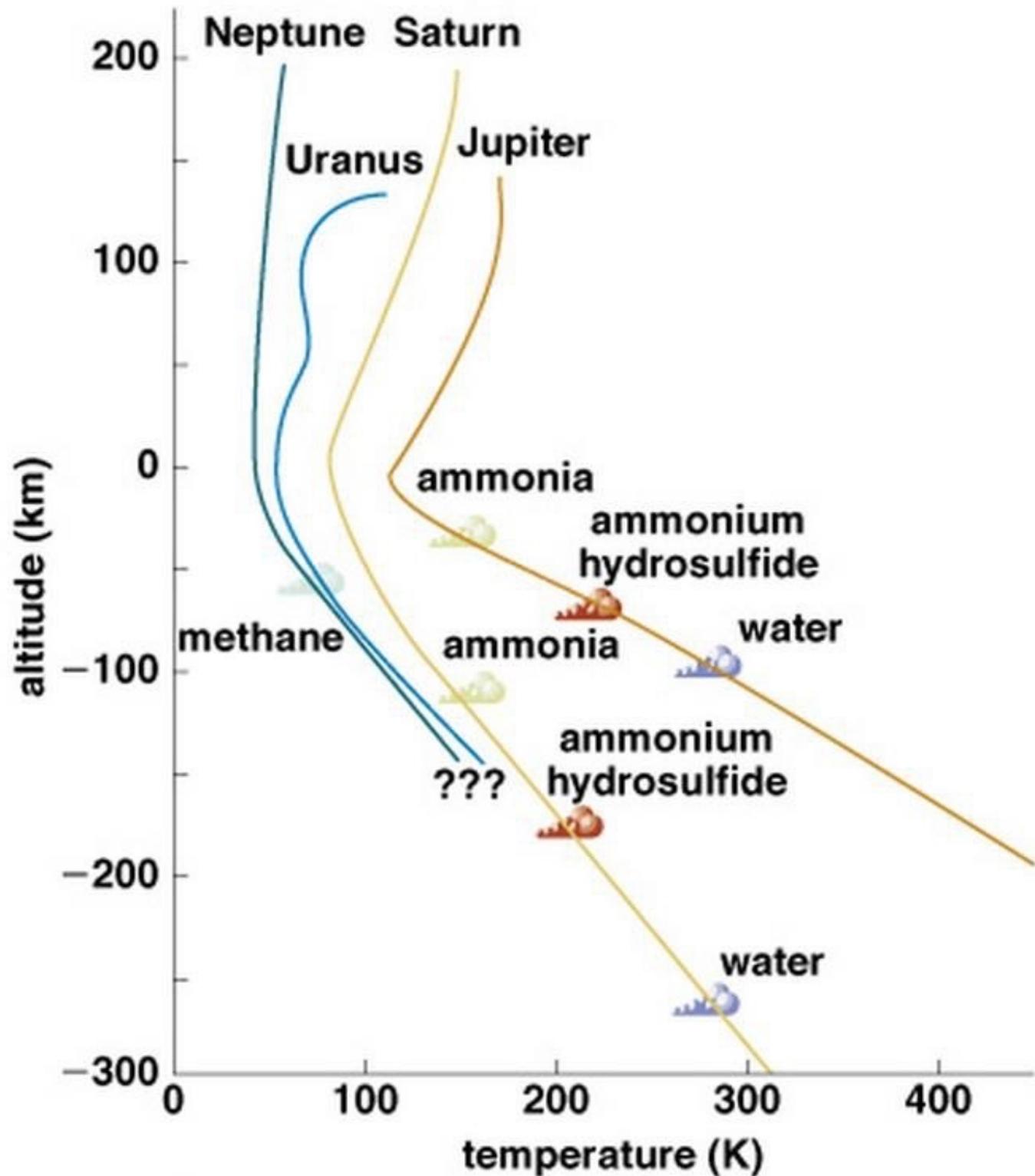
Why no Methane clouds on Jupiter & Saturn?



Cloud Decks

Why no Methane clouds on Jupiter & Saturn? **Too Hot!**

Why are cloud decks separated more on Saturn than Jupiter?



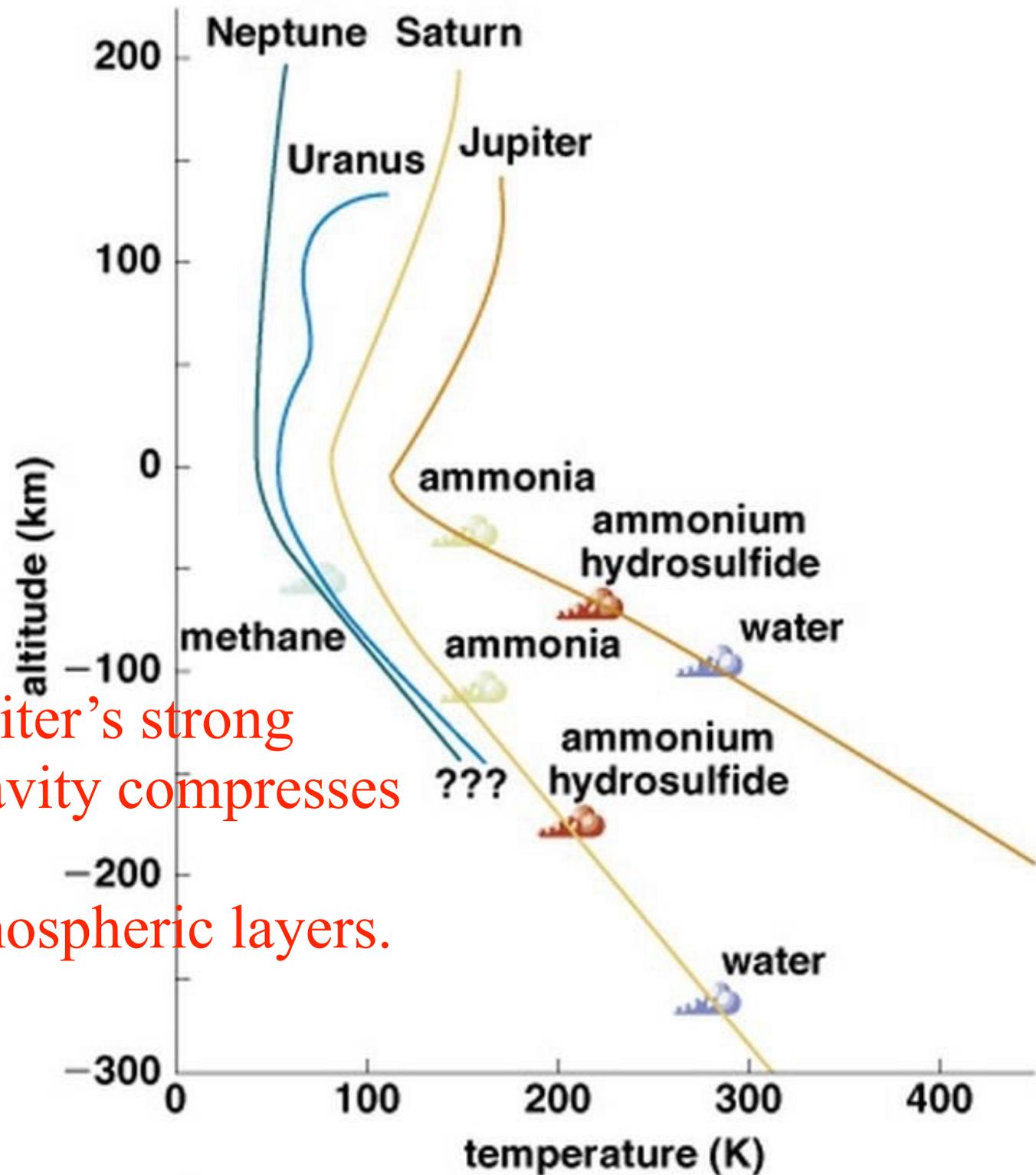
Cloud Decks

Why no Methane clouds on Jupiter & Saturn? Too Hot!

Why are cloud decks separated more on Saturn than Jupiter?

What happens below the water clouds?

Jupiter's strong Gravity compresses the atmospheric layers.



Cloud Decks

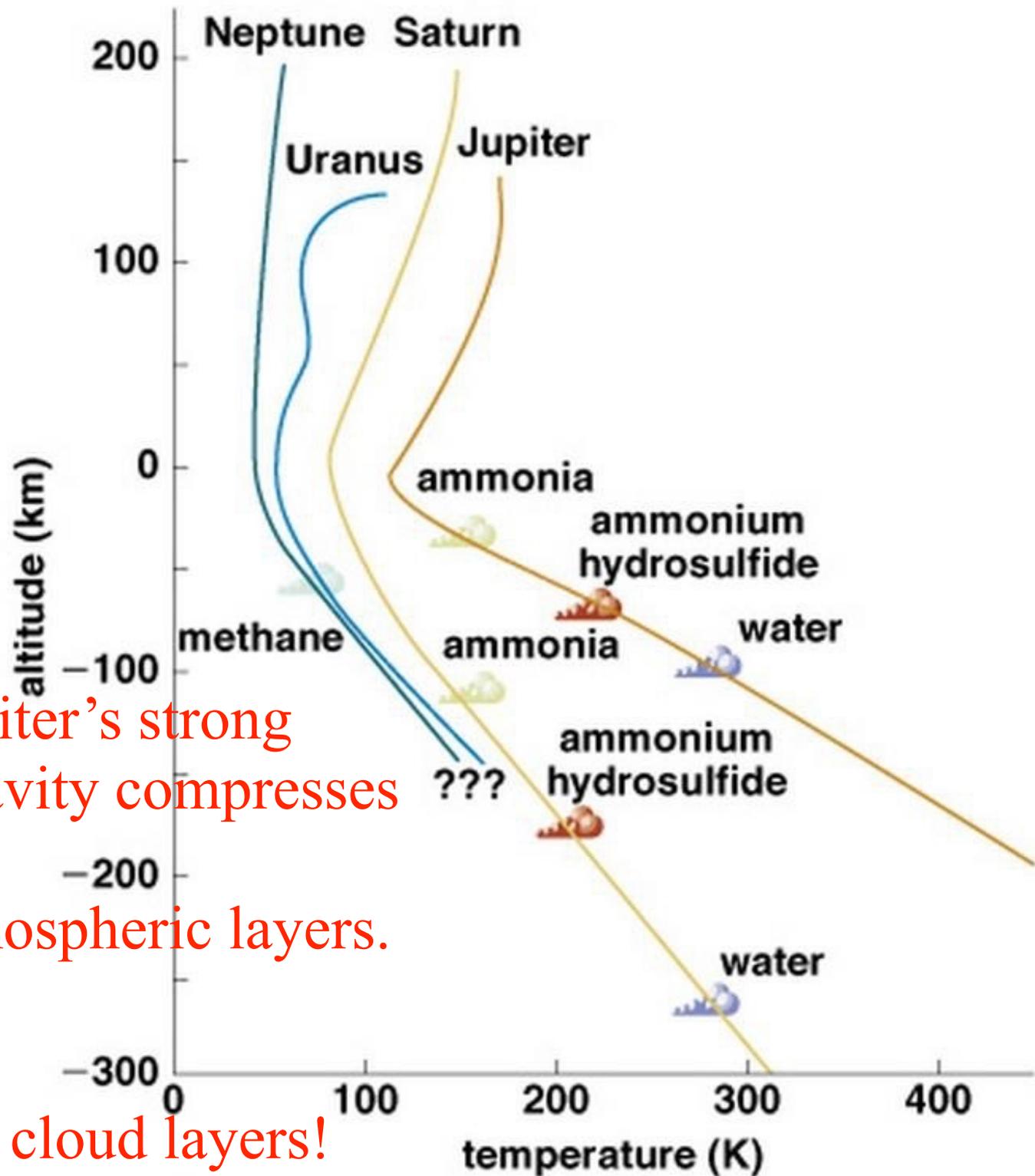
Why no Methane clouds on Jupiter & Saturn? Too Hot!

Why are cloud decks separated more on Saturn than Jupiter?

What happens below the water clouds?

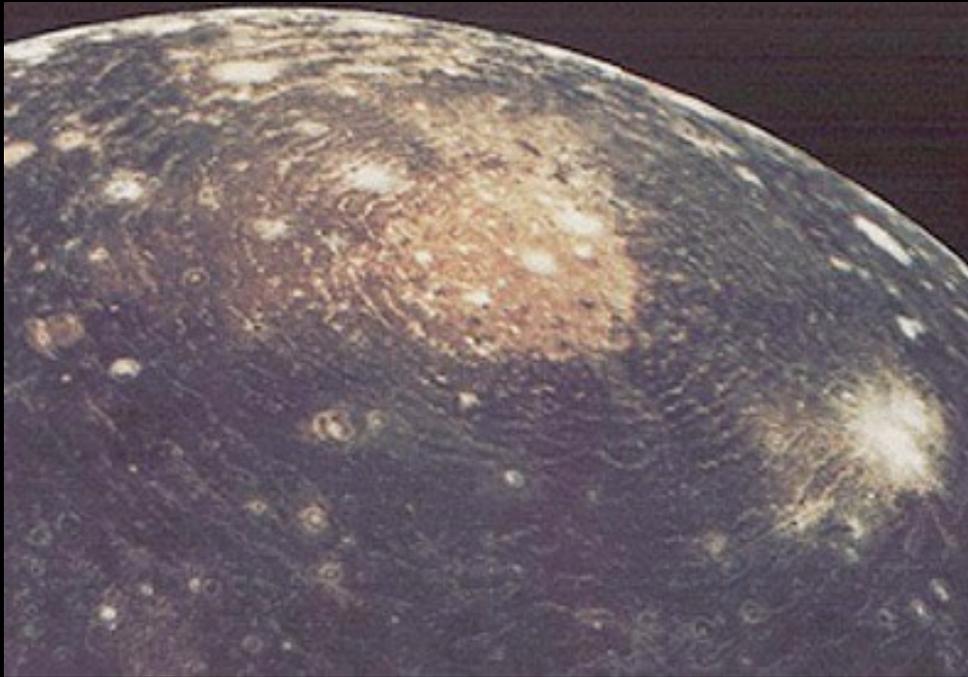
Jupiter's strong Gravity compresses the atmospheric layers.

More cloud layers!



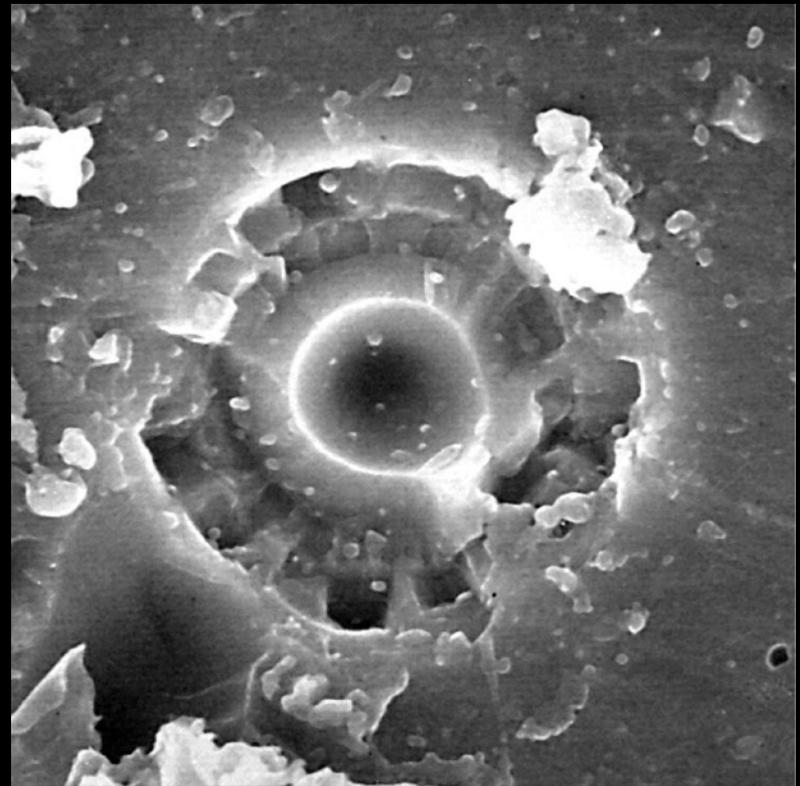
Impact Craters

Big Ones



4000 km-sized (10^7 m) impact structure on Ganymede

Small Ones



10 micron (10^{-5} m) crater in lunar glass

Impacts on the Moon



Front side of
the Moon

Back side of
the Moon



Craters give Relative Surface Ages

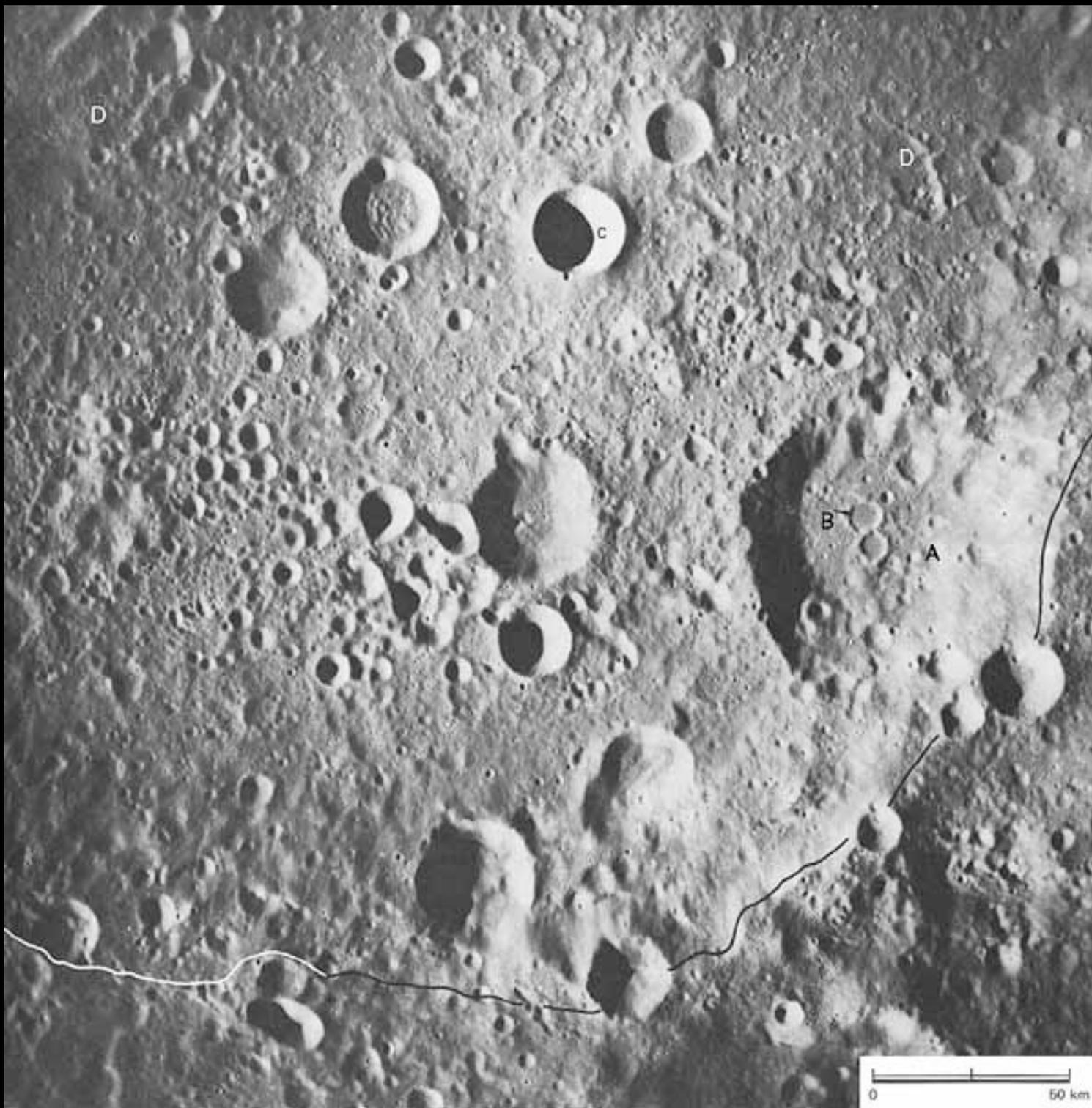


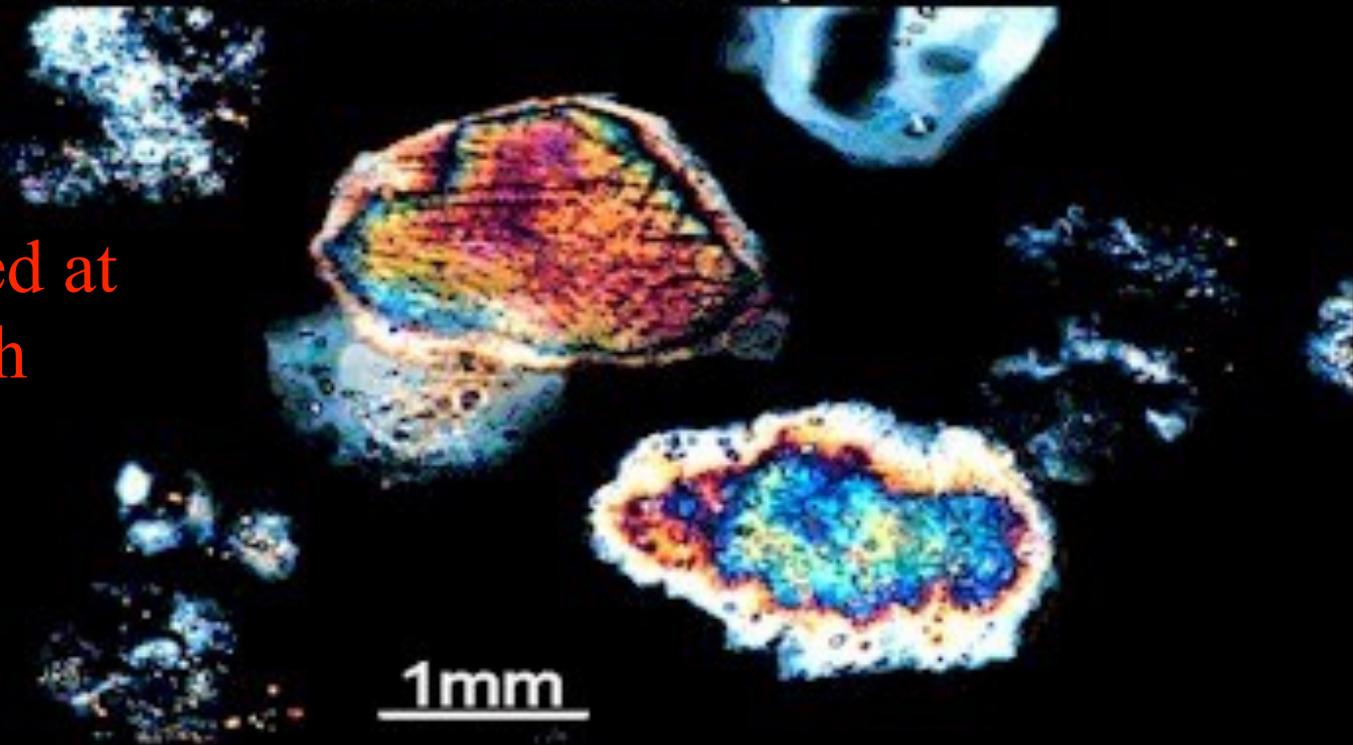
Image of
the Moon

Impacts affect Mineralogy

Impact Craters

Shock lamellae in quartz

Quartz shocked at
extremely high
pressures



(photomicrograph, polarized light, by Sean Macauley)

A Piece of Earth's Mantle

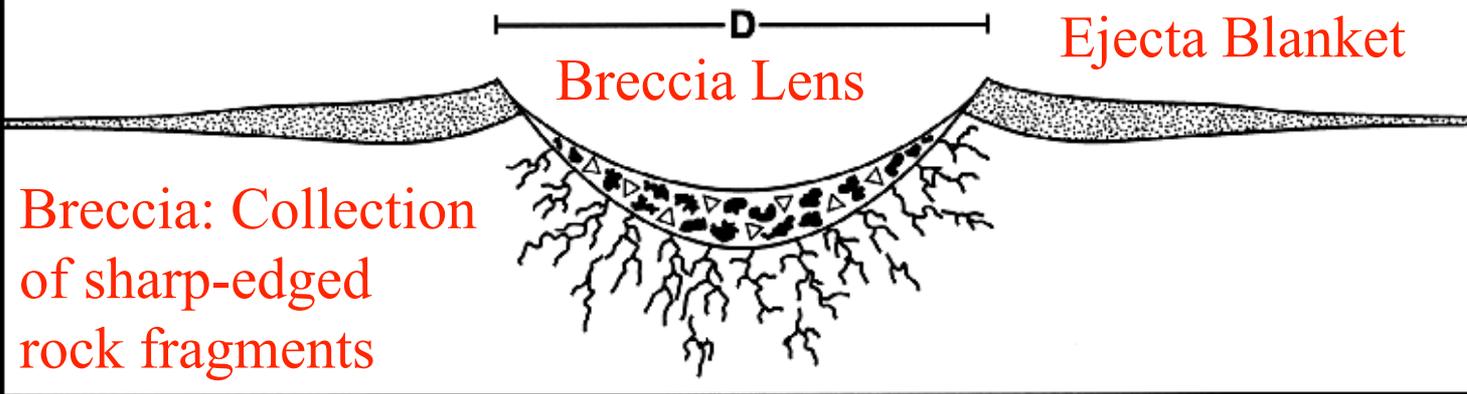
Eclogite - formed at the high pressures found ~100 km deep

These are much lower than the largest impact induced pressures.



Crater Morphology

Simple Crater



△ Breccia

■ Impact melt

■ Impact ejecta

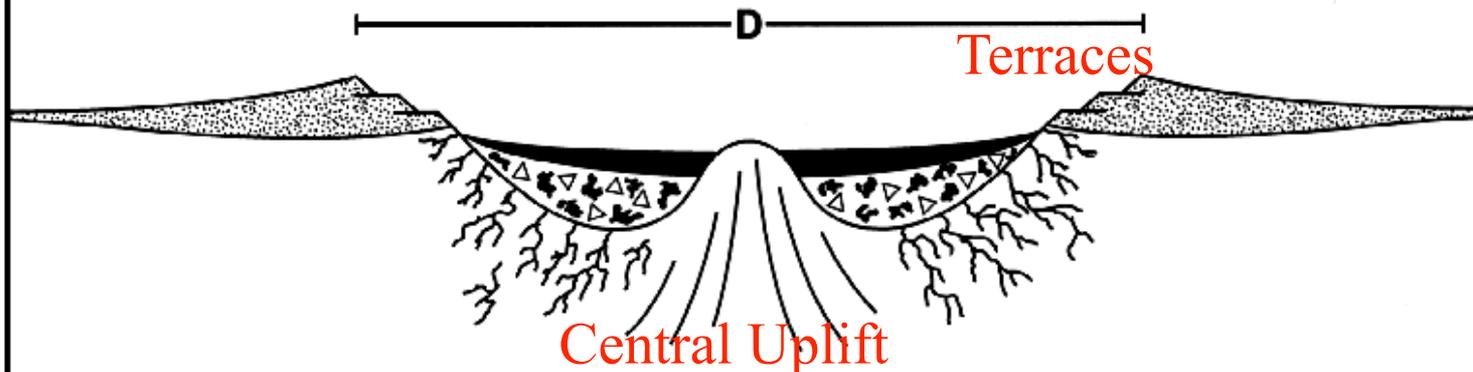


Fractured bedrock



Central peak uplift

Complex Crater



Small Craters on the Moon

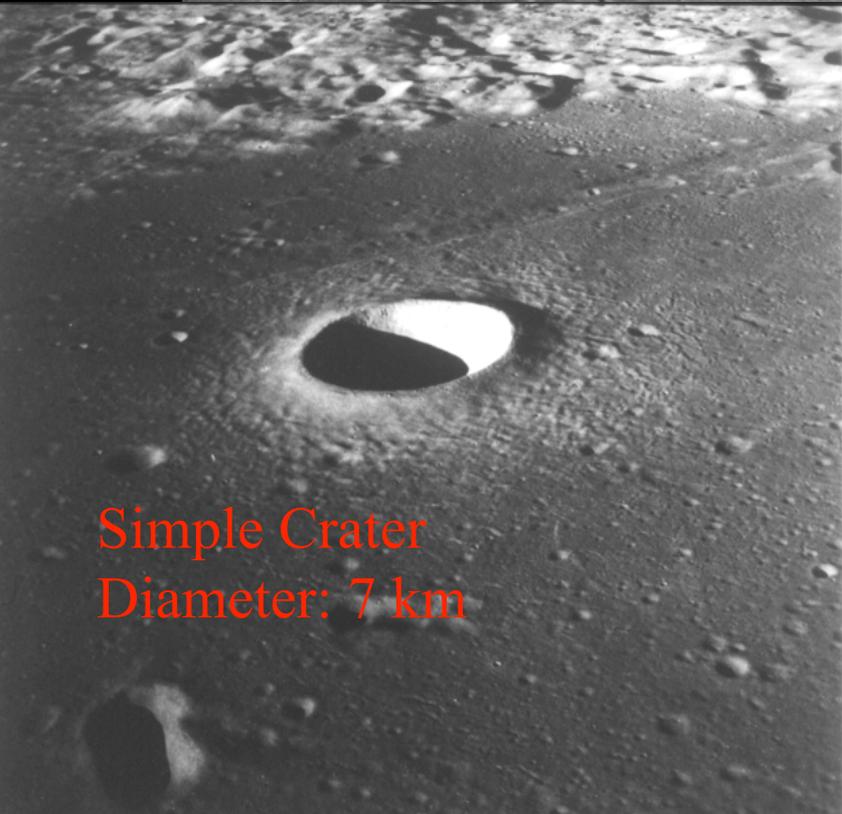
Transition Crater
Diameter: 16 km



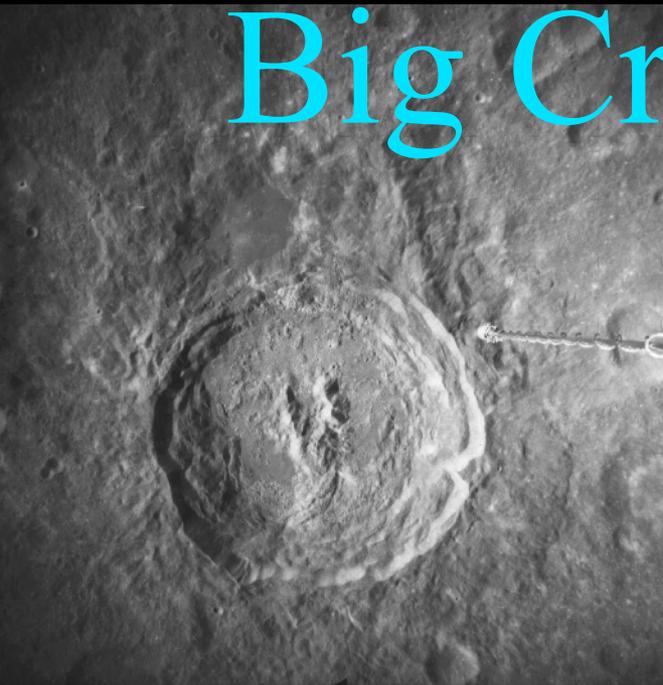
Complex Crater
Diameter: 28 km



Simple Crater
Diameter: 7 km



Big Craters on the Moon



King Crater
Diameter: 77 km



Schrodinger Crater: Diameter: 320 km

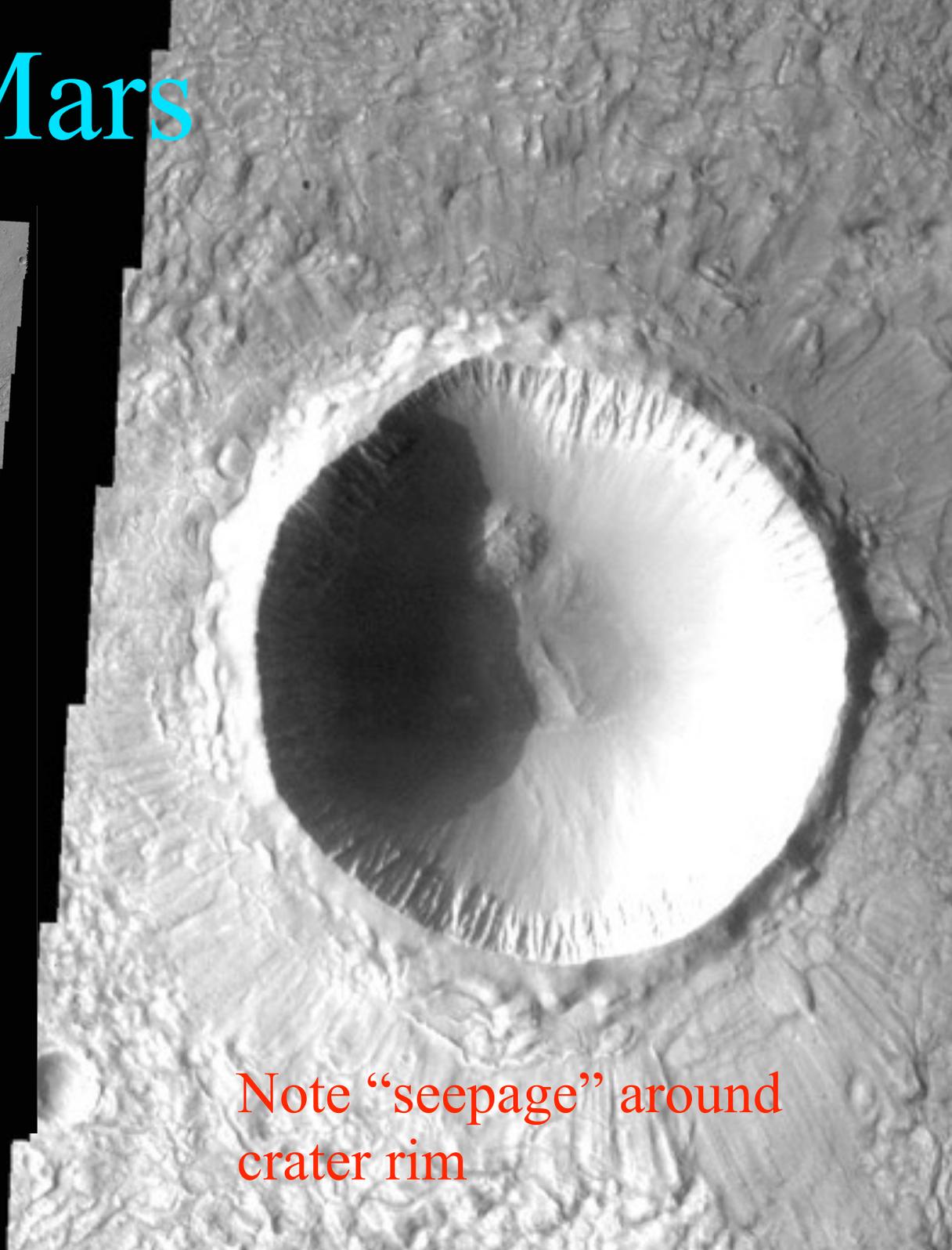


Copernicus Crater
Diameter: 93 km

Craters on Mars



Note “fluidized”
ejecta blanket



Note “seepage” around
crater rim

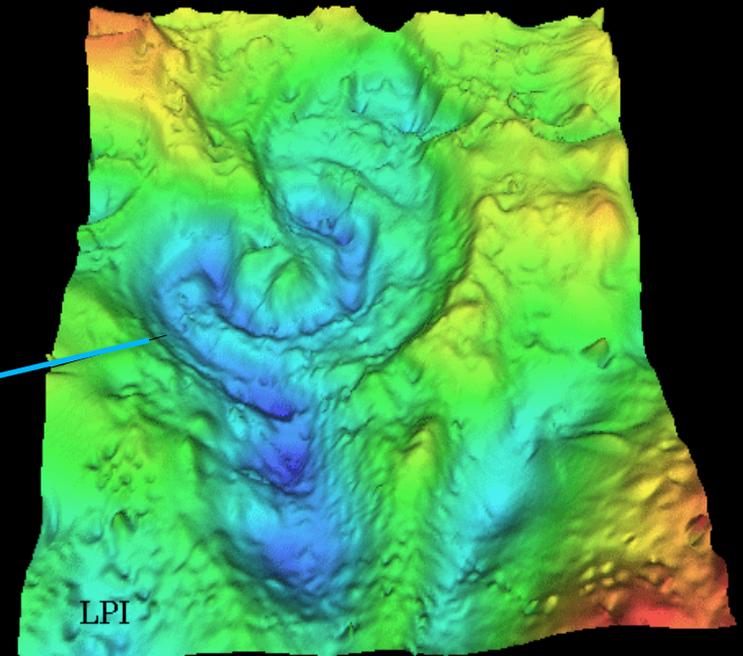
Craters on Venus



Note asymmetric
ejecta blankets

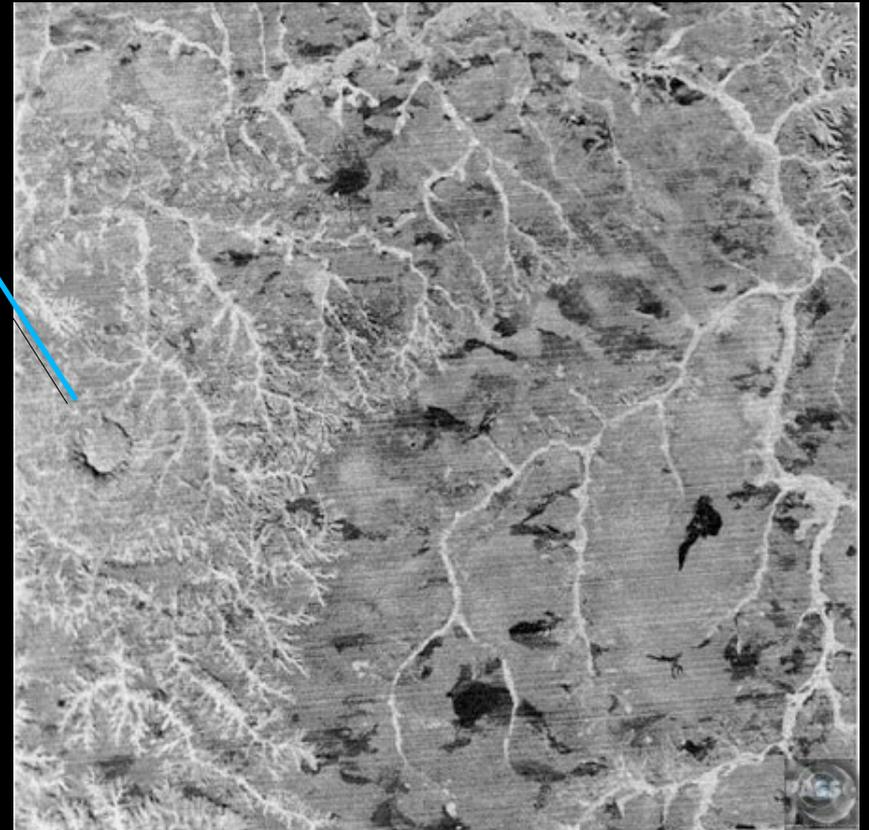
Craters on Earth

Chicxulub, Mexico
200 km diameter
Gravity Data



Craters on Earth

Serra de Cangalha
12 km diameter
Landsat Image



Craters on Earth

Meteor Crater, Arizona
1 km diameter



Craters on Earth



Why so many more in Australia?

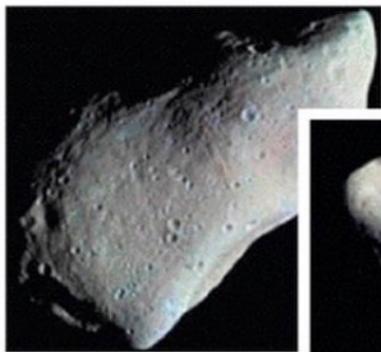
Impact Craters



Wolf Impact Crater, Australia:
Diameter 0.9 km, Age: 0.3Myr

Current Impactor Population

Some Asteroids visited
by spacecraft



Gaspra

(a)



Ida

(b)



Mathilda

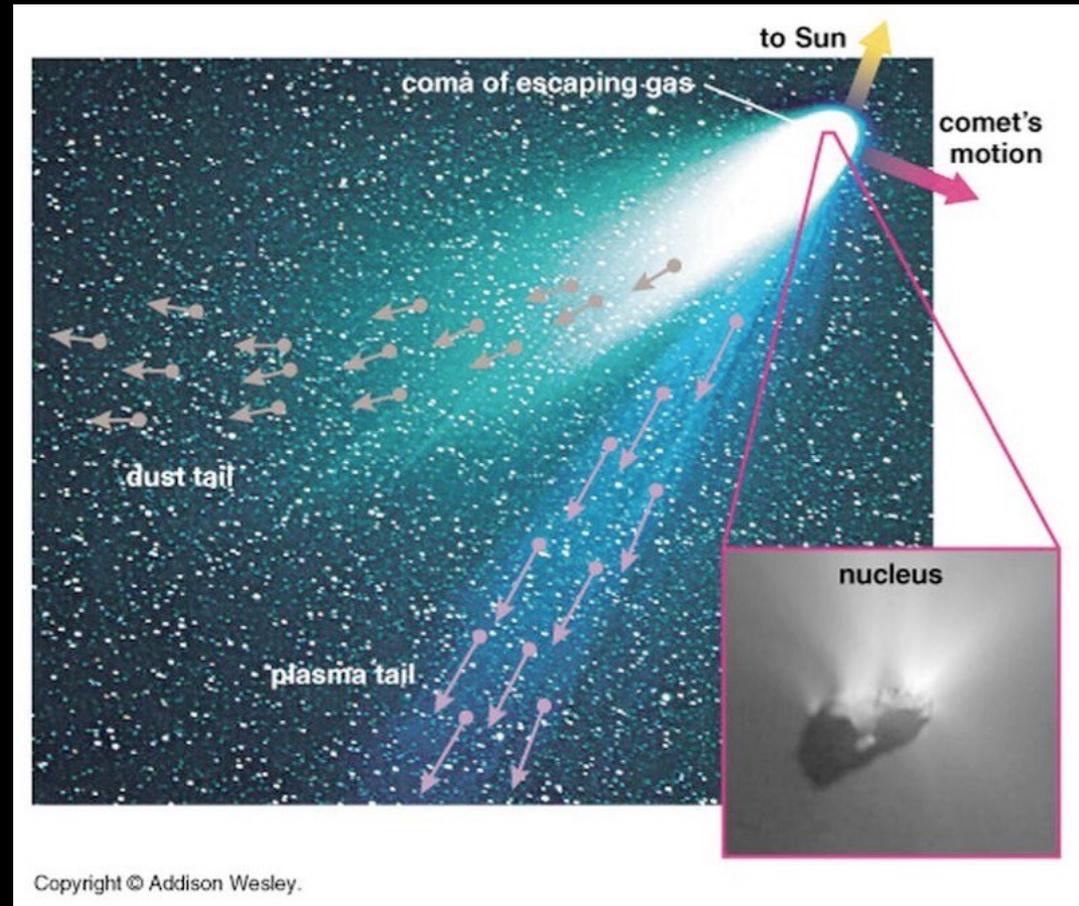
(c)



Eros

(d)

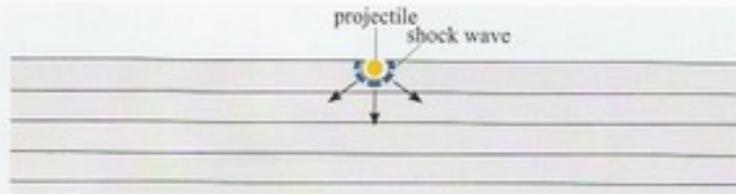
Copyright © Addison Wesley.



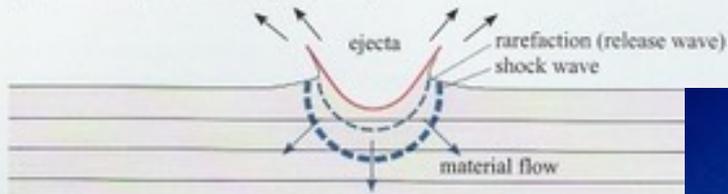
Copyright © Addison Wesley.

Comets - icy objects begin to
melt in the inner Solar System

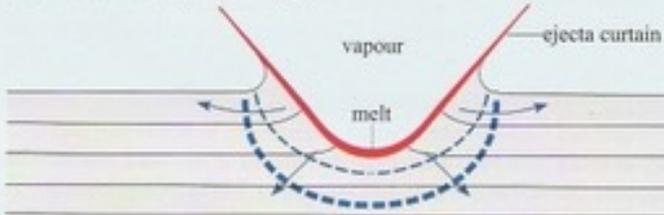
The Cratering Process



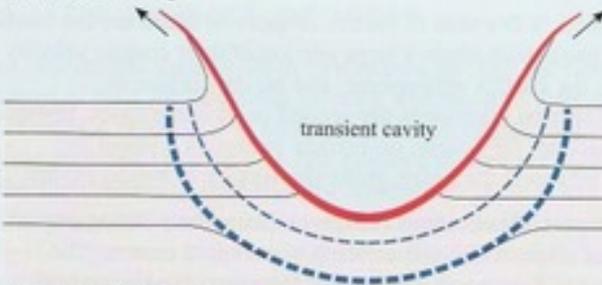
(a) contact target rocks compression stage



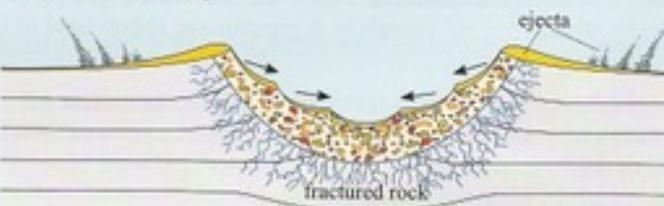
(b) end contact/compression stage



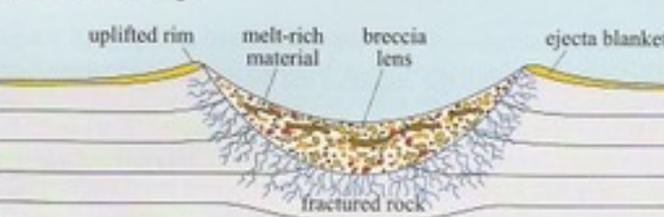
(c) excavation stage



(d) end excavation stage



(e) modification stage



(f) final crater



Three Stages:

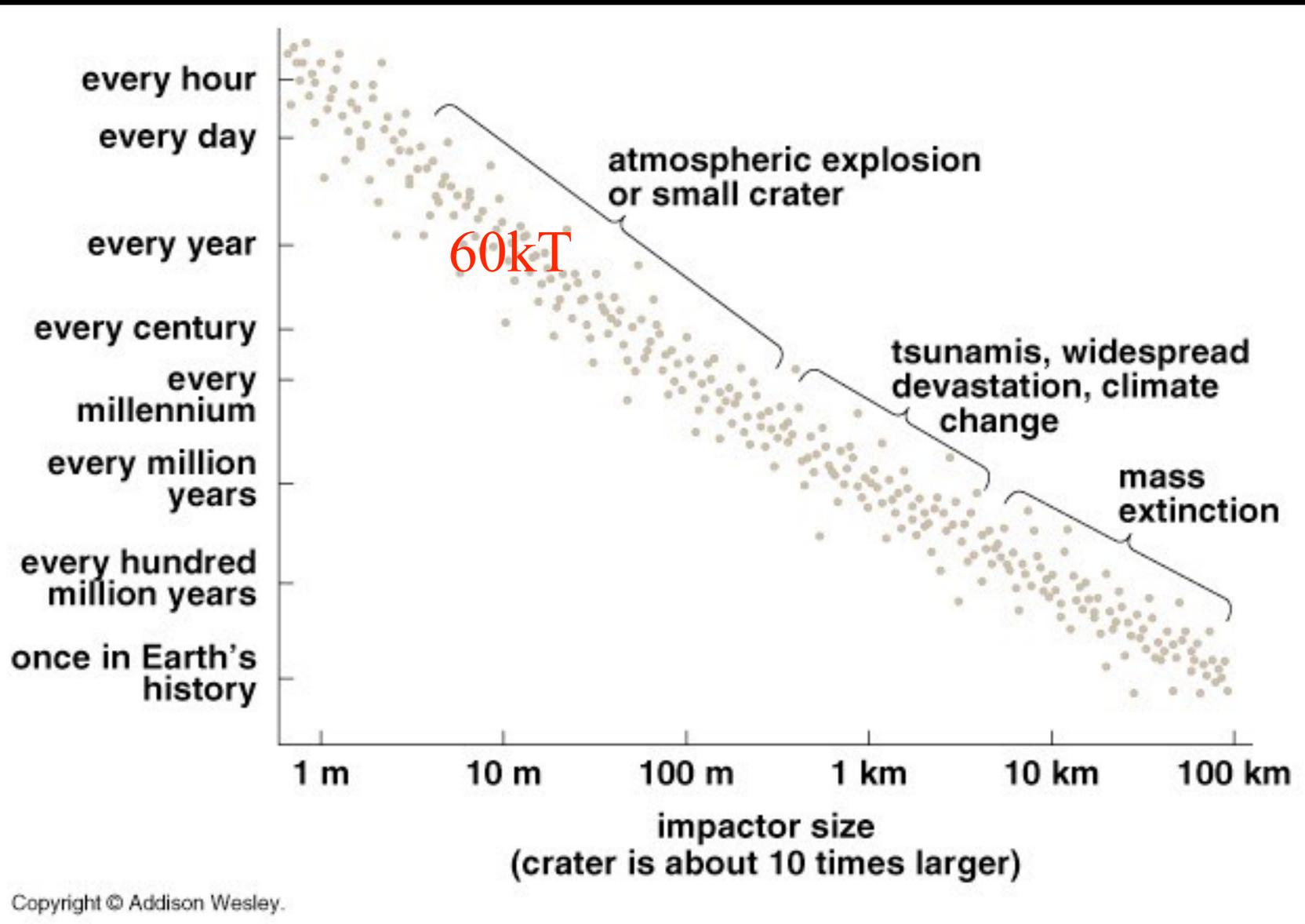
1. Contact & Compression

2. Excavation

3. Modification

NO MIDTERM 1 MATERIAL
BEYOND THIS POINT

Impact Frequency



Small impact happen much more often than large ones

Impactor Populations

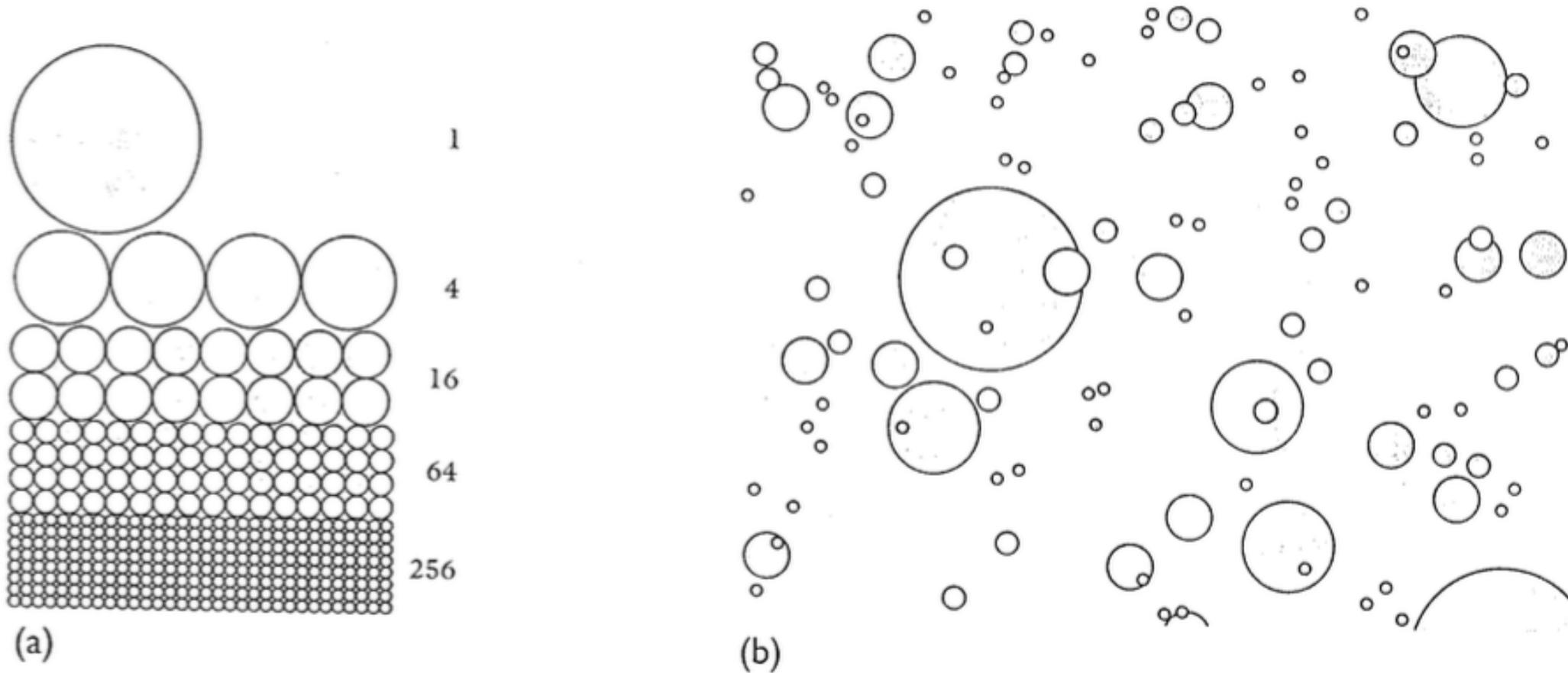


Figure 7.17 (a) The relative number of objects that could strike the Moon, according to a law in which the number of objects increases in inverse proportion to the square of the object's radius. (b) A random distribution of craters made by the population of objects shown in (a).

Estimate the Power of the Ground to stop a Meteoroid!

How much material must the meteoroid interact with to slow by 50%?

From Physics:

Momentum = (mass) (velocity) is conserved.

So velocity will be halved when mass is doubled.

Now assume that all ground in front of the impactor is plastered onto its surface.

So the impactor will penetrate into the ground by a few diameters. Observed: A few diameters.

Estimate the Power of Air to stop a Meteoroid!

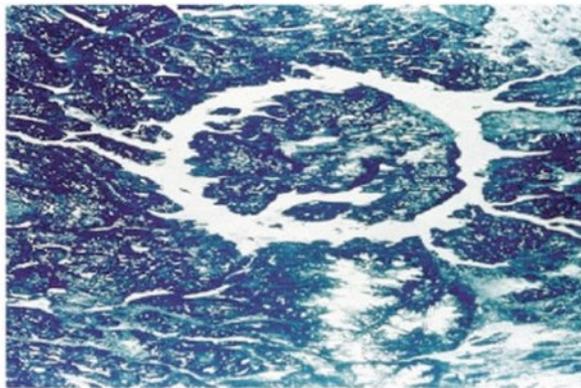
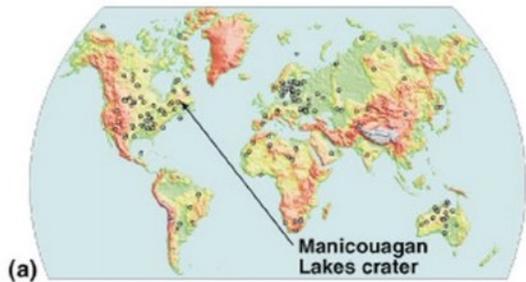
Assume that all air in front of the impactor is plastered onto its surface!

Air is 1/1000 as dense as water and is ~10km thick

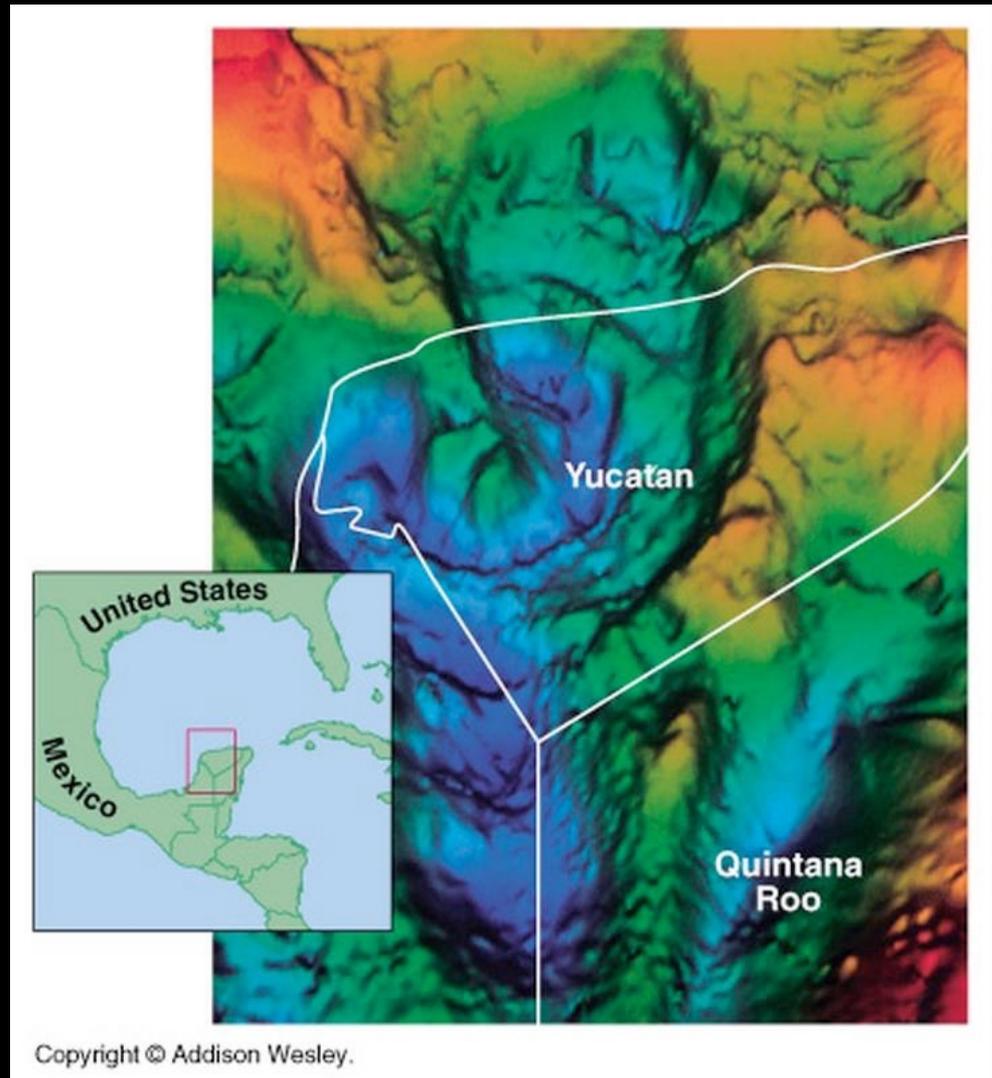
So the impactor will penetrate into the atmosphere by a thousand diameters. So $10\text{km}/1000 = 10\text{m}$ will make it through the atmosphere.

Observed: Must be ~50-100 m to make it through

Big Impact Craters on Earth



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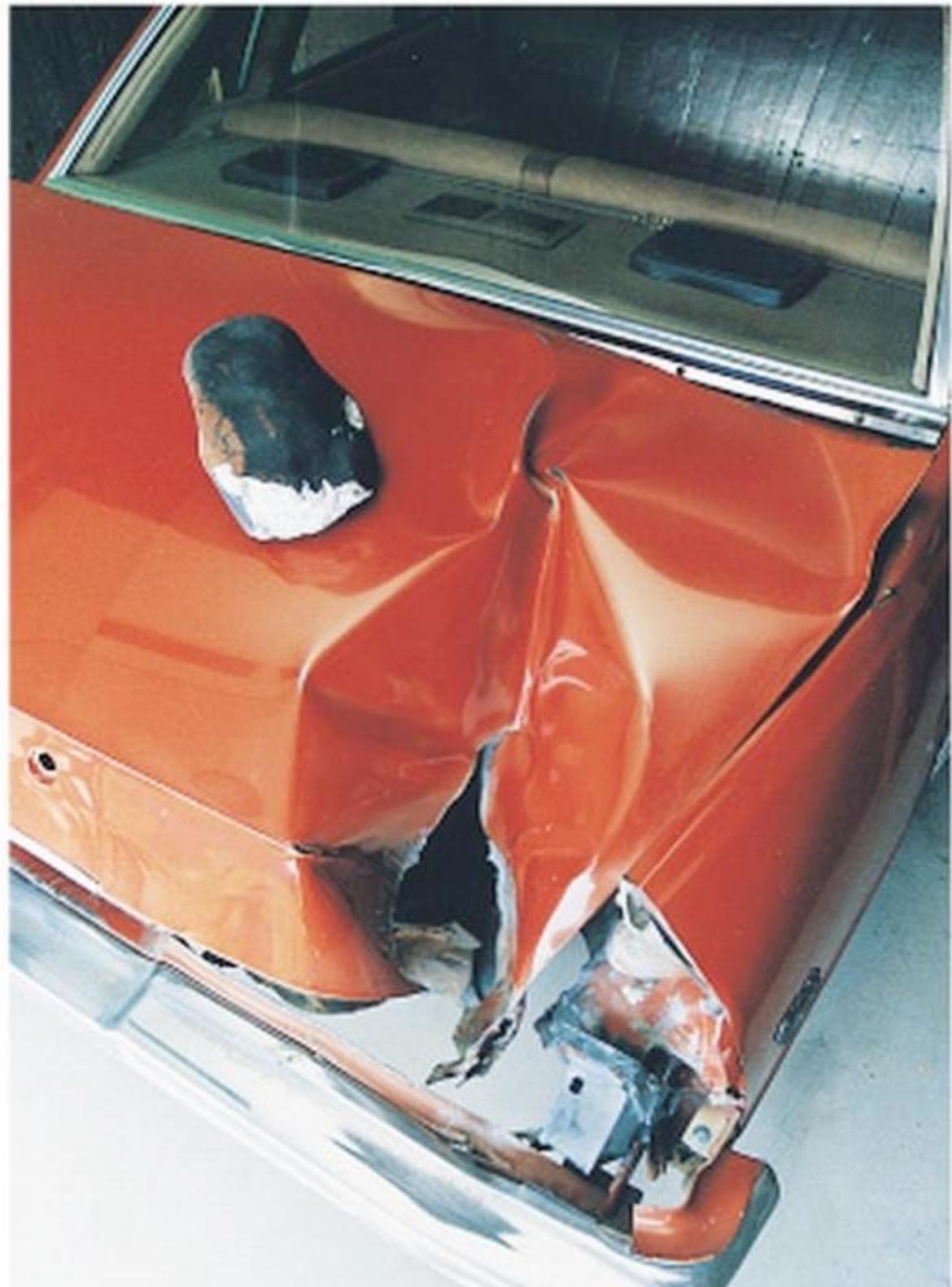


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Very Small Craters on Earth

Is your insurance
up to date?

At what speed was
this impact?



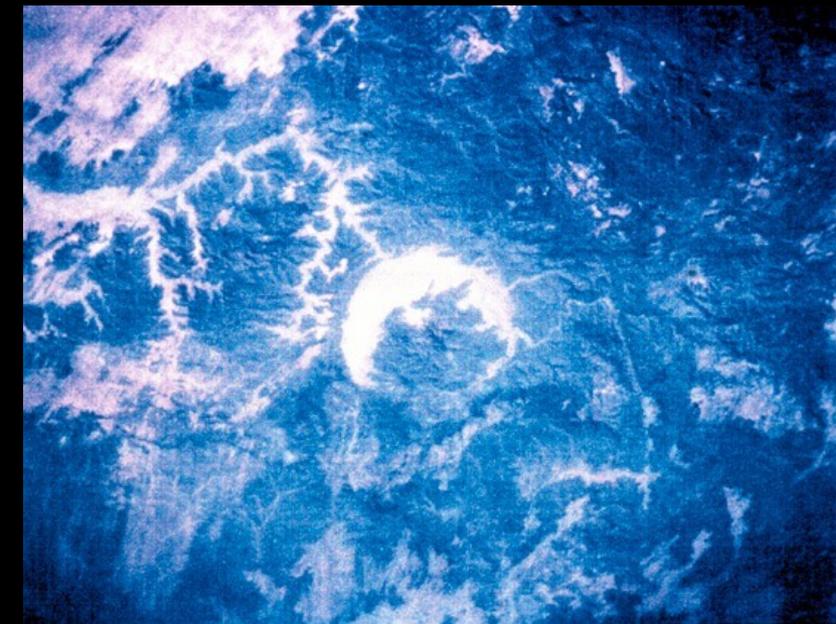


Manicouagan, Canada:
100km, 214Myr

Craters on Earth



Gosses Bluff, Australia:
22km, 142.5Myr



Gwen Fada, Chad:
14km, <350Myr



Aorounga, Chad: 10km, <350Myr

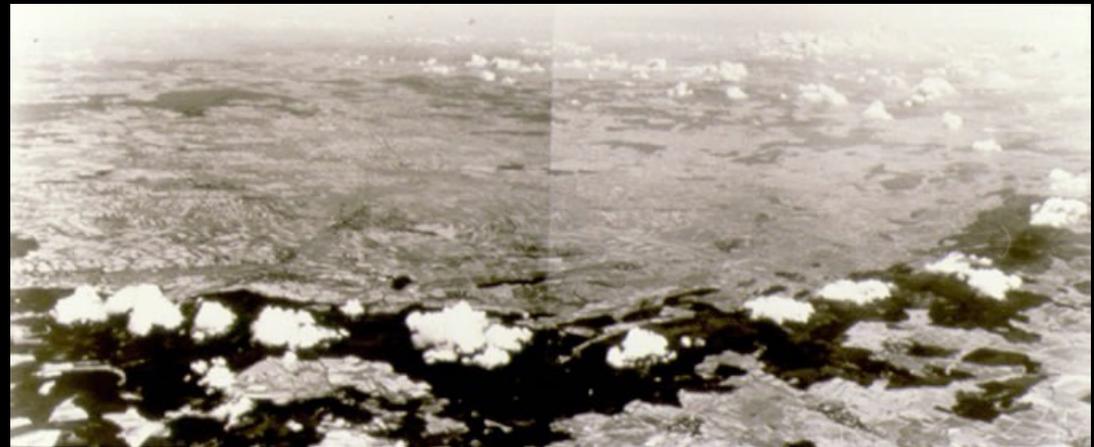


Clearwater Lakes, Canada
26km, 290Myr



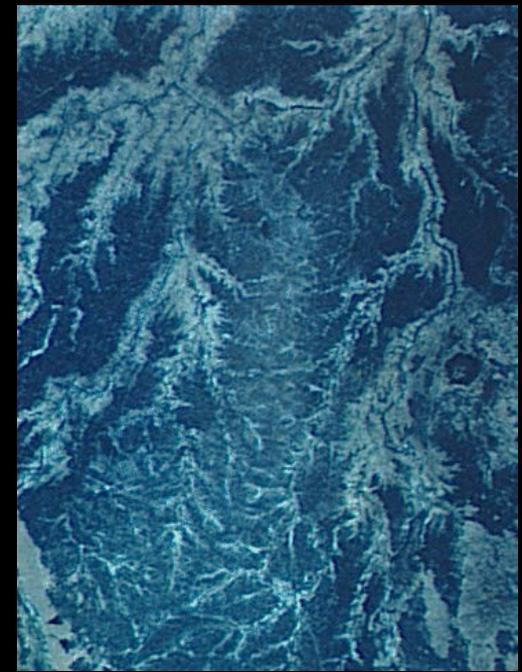
Space Shuttle!

Ries, Germany
24km, 15Myr





Mistastin Lake, Canada:
28km, 38Myr



Ramgarh, India: 5.5km, unknown



Deep Bay, Canada
5km, 100 +/- 50 Myr



Ouarkziz, Algeria: 4km, <70Myr



Roter Kamm, Namibia:
2.5km, 3.7Myr



Meteor Crater, Arizona: 1.2km, 49,000 yr

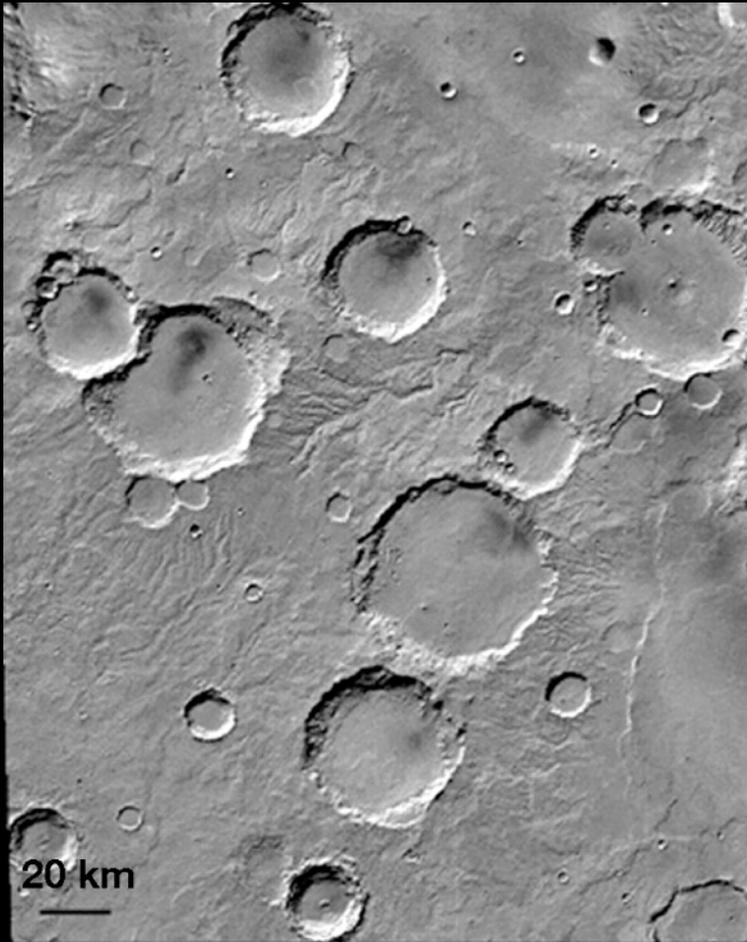


Wolf Creek, Australia: 850m, 0.3Myr



Goat Paddock, Australia: 5km, <55Myr

Craters on Mars



Viking Image, Mars
southern hemisphere

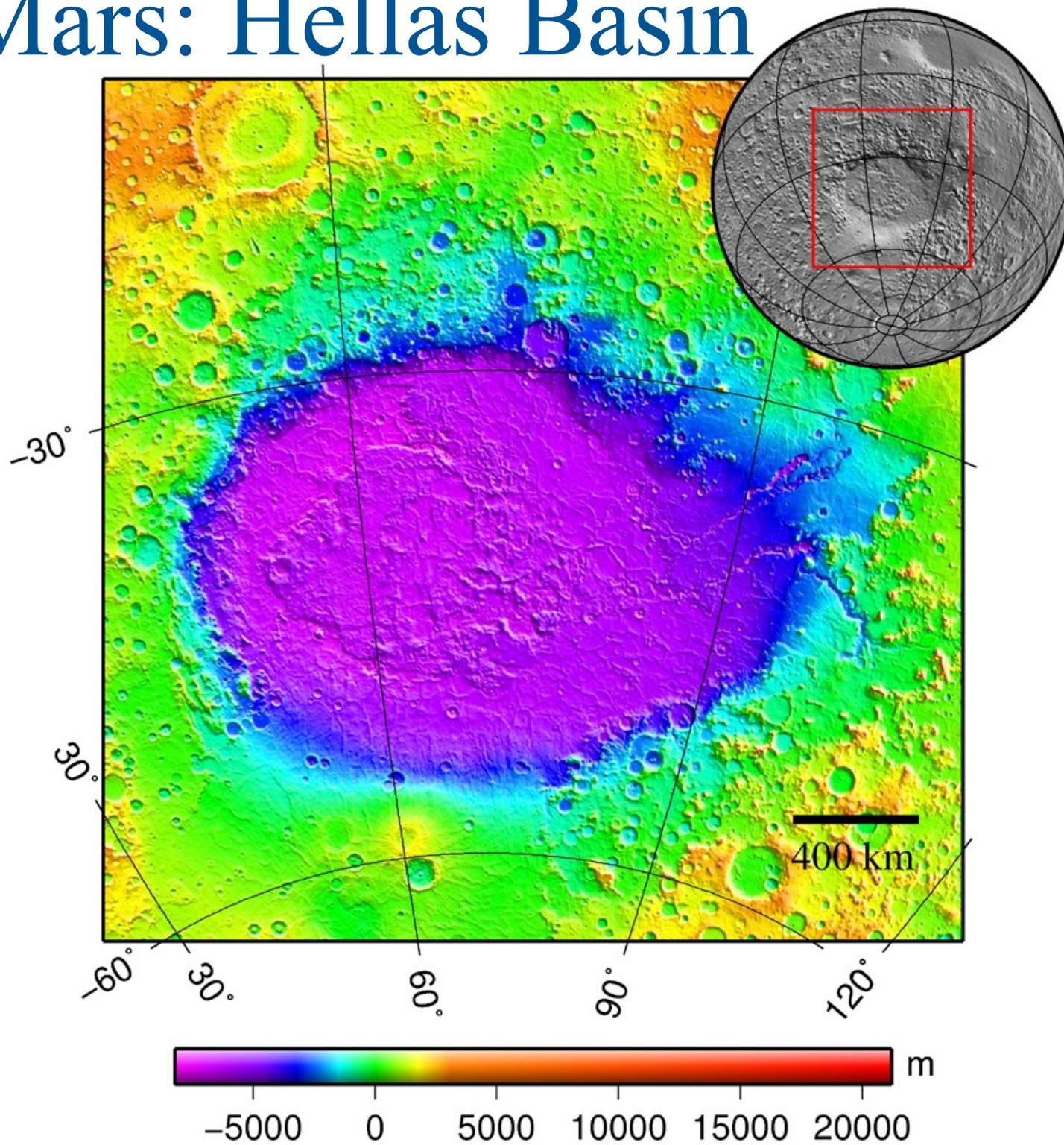


Victoria Crater, Mars: 750m



Mars, multiple strikes: 78km x 25km

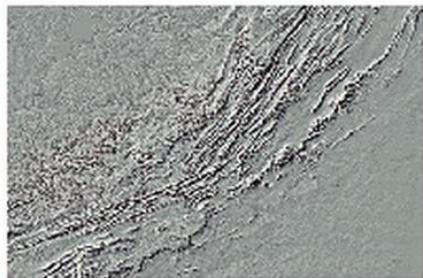
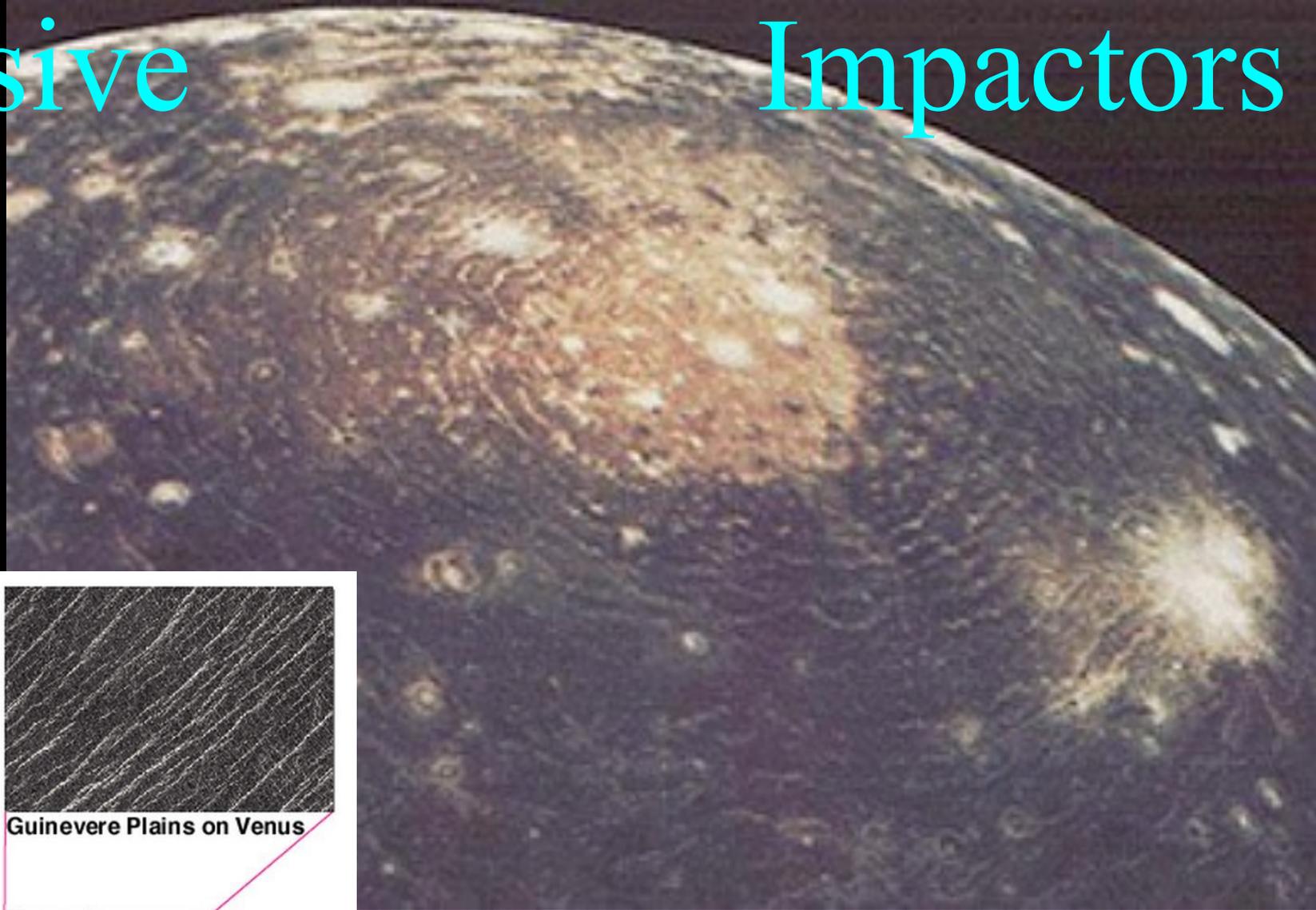
Mars: Hellas Basin



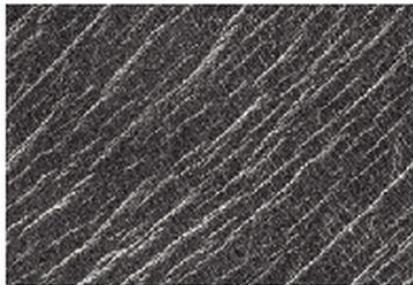
Massive

Impactors

Ganymede



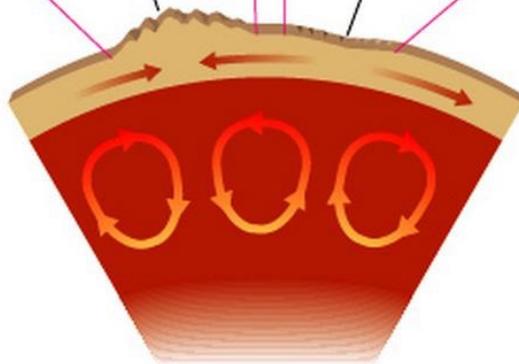
Appalachian Mountains in eastern United States



Guinevere Plains on Venus

Compression in crust can make mountains.

Extension can make cracks and valleys.



Circular features in huge multi-ringed basins are caused by an impactor that penetrates the lithosphere and sets up currents in the mantle.

Last Slide